GROUND WATER RESOURCES

Jefferson County

Open File Report 79-51 WRD

GROUND-WATER RESOURCES OF JEFFERSON COUNTY

Introduction

Approximately 67 percent of the residents of Jefferson County rely on ground water as the source of their drinking water. It is estimated that the use of ground water in the county currently approaches 0.7 billion gallons per year. For comparison, this amount would provide each resident with 118 gallons of water a day during a year. Actually, few if any households use this much water, and the rather large annual per capita use reflects the greater water requirements of the county's industries, agribusinesses, and municipalities.

The users of ground water in the county draw their supplies from several different geologic sources. Several factors must be considered in determining the availability of ground water and the adequacy of a supply source:

- distribution - having water where it is needed,
- accessbility - affects the costs for drilling wells and pumping water,
- yield - relates to the magnitude of the supply that can be sustained,
- quality - determines for what purposes the water can be used.

In terms of these factors, there are few locations in Jefferson County where the availability of ground water is not limited to some degree. The most common limitation is poor water quality, that is, highly mineralized ground water. Secondary limitations are generally related to poor distribution, small yields from some sources, and poor accessibility due to the great depths to adequate sources.

Occurrence of Ground Water in Jefferson County

The occurrence of ground water is influenced by geology -- the position and thickness of the rock units, their ability to store and transmit water, and their physical and chemical make-up. Geologic units that store and transmit water and yield appreciable amounts to wells are called aquifers. The best aquifers are usually composed of unconsolidated sand and gravel, porous sandstone, and porous or fractured limestone and dolomite. Other units with materials such as clay and silt, shale, siltstone, and mudstone yield little or no water to wells. These impermeable units are called aquicludes or aquitards and commonly separate one aquifer unit from another.

In Jefferson County there are four principal aquifers from which users obtain water supplies. The loose, unconsolidated materials near the land surface comprise the surficial aquifer. Below this there are three major rock aquifers--the Mississippian aquifer, the Devonian aquifer, and the Cambro-Ordovician aquifer. Figure 1 shows the geologic relations of these beneath the county. Each of the aquifers has its own set of geologic, hydrologic, and water quality characteristics which determine the amount and potability (suitability for drinking) of water it will yield.

Surficial Aquifers

Unconsolidated deposits at the land surface are comprised of mixtures of clay, silt, sand, gravel, and assorted boulders. Water-yielding potential of the surficial deposits is greatest in units composed mostly of sand and/or gravel. Three types of surficial aquifers are used: the alluvial aquifer, the drift aquifer, and the buried channel aquifer.

The alluvial aquifer consists mainly of the sand and gravel transported and deposited by modern streams and makes up the floodplains and terraces in major valleys. Alluvial deposits are shallow, generally less than 50-60 feet and thus may be easily contaminated by the infiltration of surface water.

The drift aquifer is the thick layer of clay to boulder size material deposited over the bedrock by glacial ice which invaded the county at lease twice in the last two million years. The composition of the glacial drift varies considerably and in many places does not yield much water. There are, however, lenses or beds of sand and gravel within the drift which are thick and widespread enough to serve as dependable water sources. These lenses are difficult to locate because they are irregular in shape and buried within the drift deposits. Usually one or two sand layers can be found in most places that will yield minimum water supplies for domestic wells.

The buried channel aquifer consists of stream alluvium that partially filled valleys that existed before the glacial period. The valleys were overridden by the glaciers and are now buried under glacial and recent alluvial deposits.

The distribution, yields, and water quality characteristics for the surficial aquifers are summarized in Figures 2, 9, and 13. An indication of accessibility can be obtained by comparing the elevations of the top (the land surface) and the bottom (the bedrock surface) of the surficial deposits from Figures 4 and 5. The thickness of the glacial drift and the depth of the buried channels are determined by subtracting the elevations at selected locations.

Rock Aquifers

Below the surficial materials is a thick sequence of layered rocks formed from deposits of rivers and shallow seas that have covered the state

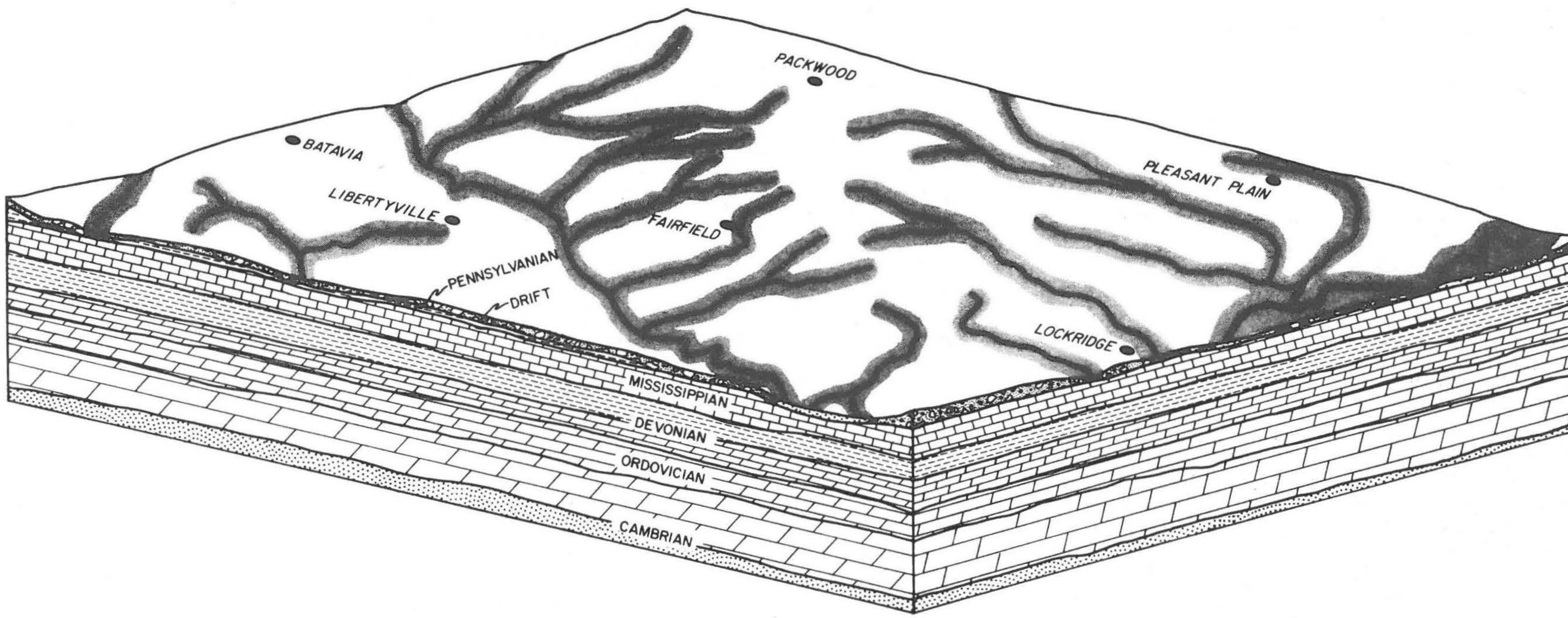


Figure 1

BLOCK DIAGRAM SHOWING THE GEOLOGY OF DALLAS COUNTY within the last 600 million years. The geologic map (Figure 3) shows the geologic units which form the top of this rock sequence. These rocks are Pennsylvanian in age and are mainly shales. Although the Pennsylvanian rocks usually act as an aquiclude, there are locally sandstone layers (particularly in the western portions of the county) which supply small yields to domestic wells. The thickness of the Pennsylvanian rocks varies from about 150 feet in the vicinity of Batavia to 0 feet in the eastern part of the county.

Underlying the Pennsylvanian aquiclude is a series of older rocks, parts of which form the three major rock aquifers in Jefferson County. This sequence and the water-bearing characteristics of the aquifers and aquicludes are shown in Table 1.

Examples of the sequence of rock units encountered in drilling existing wells at various locations in Jefferson County are indexed and illustrated in Figures 7 and 8. The geologic unit that supplies ground water and the amount of water yielded to the well are shown next to each of the well logs.

The accessibility of ground water in the rock aquifers depends first on the depth to the aquifer. The deeper a well must be, the greater the cost for well construction and pumping. The depths to and thicknesses of units at specific sites will vary somewhat because of irregularities in the elevation of the land surface and in the elevation of the tops of the underlying rock units. Estimates of depths and thicknesses can be made by comparing Figure 4 with the maps of aquifer elevations in Figures 10, 11, and 12. The range in depth below land surface to the top of the county's principal bedrock aquifers is given for each township in Figure 6.

The second factor which affects accessibility is the level to which the water will rise in the well (the static water level). Since water in the rock aquifers is under artesian pressure, the water rises in the well once it penetrates the aquifer. This rise in water level can reduce the cost of pumping. Average static water levels in Jefferson County wells are shown in Figures 10, 11, and 12.

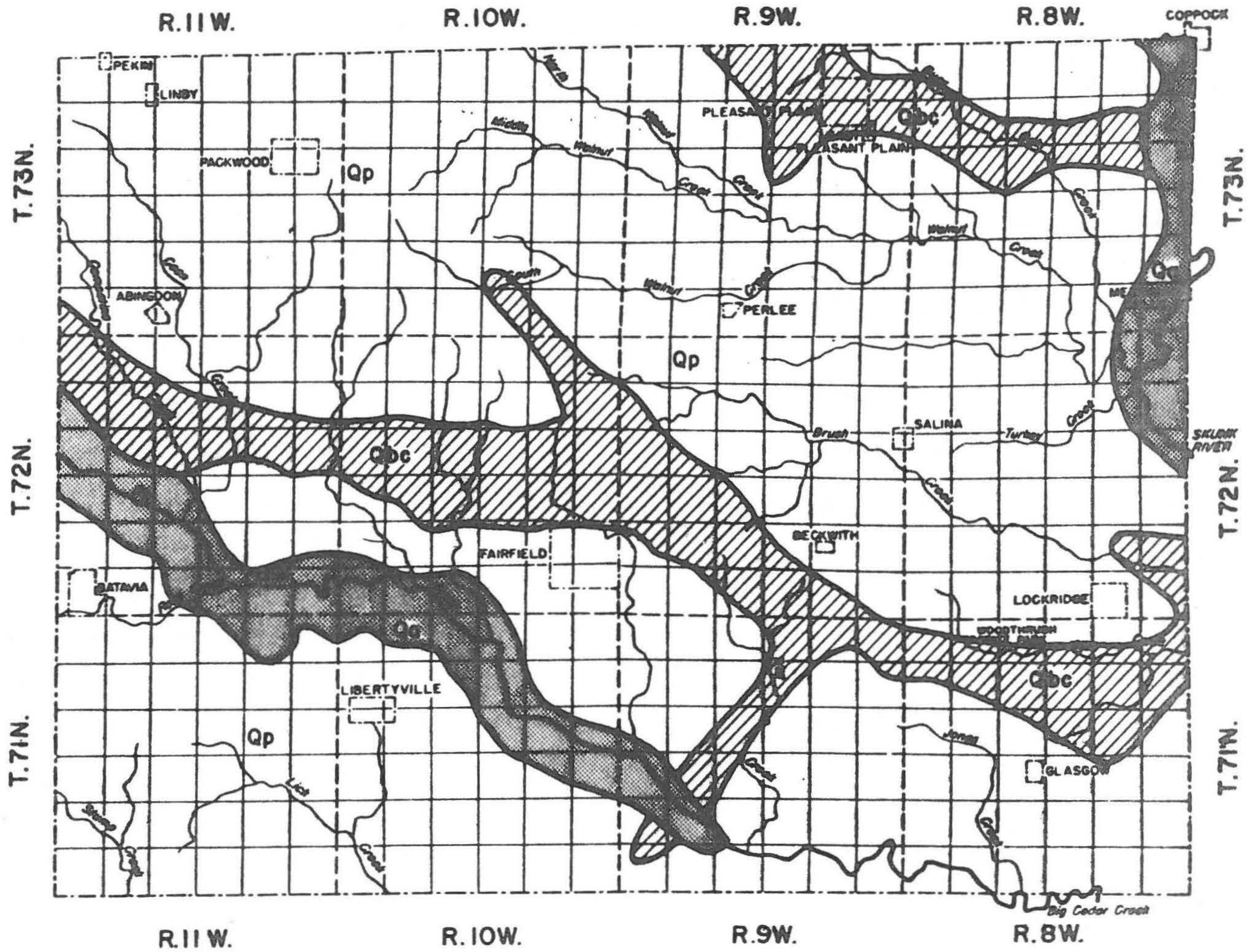
Average yields and water quality characteristics throughout the county for each of the aquifers are also summarized in the maps in Figures 10, 11, 12, 13, 14, and 15.

Table 1

GEOLOGIC AND HYDROGEOLOGIC UNITS IN JEFFERSON COUNTY

Age	Rock Unit	Description	Thickness Range	Hydrogeologic Unit	Water-Bearing Characteristics
Quaternary	Alluvium	Sand, gravel, silt, and clay	0-250 (feet)	Surficial aquifer	Fair to large yields (25 to 100 gpm)
	Glacial drift (undifferentiated)	Predominantly till containing scattered irregular bodies of sand and gravel			Low Yields (less than 10 gpm)
	Buried channel deposits	Sand, gravel, silt and clay			Small to large yields
Pennsylvanian	Des Moines Series	Shale; sandstones, mostly thin	0-150	Aquiclude	Low yields only from limestone and sandstone
Mississippian	Meramec Series	Limestone, sandy	150-400	Mississippian aquifer	Fair to low yields
	Osage Series	Limestone and dolomite, cherty			
	Kinderhook Series	Limestone, oolitic, and dolomite, cherty			
Devonian	Maple Mill Shale Sheffield Formation Lime Creek Formation	Shale; limestone in lower part	500-600	Devonian aquiclude	Does not yield water
	Cedar Valley Limestone; Wapsipinicon Formation	Limestone and dolomite; contains evaporites in southern half of Iowa		Devonian aquifer	Fair to low yields
Ordovician	Maquoketa Formation	Shale and dolomite	850-900	Maquoketa aquiclude	Does not yield water
	Galena Formation	Limestone and dolomite		Minor aquifer	Low yields
	Decorah Formation Platteville Formation	Limestone and thin shales; includes sandstone in SE Iowa		Aquiclude	Does not yield water
	St. Peter Sandstone	Sandstone		Cambrian-Ordovician aquifer	Fair yields
	Prairie du Chien Formation	Dolomite, sandy and cherty			High yields (over 500 gpm)
Cambrian	Jordan Sandstone	Sandstone	Over 1000		
	St. Lawrence Formation	Dolomite		Aquitard	Low yields
	Franconia Sandstone	Sandstone and shale			
	Dresbach Group	Sandstone		Dresbach aquifer	High to low yields
Precambrian	Undifferentiated	Coarse sandstones; crystalline rocks		Base of ground-water reservoir	Not known to yield water

Figure 2
SURFICIAL MATERIALS



Qa Alluvium

Qp Glacial Drift

Qbc Buried Channels

Figure 3

GEOLOGIC MAP

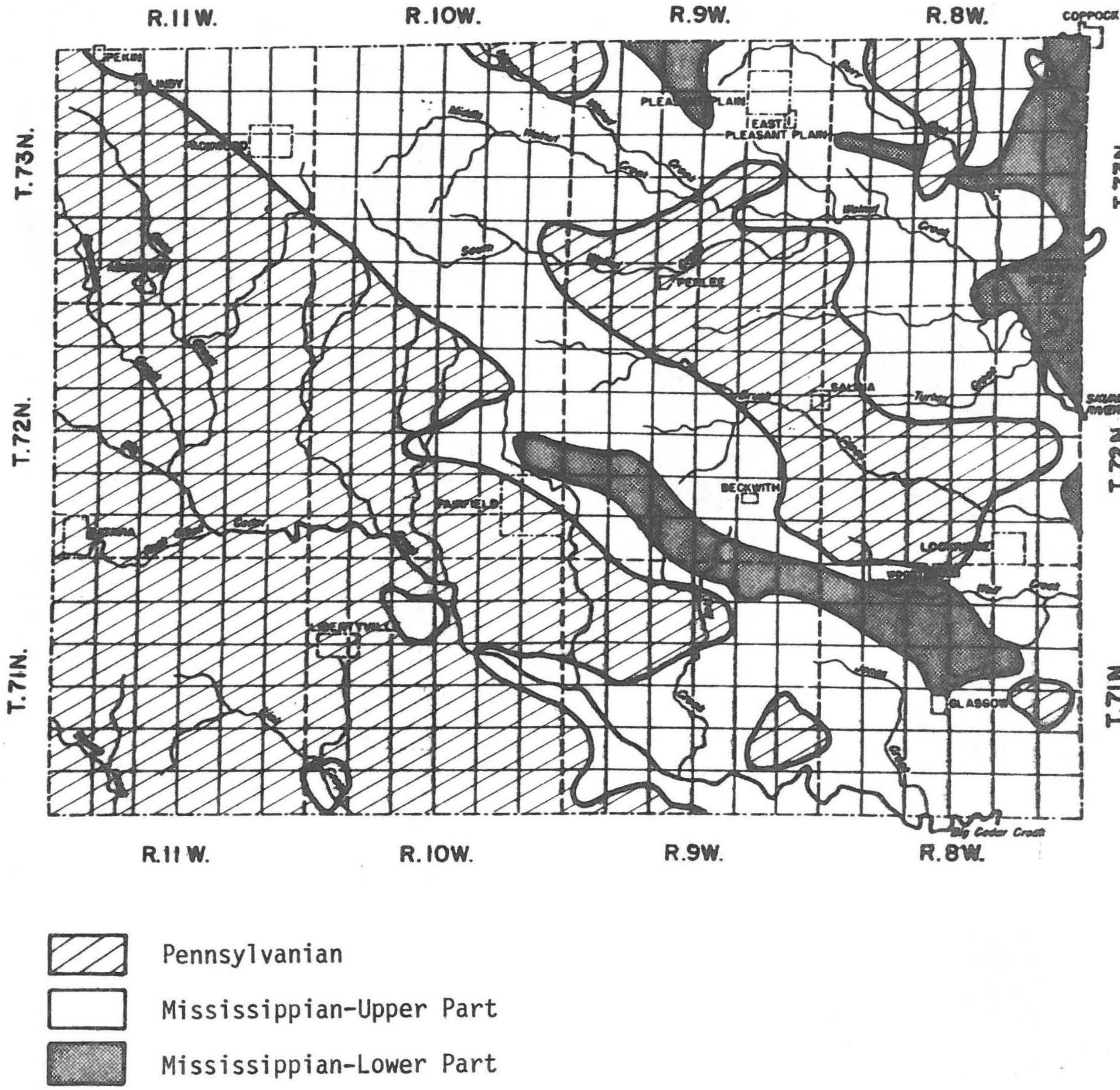


Figure 4

ELEVATION OF LAND SURFACE IN FEET ABOVE MEAN SEA LEVEL

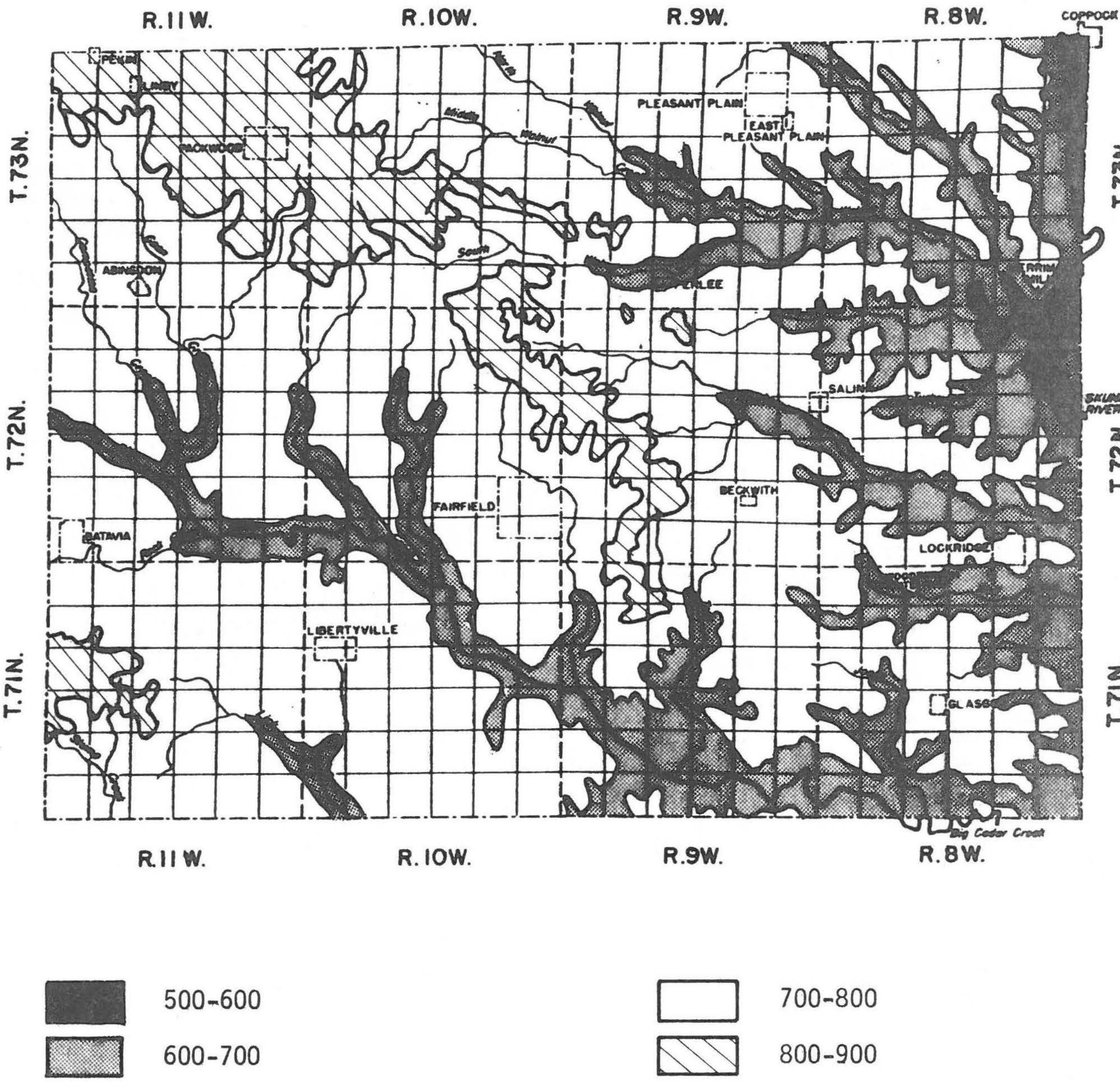
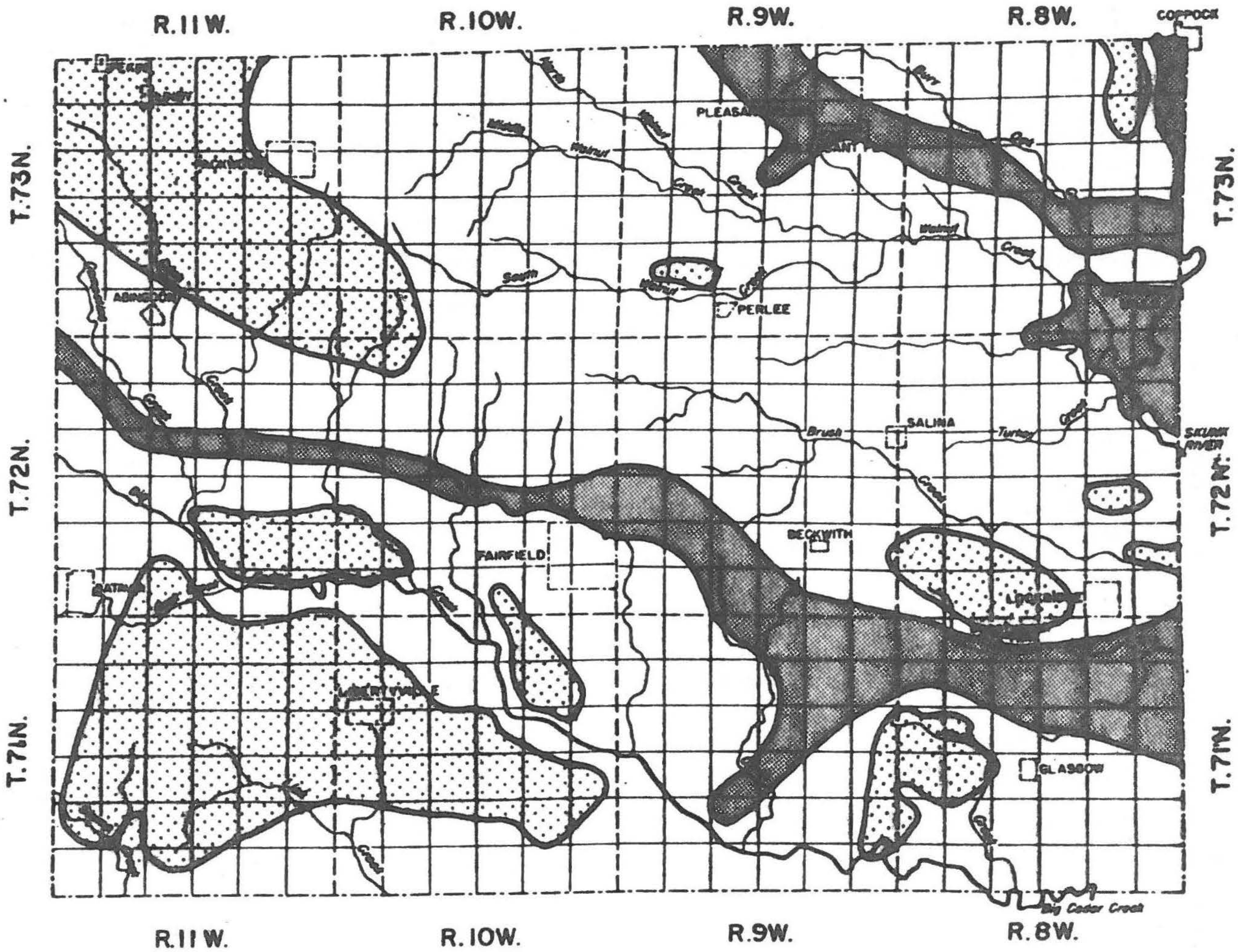


Figure 5

ELEVATION OF BEDROCK SURFACE IN FEET ABOVE MEAN SEA LEVEL



Below 600

600-700

Over 700

Figure 6

RANGE IN DEPTH TO JEFFERSON COUNTY'S PRINCIPAL ROCK AQUIFERS

	R.11W.	R.10W.	R.9W.	R.8W.
T.73N.	BEDROCK 50-150 MISSISSIPPIAN 100-300 DEVONIAN 700-800 CAMBRO-ORDOVICIAN 1800-2000	BEDROCK 50-150 MISSISSIPPIAN 100-300 DEVONIAN 700-800 CAMBRO-ORDOVICIAN 1800-2000	BEDROCK 50-150 MISSISSIPPIAN 50-200 DEVONIAN 500-700 CAMBRO-ORDOVICIAN 1700-1900	BEDROCK 50-150 MISSISSIPPIAN 50-200 DEVONIAN 400-700 CAMBRO-ORDOVICIAN 1600-1900
T.72N.	BEDROCK 50-200 MISSISSIPPIAN 100-300 DEVONIAN 600-800 CAMBRO-ORDOVICIAN 1700-1900	BEDROCK 50-200 MISSISSIPPIAN 100-300 DEVONIAN 500-900 CAMBRO-ORDOVICIAN 1700-2000	BEDROCK 50-150 MISSISSIPPIAN 100-300 DEVONIAN 500-900 CAMBRO-ORDOVICIAN 1700-2000	BEDROCK 50-150 MISSISSIPPIAN 100-200 DEVONIAN 400-700 CAMBRO-ORDOVICIAN 1600-1900
T.71N.	BEDROCK 50-150 MISSISSIPPIAN 100-300 DEVONIAN 600-900 CAMBRO-ORDOVICIAN 1700-2000	BEDROCK 50-150 MISSISSIPPIAN 100-200 DEVONIAN 600-800 CAMBRO-ORDOVICIAN 1700-1900	BEDROCK 0-150 MISSISSIPPIAN 0-150 DEVONIAN 600-900 CAMBRO-ORDOVICIAN 1700-2000	BEDROCK 0-150 MISSISSIPPIAN 0-150 DEVONIAN 600-800 CAMBRO-ORDOVICIAN 1700-1900

Figure 7

INDEX MAP FOR TYPICAL WELLS IN JEFFERSON COUNTY

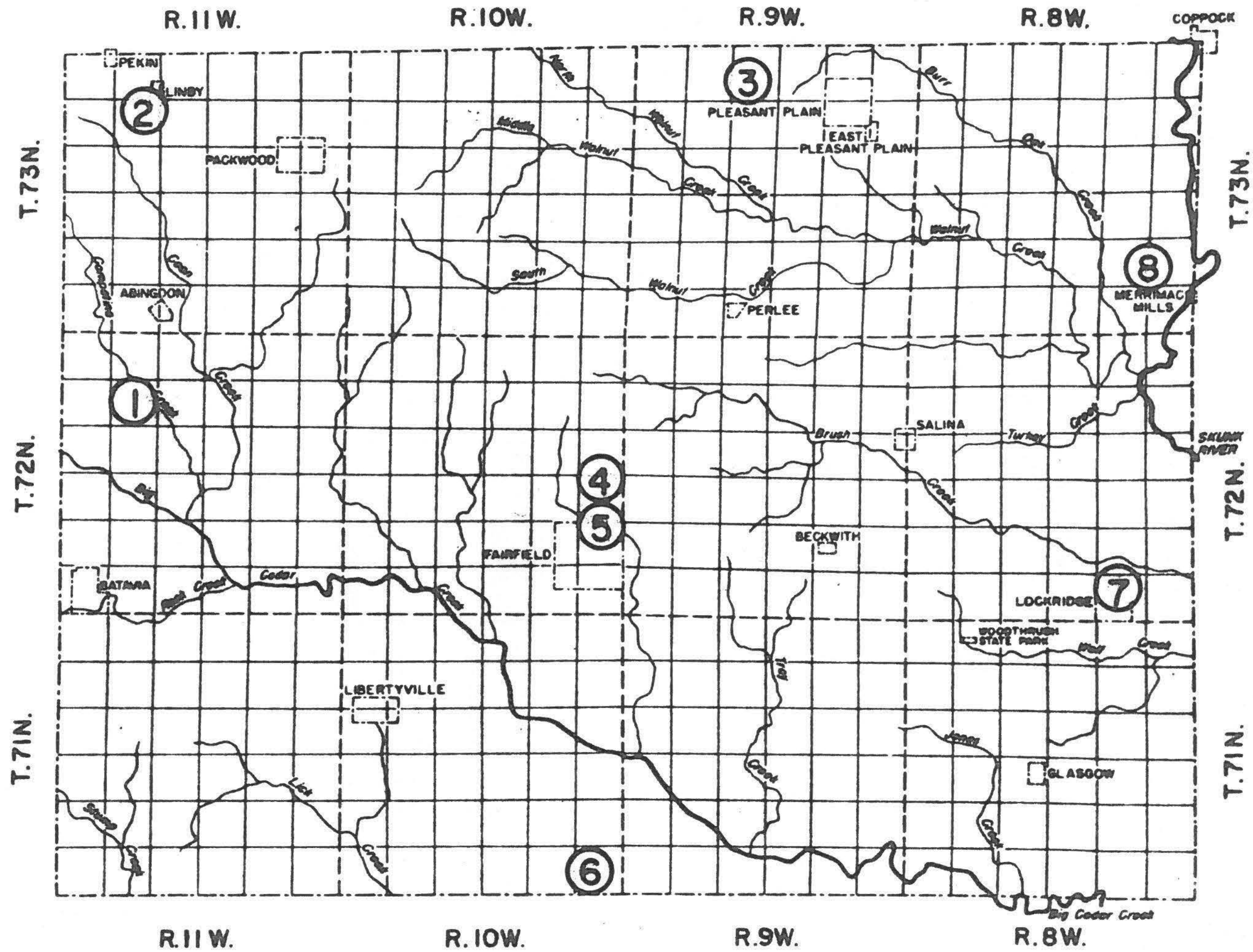


Figure 8

TYPICAL WELLS IN JEFFERSON COUNTY

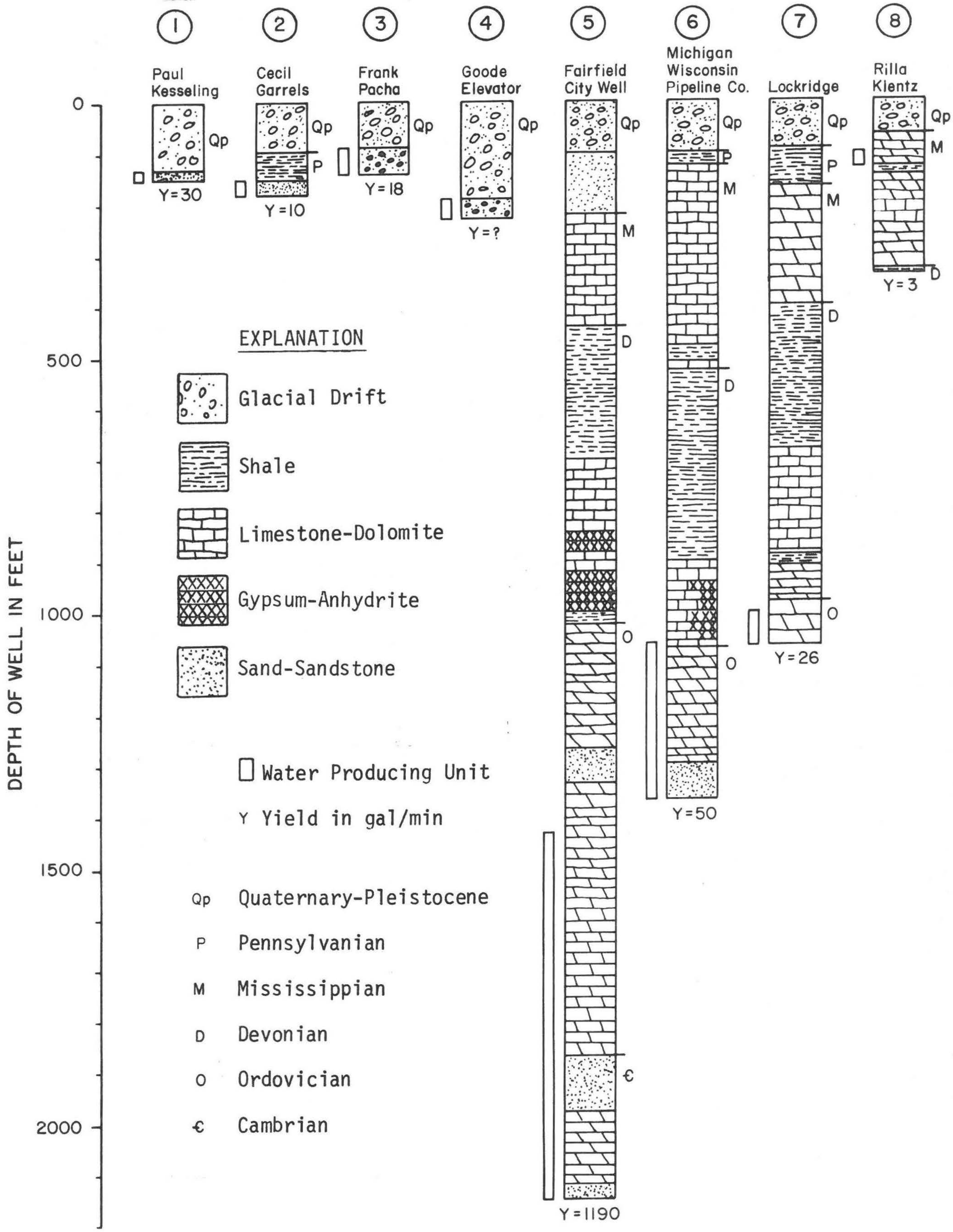
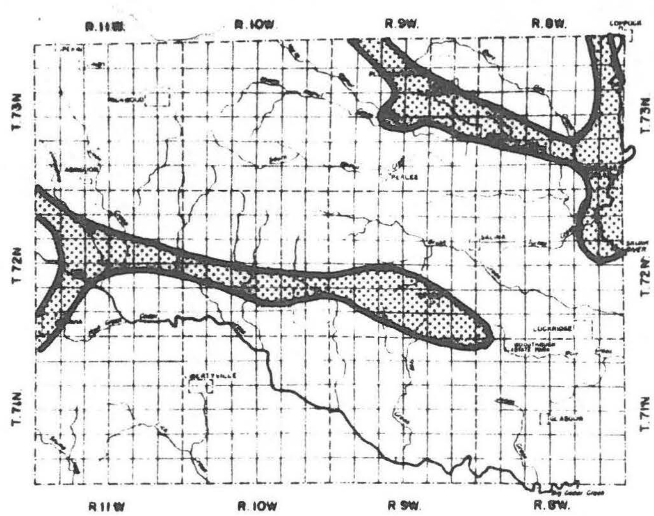


Figure 9
SURFICIAL AQUIFERS

Water Levels

Water levels in the surficial aquifers are difficult to analyze. Water rises to different levels in wells drilled into alluvial, buried-channel, and drift aquifers. The water table in the drift aquifer generally slopes from high land areas toward the streams and, changes noticeably throughout the year. Levels in drift and buried-channel aquifers respond rapidly to recharge from precipitation. Water levels in the alluvial aquifer fluctuate somewhat in the same way as those in the drift and buried-channel aquifers; however, the main influence on the alluvial aquifer is the stage (level) of the associated streams. Water levels will be high during periods of high stream stage and low during the low-stage periods.

Water levels in the drift aquifers commonly are from 10 to 50 feet below the land surface, and those in the buried-channel aquifer have been reported to be as low as 175 feet below the land surface. The water levels in alluvial wells are from 4 to 20 feet below the flood-plain surface, and the depth to the water surface will be accordingly deeper in wells located on terrace surfaces.



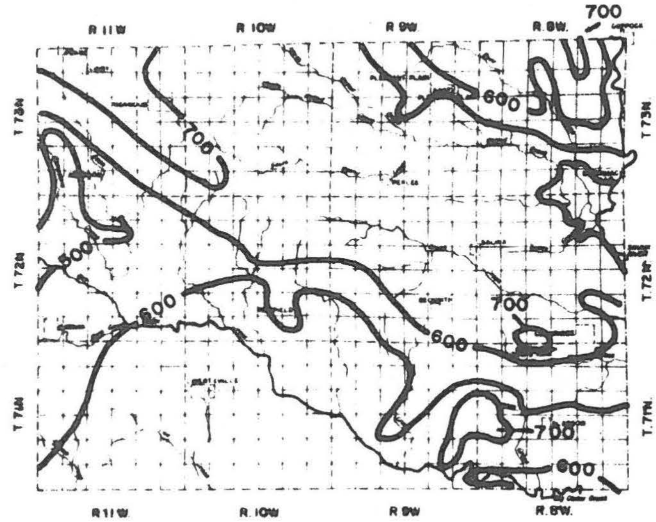
Water yields to wells in gallons per minute

Below 10

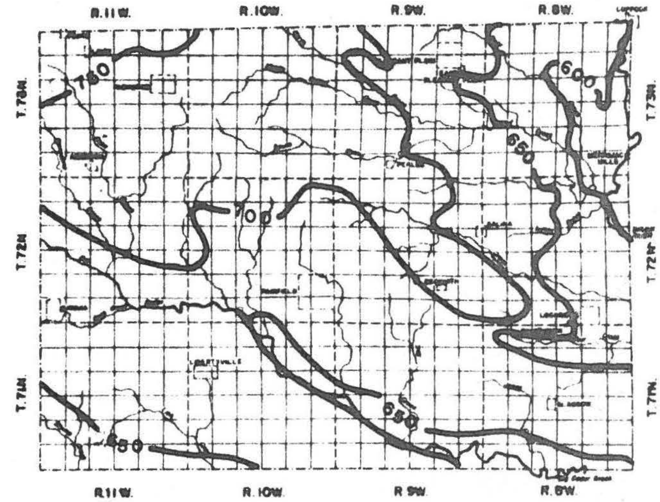
20-100

Figure 10

MISSISSIPPIAN AQUIFER

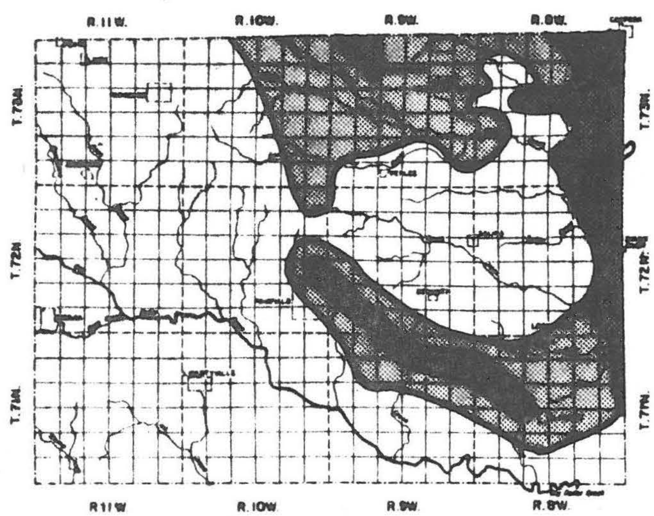


Elevation of Mississippian Aquifer in feet above mean sea level



Water levels in wells in feet above mean sea level

Upper Part

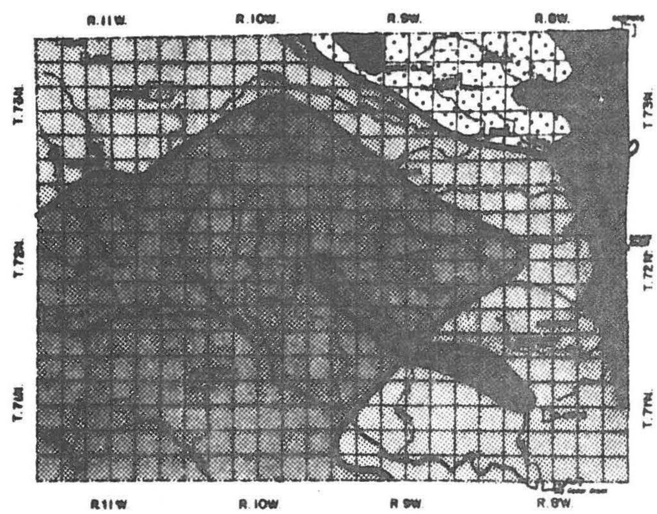


Water yields to wells in gallons per minute

Below 20

20-50

Aquifer not present



Dissolved solids content in milligrams per liter (mg/l)*

less than 500

500-1000

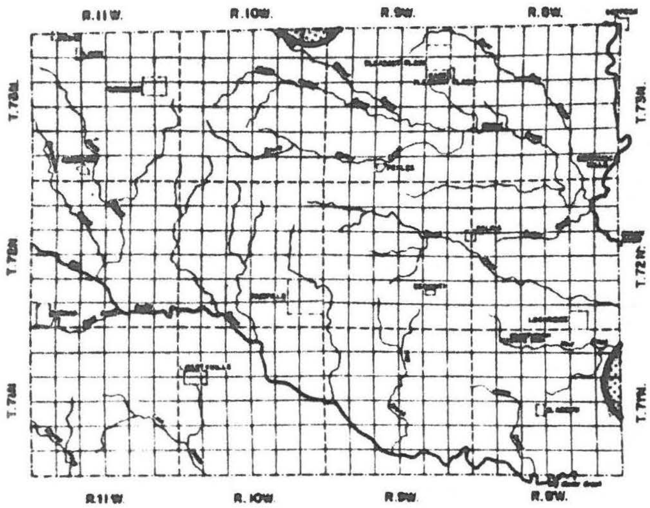
1000-2500

Aquifer not present

*Other water quality data in Figure 13

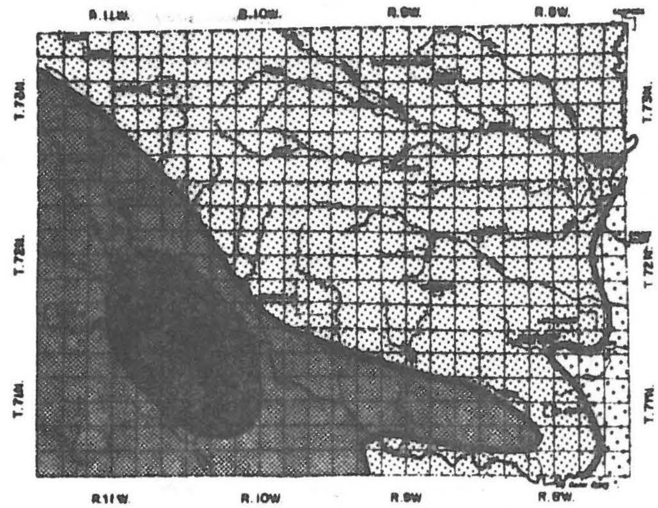
MISSISSIPPIAN AQUIFER

Lower Part



Water yields to wells in gallons per minute

- Below 20
- 20-50

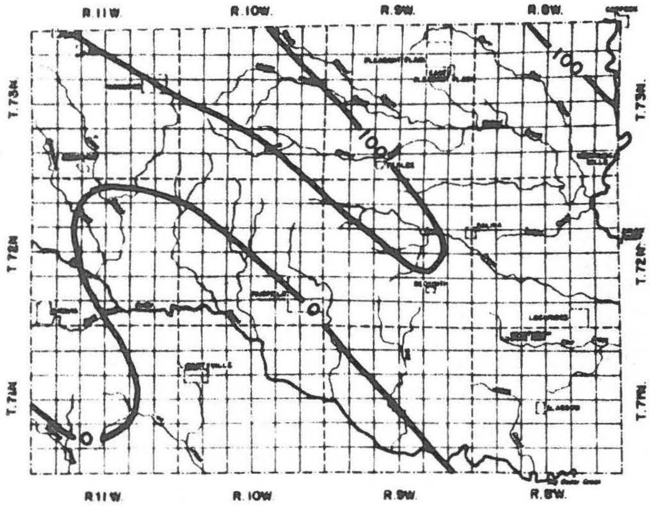


Dissolved solids content in milligrams per liter (mg/l)*

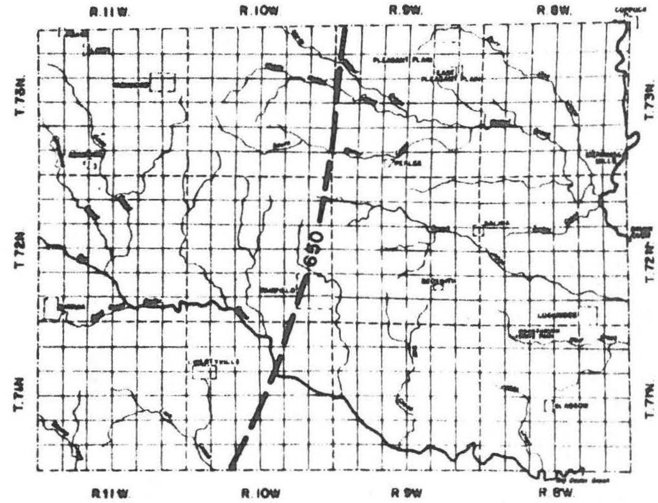
- 500-1000
- 1000-2500
- 2500-5000
- 5000-7500

*Other water quality data in Figure 14

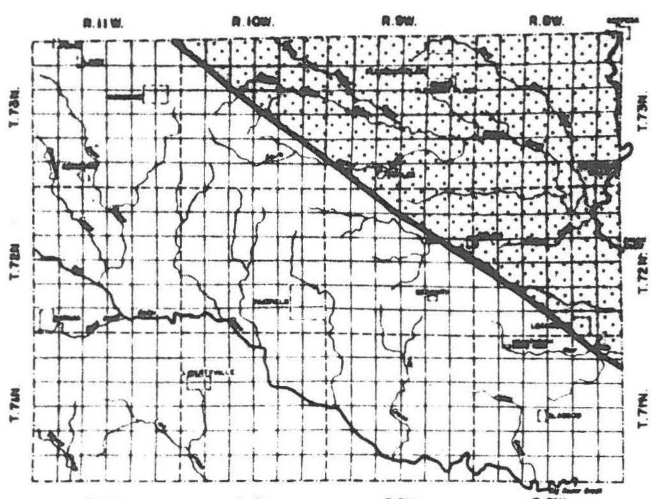
Figure 11
DEVONIAN AQUIFER



Elevation of Devonian Aquifer in feet above mean sea level

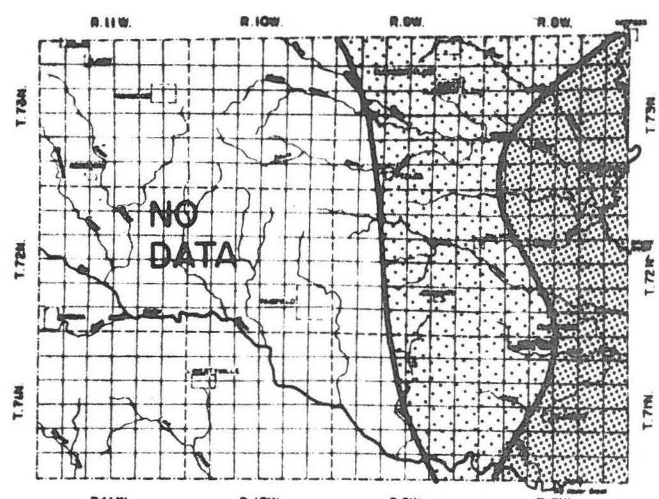


Water levels in wells in feet above mean sea level



Water yields to wells in gallons per minute

- Below 20
- 20-50

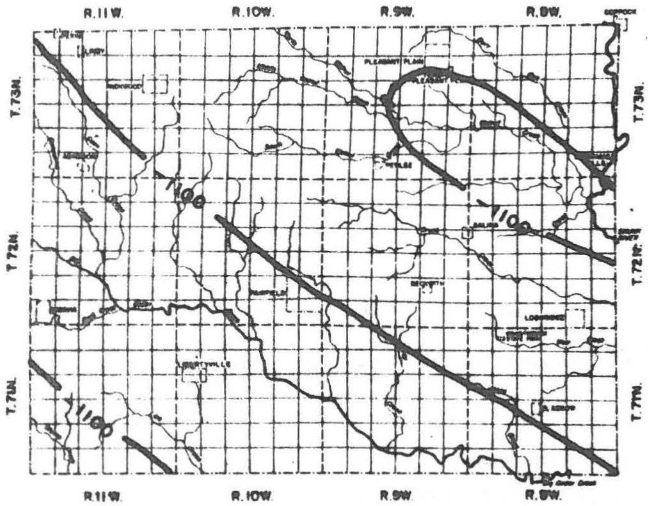


Dissolved solids content in milligrams per liter (mg/l)*

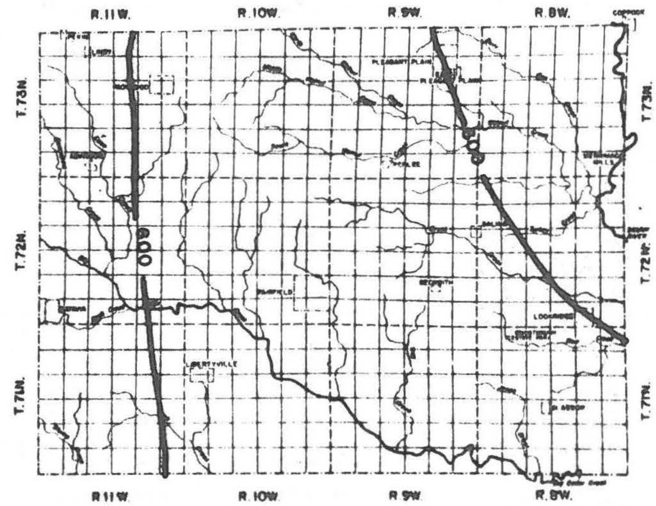
- 5000-7500
- Over 10,000

*Other water quality data in Figure 15

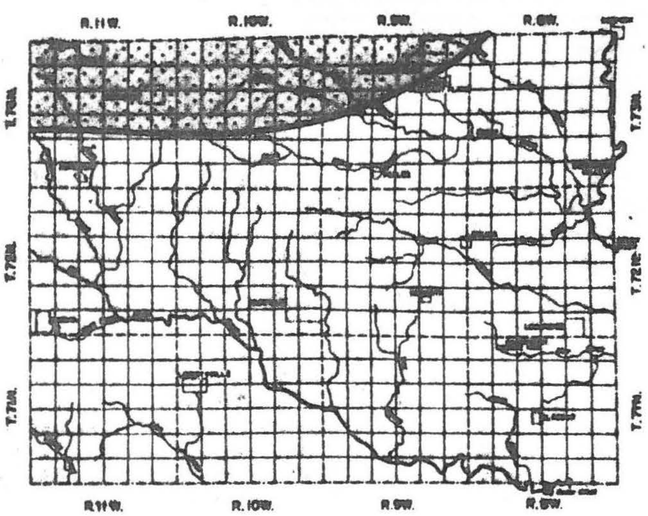
Figure 12
CAMBRO-ORDOVICIAN (JORDAN) AQUIFER



Elevation of Jordan aquifer in feet above sea level



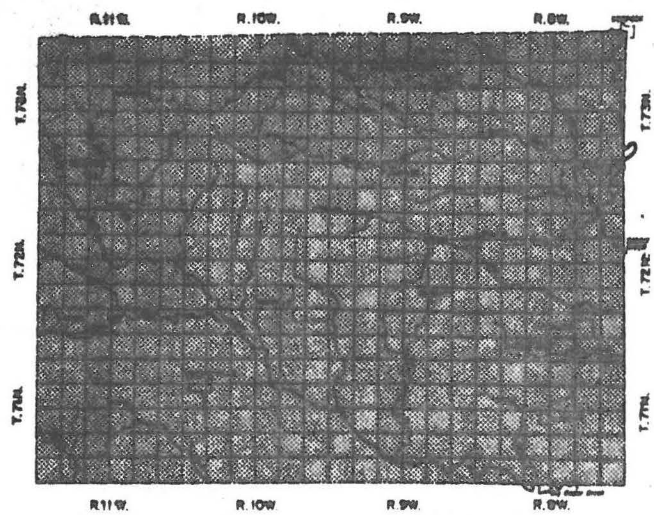
Water levels in wells in feet above mean sea level



Water yields to wells in gallons per minute

200-500

more than 1000



Dissolved solids content in milligrams per liter (mg/l)*

1000-1500

*Other water quality data in Figure 15

Table 2

SIGNIFICANCE OF MINERAL CONSTITUENTS AND PHYSICAL PROPERTIES OF WATER

Constituent or Property	Maximum Recommended Concentration	Significance
Iron (Fe)	0.3 mg/l	Objectional as it causes red and brown staining of clothing and porcelain. High concentrations affect the color and taste of beverages.
Manganese (Mn)	0.05 mg/l	Objectionable for the same reasons as iron. When both iron and manganese are present, it is recommended that the total concentration not exceed 0.3 mg/l.
Calcium (Ca) and Magnesium (Mg)		Principal causes for hardness and scale-forming properties of water. They reduce the lathering ability of soap.
Sodium (Na) and Potassium (K)		Impart a salty or brackish taste when combined with chloride. Sodium salts cause foaming in boilers.
Sulfate (SO_4)	250 mg/l	Commonly has a laxative effect when the concentration is 600 to 1,000 mg/l, particularly when combined with magnesium or sodium. The effect is much less when combined with calcium. This laxative effect is commonly noted by newcomers, but they become acclimated to the water in a short time. The effect is noticeable in almost all persons when concentrations exceed 750 mg/l. Sulfate combined with calcium forms a hard scale in boilers and water heaters.
Chloride (Cl)	250 mg/l	Large amounts combined with sodium impart a salty taste.
Fluoride (F)	2.0 mg/l	In central Iowa, concentrations of 0.8 to 1.3 mg/l are considered to play a part in the reduction of tooth decay. However, concentrations over 2.0 mg/l will cause the mottling of the enamel of children's teeth.
Nitrate (NO_3)	45 mg/l	Waters with high nitrate content should not be used for infant feeding as it may cause methemoglobinemia or cyanosis. High concentrations suggest organic pollution from sewage, decayed organic matter, nitrate in the soil, or chemical fertilizer.
Dissolved solids	500 mg/l	This refers to all of the material in water that is in solution. It affects the chemical and physical properties of water for many uses. Amounts over 2,000 mg/l will have a laxative effect on most persons. Amounts up to 1,000 mg/l are generally considered acceptable for drinking purposes if no other water is available.
Hardness (as $CaCO_3$)		This affects the lathering ability of soap. It is generally produced by calcium and magnesium. Hardness is expressed in milligrams per liter equivalent to $CaCO_3$ as if all the hardness were caused by this compound. Water becomes objectionable for domestic use when the hardness is above 100 mg/l; however, it can be treated readily by softening.
Temperature		Affects the desirability and economy of water use, especially for industrial cooling and air conditioning. Most users want a water with a low and constant temperature.

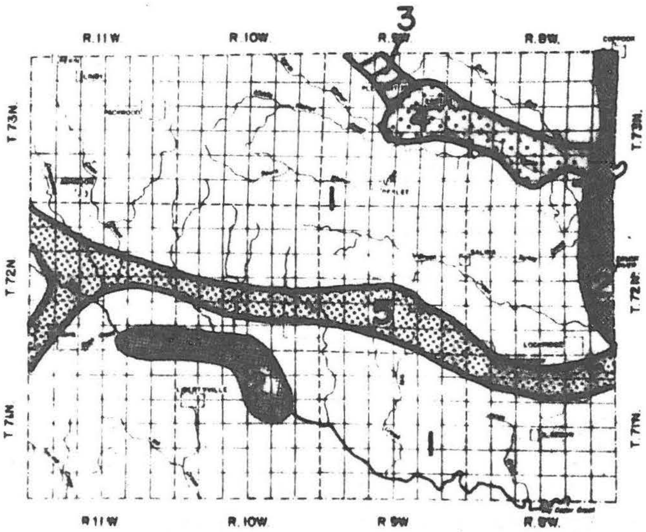
To the user, the quality of ground water is as important as the amount of water that an aquifer will yield. As ground water moves through soil and rock materials, it dissolves some of the minerals which, in turn, affect water quality. In addition to mineral content, bacterial and chemical contamination may be introduced through poorly constructed wells and seepage from other pollution sources.

Recommended standards for common water constituents are described in the table above. These are nationally accepted as guidelines for acceptable drinking water supplies. Limits for uses other than drinking often differ from these. For instance, water that is unacceptable for drinking and household use may be completely satisfactory for industrial cooling.

From analyses of ground water, averages (A) and ranges (R) of values in milligrams per liter (mg/l) for several mineral constituents are summarized in Figures 13, 14, and 15 for the 4 major aquifers in Jefferson County. Recommended concentrations for some constituents are often exceeded without obvious ill effects, although the water may be unpalatable. Water quality analyses for individual wells should be obtained to determine if concentrations of constituents that affect health are exceeded.

Figure 13
CHEMICAL CHARACTER OF GROUND WATER

Surficial Aquifers

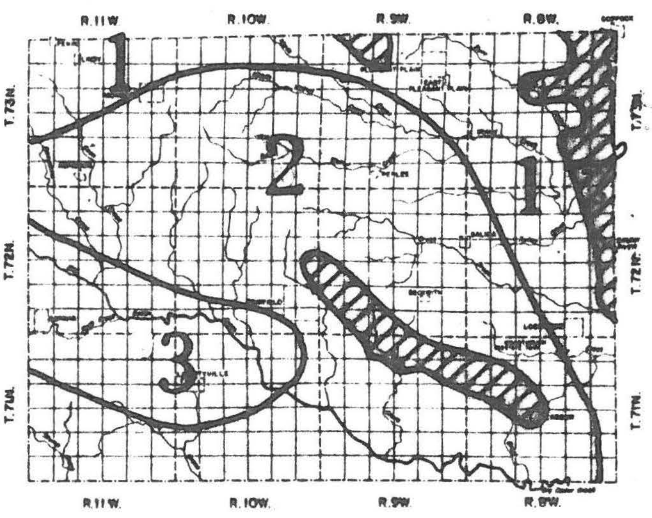


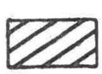
Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium + potassium (Na+K)	Bicarbonate (HCO_3)	Sulfate (SO_4)	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as $CaCO_3$)
					Drift Aquifer					
1	A	103	36	40	506	74	2.7	0.4	547	406
	R	77-150	24-61	13-88	304-739	11-393	0.5-7.5	0.2-0.7	357-982	292-625
					Alluvial Aquifer					
2	A	66	19	15	272	39	8	0.3	306	243
	R	41-85	11-29	7.2-56	149-407	8.2-113	.5-31	.2-1.0	221-436	168-330
					Water from the buried-channel aquifer					
3	A	96	32	45	499	49	7.8	.2	501	374
	R	55-135	17-49	3.5-133	305-673	.1-140	.5-36	0-.6	311-676	259-489
4	A	53	20	153	579	46	12	.4	591	215
	R	24-88	3.1 -39	87-244	410-803	.2-170	2.5-28	0-1.0	357-715	78-360
5	A	160	52	95	676	250	2.0	.4	946	616
	R	128-175	46-61	42-130	548-775	110-362	.5-4.0	.2-.9	782-1120	510-685

The alluvial and drift aquifers yield good quality water. Water from buried channels has a higher dissolved solids content, but it is less than that from rock aquifers. Water temperatures average 54°F and normally do not vary more than 6°.

Mississippian Aquifer

Upper Part





Aquifer not present

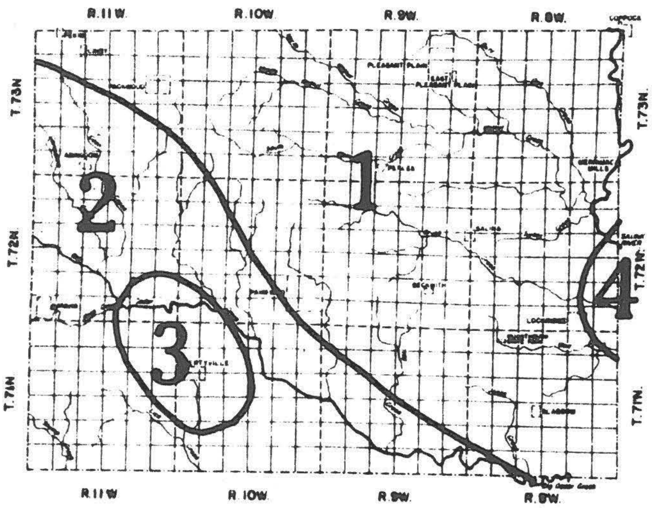
Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium + Potassium (Na+K)	Bicarbonate (HCO_3)	Sulfate (SO_4)	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as $CaCO_3$)
1	A	104	32	43	472	91	4.5	0.4	537	399
	R	66-156	19-50	8.8-118	332-802	5.8-240	0.5-24	0-1.8	345-737	315-545
2	A	299	92	125	281	1100	8.5	.8	1950	1130
	R	197-547	47-124	33-214	44-434	750-1590	2-22	.3-1.6	1420-2740	723-1560
3	A	61	31	404	608	587	22	2.1	1400	281
	R	54-69	25-37	275-513	471-736	400-710	20-25	.9-3.5	1060-1490	240-324

Water in the upper part of the Mississippian aquifer is more highly mineralized than that typically found in the surficial aquifers and is usually very hard. With exception of along the northern and eastern margins of the county, dissolved solids content is high, particularly in the central portion of the county. For most of the county, sulfate concentrations exceed recommended standards. Average water temperature is 55°F with a range of 51°-60°.

Figure 14

Mississippian Aquifer

Lower Part



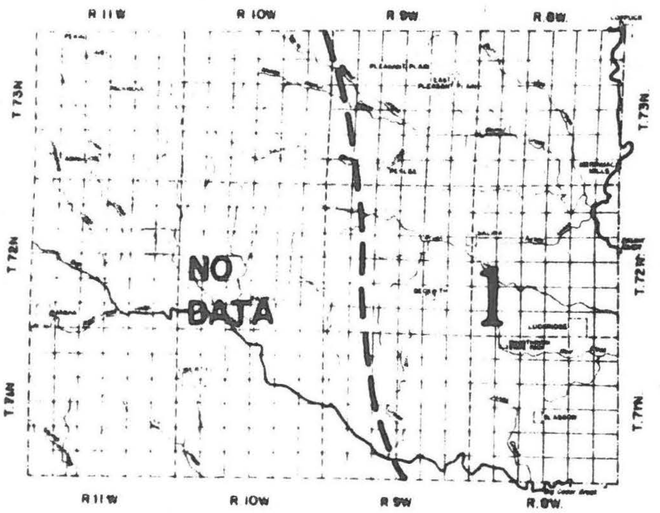
Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium + Potassium (Na+K)	Bicarbonate (HCO_3)	Sulfate (SO_4)	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as $CaCO_3$)
1	A	78	38	276	592	354	57	1.6	1110	355
	R	38-128	19-68	143-489	465-754	260-560	.5-150	.5-2.5	879-1480	176-581
2	A	102	52	718	459	1340	169	2.6	2710	469
	R	35-193	15-107	451-994	266-595	920-1660	19-365	.5-6	2220-3250	148-891
3	A	215	82	1375	956	2520	610	1.8	5260	875
	R	102-452	58-120	1040-1630	378-3640	1800-4020	290-930	1.4-3	4070-7330	497-1620
4	A	90	37	42	504	46	6	0.4	490	380
	R	30-160	10-61	9-107	298-710	1-186	0-69	0-1.2	280-800	160-575

Water in the lower part of the Mississippian aquifer is generally of poorer quality than found in the upper part. Throughout the county the water is exceptionally hard, and with the exception of the extreme east-central part of the county, greatly exceeds recommended standards for sulfate and total dissolved solids. Average water temperature is 55°F with a range of 51°-60°.

Figure 15

CHEMICAL CHARACTER OF GROUND WATER

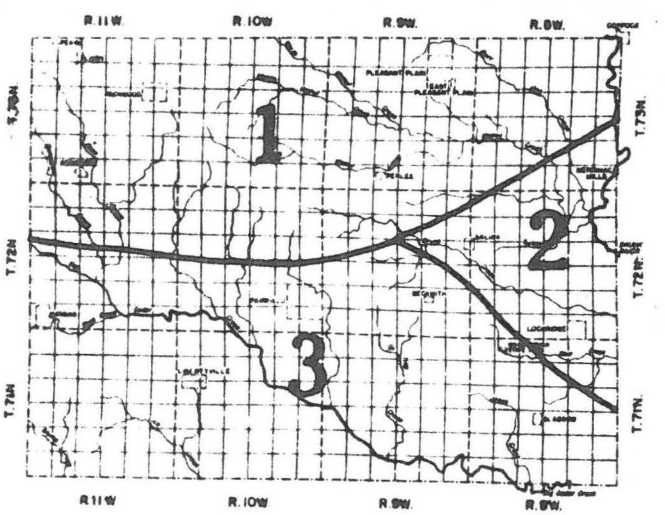
Devonian Aquifer



Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium & Potassium (Na+K)	Bicarbonate (HCO_3)	Sulfate (SO_4)	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as $CaCO_3$)
1	A	492	138	2220	289	4580	1100	2.8	9570	1800
	R	441-617	97-155	1850-2700	183-378	4030-5000	550-1740	1.9-5.0	8240-11,100	1580-2180

Water in the Devonian aquifer is highly mineralized and objectionably hard wherever it has been encountered in the county. Because of extremely high total dissolved solids, primarily sodium-potassium, chloride, and sulfate, the water from the Devonian is unfit for human or animal consumption. The average water temperature is 60°F and ranges between 54° and 64° in southeast Iowa.

Cambro-Ordovician Aquifer



Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium + Potassium (Na+K)	Bicarbonate (HCO_3)	Sulfate (SO_4)	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as $CaCO_3$)
1	A	106	50	202	304	552	52	1.2	1180	470
	R	98-116	46-54	192-211	283-337	520-600	38-60	1.0-1.5	1120-1240	452-510
2	A	93	42	232	295	520	79	1.4	1130	406
	R	86-105	37-47	223-249	288-305	489-543	69-85	1.2-1.6	1110-1150	372-455
3	A	84	34	267	298	476	124	1.7	1160	349
	R	78-92	26-41	247-283	283-317	455-500	100-148	1.0-2.2	1100-1220	322-388

This deep aquifer yields water of relatively good quality compared to the other rock aquifers. However, the water is noticeably hard and exceeds recommended standards for sulfate and dissolved solids. Water temperatures are higher than other rock aquifer sources averaging 72°F and ranging up to about 75°.

RECOMMENDATIONS FOR PRIVATE WATER WELLS

Contracting for Well Construction

To protect your investment and guarantee satisfactory well completion, it is a good idea to have a written agreement with the well driller. The agreement should specify in detail:

size of well, casing specifications, and types of screen and well seal

methods of eliminating surface and subsurface contamination

disinfection procedure to be used

type of well development if necessary

test pumping procedure to be used

date for completion

itemized cost list including charges for drilling per foot, for materials per unit, and for other operations such as developing and test pumping

guarantee of materials, workmanship, and that all work will comply with current recommended methods

liability insurance for owner and driller

Well Location

A well should be located where it will be least subject to contamination from nearby sources of pollution. The Iowa State Department of Health recommends minimum distances between a new well and pollution sources, such as cesspools (150 ft.), septic tanks (50 ft.), and barnyards (50-100 ft. and down slope from well). Greater distances should be provided where possible.

The well location should not be subject to flooding or surface water contamination. Select a well-drained site, extend the well casing a few feet above the ground, and mound earth around it. Diversion terraces or ditches may be necessary on slopes above a well to divert surface runoff around the well site.

In the construction of all wells care should be taken to seal or grout the area between the well bore and the well casing (the annulus) as appropriate so that surface water and other pollutants cannot seep into the well and contaminate the aquifer.

Locate a well where it will be accessible for maintenance, inspection, and repairs. If a pump house is located some distance from major buildings and wired separately for power, continued use of the water supply will be less jeopardized by fire in major buildings.

Water Treatment

Water taken from a private well should ideally be tested every six months. The University Hygienic Laboratory will do tests for coliform bacteria, nitrate, iron, hardness, and iron bacteria in drinking water for private individuals. Special bottles must be used for collecting and sending water samples to the laboratory. A sample kit can be obtained by writing to the University Hygienic Laboratory, Medical Laboratories Building, Iowa City, Iowa 52242. Indicate whether your water has been treated with chlorine, iodine, or bromine, for different sample bottles must be used for treated and untreated water. The charge for the bacterial test is $3.00; for iron, hardness and nitrate, it is $3.00; and for iron bacteria, $5.00. If your well is determined to be unsafe, advice for correcting the problem can be obtained from your county or state Department of Health. Several certified private laboratories also run water analyses.

Shock chlorination is recommended following the construction and installation of a well and distribution system and anytime these are opened for repairs or remodeling. A strong chlorine solution is placed in the well and complete distribution system to kill nuisance and disease-causing organisms. If the first shock chlorination does not rid the water supply of bacteria it should be repeated, if this does not solve the problem the well should be abandoned or the water should be continuously disinfected with proper chlorination equipment.

Since most of the ground waters in Jefferson County are mineralized, water softening and iron removal equipment may make water more palatable and pleasant to use. Softened water contains increased sodium; contact your physician before using a softener if you are on a sodium-restricted diet. Chlorination followed by filtration will remove most forms of iron and iron bacteria. Iron bacteria has no adverse effect on health but will plug wells, water lines, and equipment and cause tastes and odors. Iron removal equipment can be used if problems persist.

Well Abandonment

Wells taken out of service provide easy access for pollution to enter aquifers supplying water to other wells in the vicinity. Unprotected wells may also cause personal injury. Proper abandonment procedures should be followed to restore the natural conditions that existed before well construction and prevent any future contamination. Permanent abandonment requires careful sealing. The well should be filled with concrete, cement grout, or sealing clays throughout its entire length. Before dug or bored wells are filled at least the top 10 feet of lining should be removed so surface waters will not penetrate the subsurface through a porous lining or follow cracks in or around the lining. The site should be completely filled and mounded with compacted earth.

ABANDONED WELLS SHOULD NEVER BE USED FOR DISPOSAL OF SEWAGE OR OTHER WASTES.

SOURCES OF ADDITIONAL INFORMATION

In planning the development of a ground water supply or contracting for the drilling of a new well additional or more specific information is often required. This report section lists several sources and types of additional information.

State Agencies That May Be Consulted

Iowa Geological Survey[1]	123 North Capitol Iowa City 52242	(319) 338-1173
State Health Department[2,6]	Lucas Building Des Moines 50319	(515) 281-5787
Iowa Natural Resources Council[3]	Wallace Building Des Moines 50319	(515) 281-5914
Iowa Dept. of Environ. Quality[4]	Wallace Building Des Moines 50319	(515) 281-8854
University Hygienic Laboratory[5]	270 Medical Lab. Bldg. Iowa City 52242	(515) 353-5990
Cooperative Extension Service in Agriculture and Home Economics[6]	110 Curtis Hall, ISU Ames 50011	(515)294-4569

Functions:

[1] Geologic and ground water data repository, consultant on well problems, water development and related services.

[2] Drinking water quality, public and private water supplies

[3] Water withdrawal regulation and Water Permits for well withdrawing more than 5000 gpd.

[4] Municipal supply regulation and well construction permits

[5] Water quality analysis

[6] Advice on water systems design and maintenance

Well Drillers and Contractors

The listing provided here was drawn from an Iowa Geological Survey mailing list, those selected are within a radius of 50 miles of Jefferson County. For a state-wide listing contact either the Iowa Water Well Drillers Association, 4350 Hopewell Ave., Bettendorf, Iowa 51712, (319) 355-7528 or the Iowa Geological Survey.

Bailey Well Co.
203 East Main
New London, Iowa 52645

Bruinekool Well Co.
Oskaloosa, Iowa 52577

Detrick Well Co.
R. R. #1
New London, 52645

Gingerich Well Co.
Kalona, Iowa 52247

Kramer Well Co.
Mt. Pleasant, Iowa 52641

Latta & Sons Well Drilling
Riverside, Iowa 52327

Lyon Well Co.
Salem, Iowa 52649

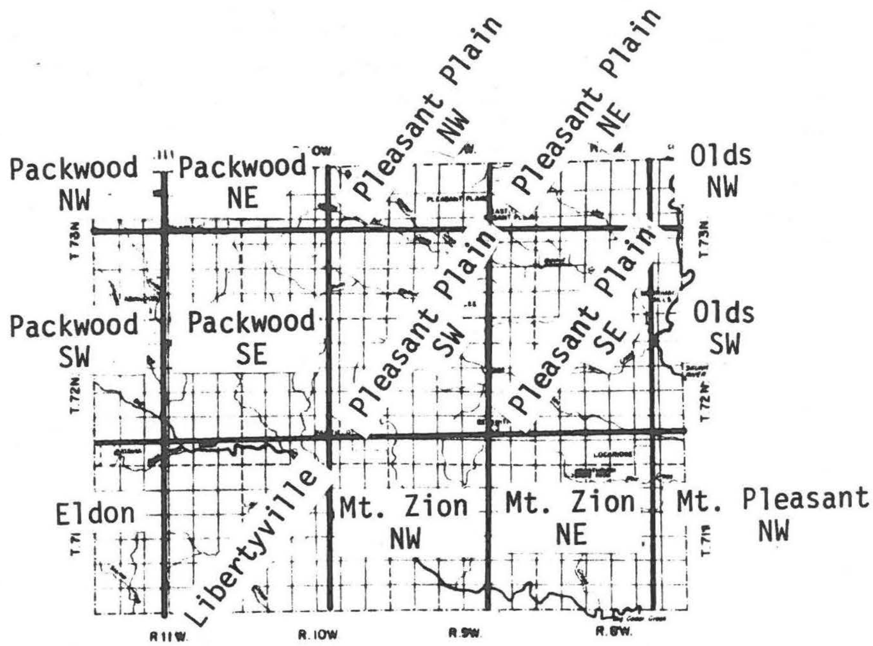
Miller & Son Well Co.
Kalona, Iowa 52247

Schlicher Bros. Well
Hwy. 34 West
Fairfield, Iowa 52556

Schmeiser Well Co.
1111 Hageman St.
Burlington, Iowa 52601

Wilson Well Co.
R.R. #3
Burlington, Iowa 52601

Topographic Maps (Available from the Iowa Geological Survey)



Map Title	Date	Scale	Contour Interval
	(Published)		
Libertyville	1965	1:24,000	10 ft.
Douds	1965	1:24,000	10 ft.
	(Preliminary)		
Packwood SW		1:24,000	10 ft.
Packwood NW		1:24,000	10 ft.
Packwood NE		1:24,000	10 ft.
Packwood SE		1:24,000	10 ft.
Pleasant Plain SW		1:24,000	10 ft.
Pleasant Plain NW		1:24,000	10 ft.
Pleasant Plain NE		1:24,000	10 ft.
Pleasant Plain SE		1:24,000	10 ft.
Mt. Zion NW		1:24,000	10 ft.
Mt. Zion NE		1:24,000	10 ft.
Olds NW		1:24,000	10 ft.
Olds SW		1:24,000	10 ft.
Mt. Pleasant NW		1:24,000	10 ft.

Useful Reference Materials

Coble, R. W., and Roberts, J.V., 1971, The water resources of Southeast Iowa, Iowa Geological Survey, Water Atlas No. 4.

Horick, P.J., and Steinhilber, W.L., 1973, Mississippian aquifer of Iowa, Iowa Geological Survey, Misc. Map Series No. 3.

Horick, P.J., and Steinhilber, W.L., 1978, Jordan aquifer of Iowa, Iowa Geological Survey, Misc. Map Series No. 6.

Iowa State Department of Health, 1971, Sanitary standards for water wells, State Department of Health, Environmental Engineering Service.

Van Eck, O.J., 1971, Optimal well plugging procedures, Iowa Geological Survey Public Information Circular No. 1.

Van Eck, O.J., 1978, Plugging procedures for domestic wells, Iowa Geological Survey, Public Information Circular No. 11.