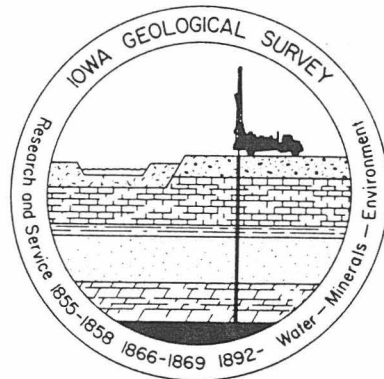


# GROUND WATER RESOURCES



## Van Buren County

Open File Report 80-89 WRD

Compiled by DONIVAN L. GORDON

## GROUND-WATER RESOURCES OF VAN BUREN COUNTY

### Introduction

Approximately 80 percent of the water used in Van Buren County comes from underground sources. It is estimated that the use of ground water in the county currently approaches 0.4 billion gallons per year. For comparison, this amount would provide each resident with 109 gallons of water a day during a year. Actually, few if any households use this much water, and this greater per capita use reflects the county's commercial, agribusiness, and municipal needs.

The users of ground water in the county draw their supplies from several different geologic sources. Several factors must be considered in determining the availability of ground water and the adequacy of a supply source:

distribution - having water where it is needed,

accessibility - affects the costs for drilling wells and pumping water,

yield - relates to the magnitude of the supply that can be sustained,

quality - determines for what purposes the water can be used.

In terms of these factors, there are few locations in Van Buren County where the availability of ground water is not limited to some degree. The most common limitation is poor water quality, that is, highly mineralized ground water. Secondary limitations are generally related to poor distribution, small yields from some sources, and poor accessibility due to the great depths to adequate sources.

### Occurrence of Ground Water in Van Buren County

The occurrence of ground water is influenced by geology -- the position and thickness of the rock units, their ability to store and transmit water, and their physical and chemical make-up. Geologic units that store and transmit water and yield appreciable amounts to wells are called aquifers. The best aquifers are usually composed of unconsolidated sand and gravel, porous sandstone, and porous or fractured limestone and dolomite. Other units with materials such as clay and silt, shale, siltstone, and mudstone yield little or no water to wells. These impermeable units are called aquicludes or aquitards and commonly separate one aquifer unit from another.

In Van Buren County there are four principal aquifers from which users obtain water supplies. The loose, unconsolidated materials near the land surface comprise the surficial aquifer. Below this there are three major rock aquifers -- the Mississippian aquifer, the Devonian aquifer, and the Cambro-Ordovician aquifer. Although a principal water supply source in several counties, the Devonian aquifer in Van Buren cannot be considered for most purposes because its water is highly mineralized. Figure 1 shows the geologic relations of these beneath the county. Each of the aquifers has its own set of geologic, hydrologic, and water quality characteristics which determine the amount and potability (suitability for drinking) of water it will yield.

### Surficial Aquifers

Unconsolidated deposits at the land surface are comprised of mixtures of clay, silt, sand, gravel, and assorted boulders. The water-yielding potential of the surficial deposits is greatest in units composed mostly of sand and/or gravel. Three types of surficial aquifers are used: the alluvial aquifer, the drift aquifer, and the buried-channel aquifer.

The alluvial aquifer consists mainly of the sand and gravel transported and deposited by modern streams and makes up the floodplains and terraces in major valleys. Alluvial deposits are shallow, generally less than 50-60 feet and may be contaminated by percolating surface water.

The drift aquifer is the thick layer of soils materials deposited over the bedrock by glacial ice which invaded the county at least twice in the last two million years. The composition of the glacial drift varies considerably, and where it is silty and clayey it does not yield much water. Locally there are lenses or beds of sand and gravel within the drift which are thick and widespread enough to serve as dependable water sources. These lenses are difficult to locate because they are irregular in shape and buried within the drift deposits. Usually one or two sand layers can be found in most places that will yield enough water for domestic needs.

Buried-channel aquifers consist of stream alluvium that partially fills valleys that existed before the glacial period. The valleys, overridden by glaciers, are now buried under glacial and more recent alluvial deposits.

The distribution, yields, and water quality characteristics for the surficial aquifers are summarized in Figures 2, 9, and 13. An indication of aquifer depths can be obtained by comparing the elevations of the top (the land surface) and the bottom (the bedrock surface) of the surficial deposits from Figures 4 and 5. The thickness of the glacial drift and the depth of the buried channels are determined by subtracting bedrock from land surface elevations at selected locations.

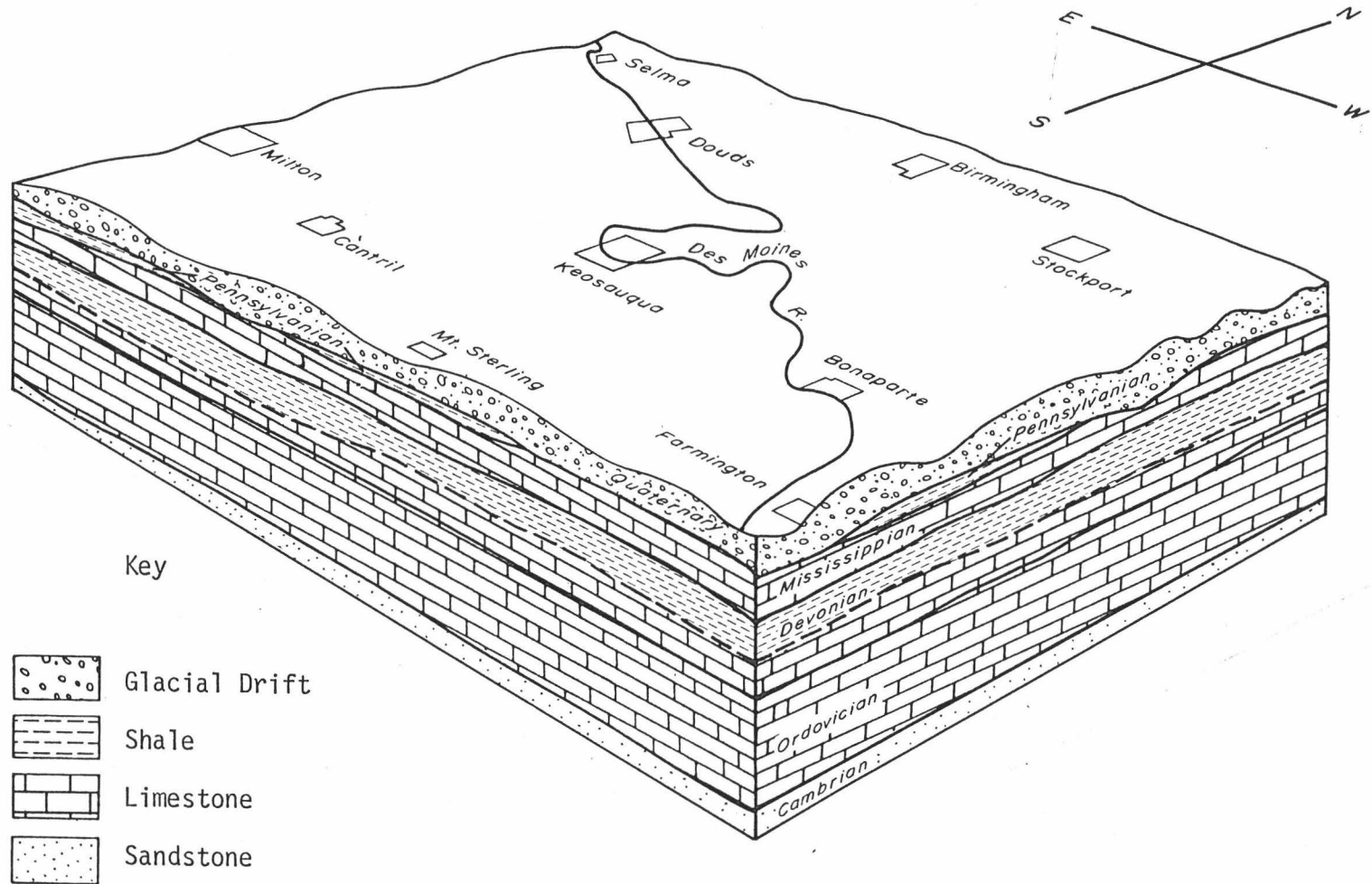


Figure 1

BLOCK DIAGRAM SHOWING THE GEOLOGY OF VAN BUREN COUNTY

## Rock Aquifers

Below the surficial materials is a thick sequence of layered rocks formed from deposits of rivers and shallow seas that alternately covered the state within the last 600 million years. The geologic map (Figure 3) shows the geologic units at the surface of this sequence. These rocks are Mississippian and Pennsylvanian in age. The Pennsylvanian rocks are principally shales and act as an aquiclude. The thickness of the Pennsylvanian rocks varies from 0 feet where Mississippian rocks are uppermost to a maximum of approximately 90-100 feet.

Underlying the Pennsylvanian aquiclude is a sequence of older rocks, parts of which form the three major rock aquifers in Van Buren County. This sequence and the water-bearing characteristics of the aquifers and aquicludes are shown in Table 1.

Examples of the sequence of rock units encountered in drilling existing wells at various locations in Van Buren County are indexed and illustrated in Figure 7 and 8. The geologic unit that supplies ground water and the amount of water yielded to the well are shown next to each of the well logs.

The accessibility of ground water in the rock aquifers depends first on the depth to the aquifer. The deeper a well must be, the greater the cost for well construction and pumping. The depths to and thicknesses of units at specific sites will vary somewhat because of irregularities in the elevation of the land surface and in the elevation of the tops of the underlying rock units. Estimates of depths and thicknesses can be made by comparing Figure 4 with the maps of aquifer elevations in Figures 10, 11, and 12. The range in depth below land surface to the top of the county's principal bedrock aquifers is given for each township in Figure 6.

The second factor which affects accessibility is the level to which the water will rise in the well (the static water level). Since water in the rock aquifers is under artesian pressure, the water rises in the well once it penetrates the aquifer. This rise in water level can reduce the cost of pumping. Average static water levels in Van Buren County wells are shown in Figures 10, 11, and 12.

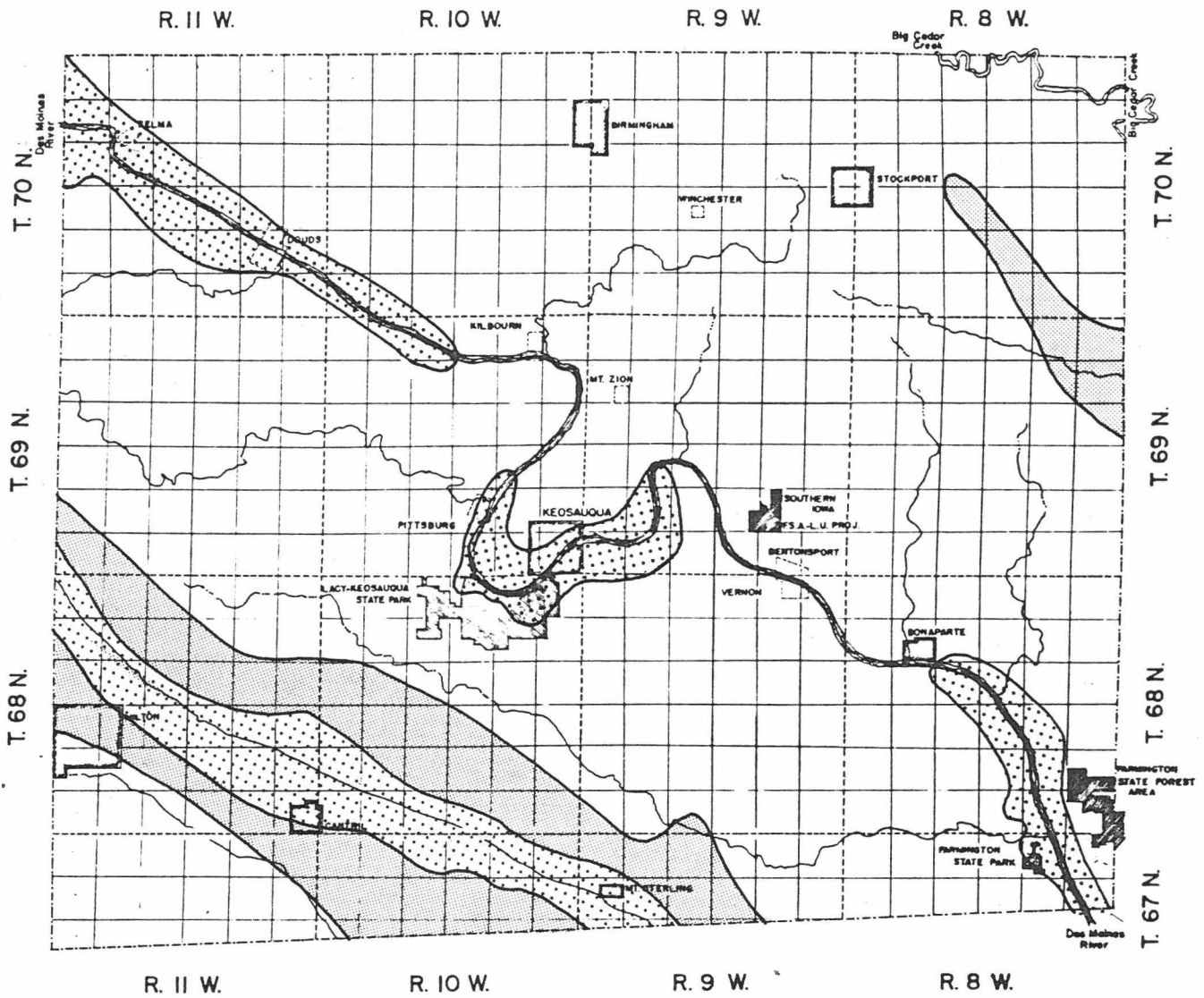
Average yields and water quality characteristics throughout the county for each of the aquifers are also summarized in the maps in Figures 10, 11, 12, 13, 14, and 15.

Table 1

## GEOLOGIC AND HYDROGEOLOGIC UNITS IN VAN BUREN COUNTY

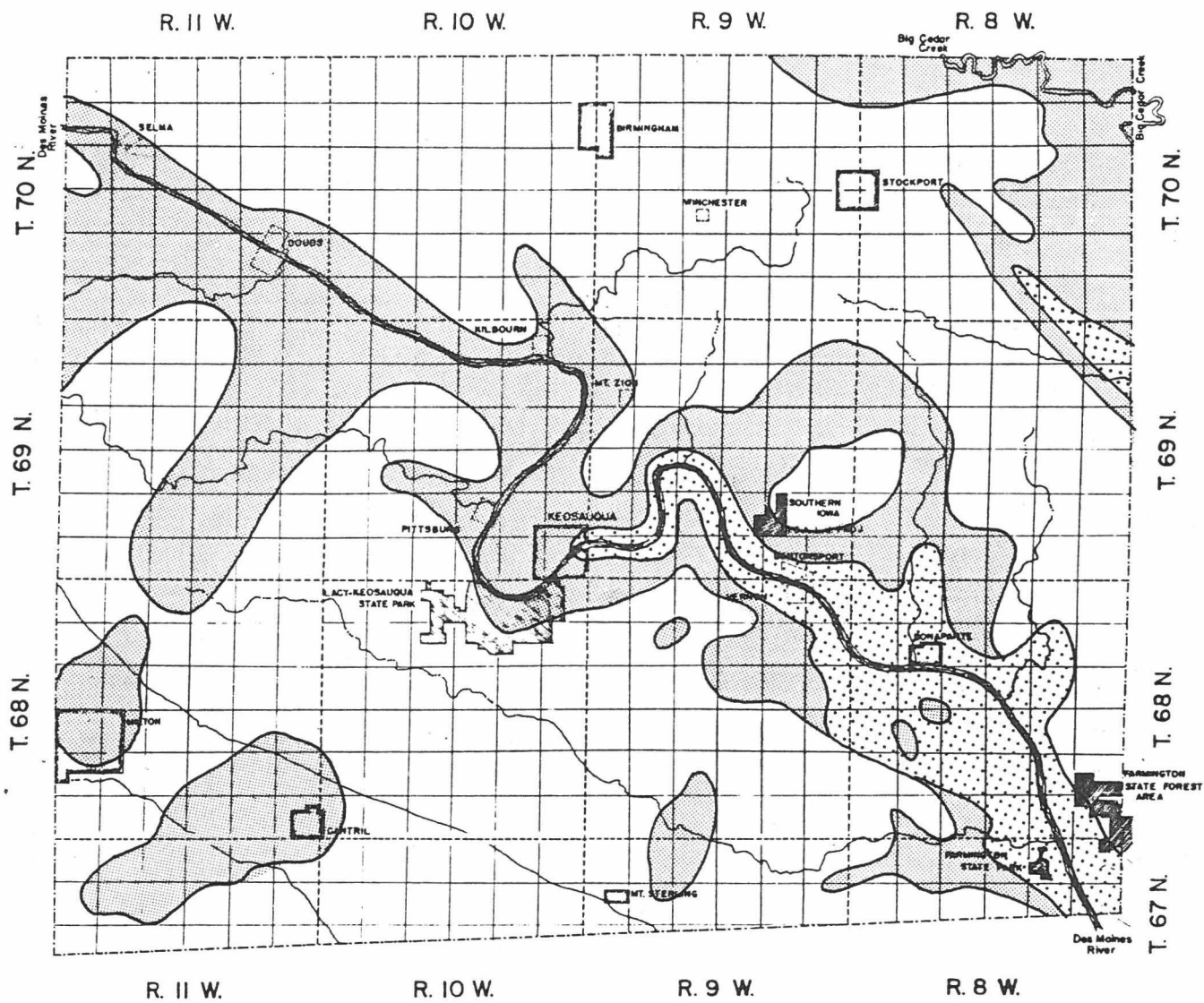
Age	Rock Unit	Description	Thickness Range	Hydrogeologic Unit	Water-Bearing Characteristics
Quaternary	Alluvium	Sand, gravel, silt, and clay		Surficial aquifer	Fair to large yields (25 to 100 gpm)
	Glacial drift (undifferentiated)	Predominantly till containing scattered irregular bodies of sand and gravel	0-300 (feet)		Low yields (less than 10 gpm)
	Buried channel deposits	Sand, gravel, silt and clay			Small to large yields
Pennsylvanian	Cherokee Series	Shale; sandstone, mostly thin	0-100	Aquiclude	Does not yield water
Mississippian	Meramec Series	Limestone, sandy	250-375	Mississippian aquifer	Fair to low yields
	Osage Series	Limestone and dolomite, cherty			
	Kinderhook Series	Limestone, oolitic, and dolomite, cherty			
Devonian	Maple Mill Shale Sheffield Formation Lime Creek Formation	Shale; limestone in lower part	250-325	Devonian aquiclude	Does not yield water
	Cedar Valley Lime- stone; Wapsipinicon Formation	Limestone and dolomite; contains evaporites in southern half of Iowa	150-225	Devonian aquifer	Fair to low yields
Ordovician	Galena Formation	Limestone and dolomite	800-850	Minor aquifer	Low yields
	Decorah Formation Platteville Forma- tion	Limestone and thin shales; includes sandstone in SE Iowa		Aquiclude	Does not yield water
	St. Peter Sandstone	Sandstone		Cambrian-Ordovician aquifer	Fair yields
	Prairie du Chien Formation	Dolomite, sandy and cherty			High yields (over 500 gpm)
Cambrian	Jordan Sandstone	Sandstone	20-30		
	St. Lawrence Formation	Dolomite			
	Franconia Sandstone	Sandstone and shale		Aquitard	No data
	Dresbach Group	Sandstone		Dresbach aquifer	No data
Precambrian	Undifferentiated	Coarse sandstones; crystalline rocks		Base of ground- water reservoir	Not known to yield water

Figure 2  
SURFICIAL MATERIALS



-  Alluvium
-  Glacial Drift
-  Buried Channels

Figure 3  
GEOLOGIC MAP



- Pennsylvanian
- Mississippian-Upper Part
- Mississippian-Lower Part



Figure 4

ELEVATION OF LAND SURFACE IN FEET ABOVE MEAN SEA LEVEL

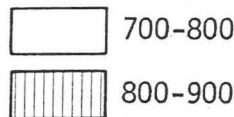
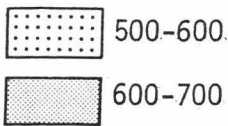
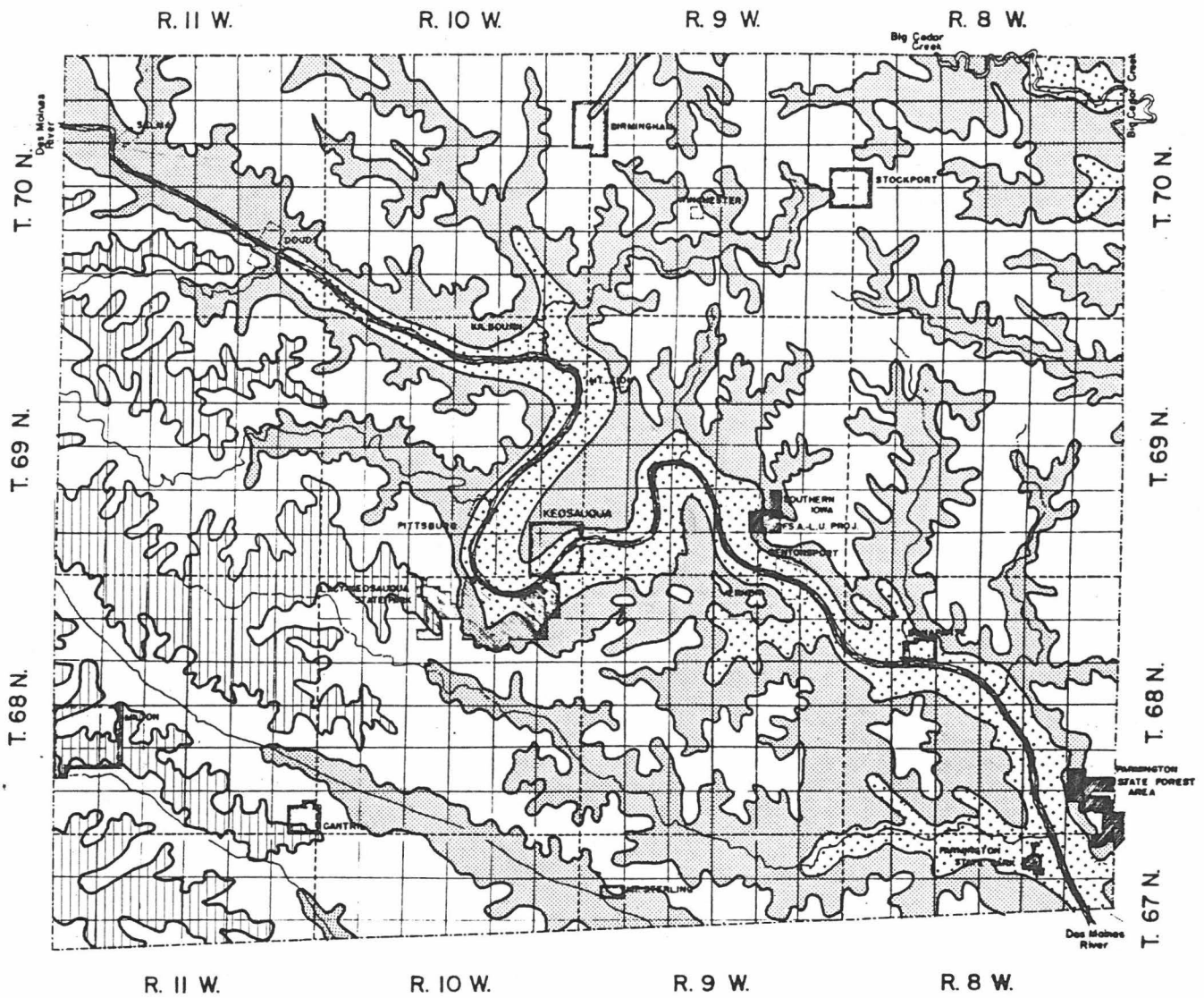
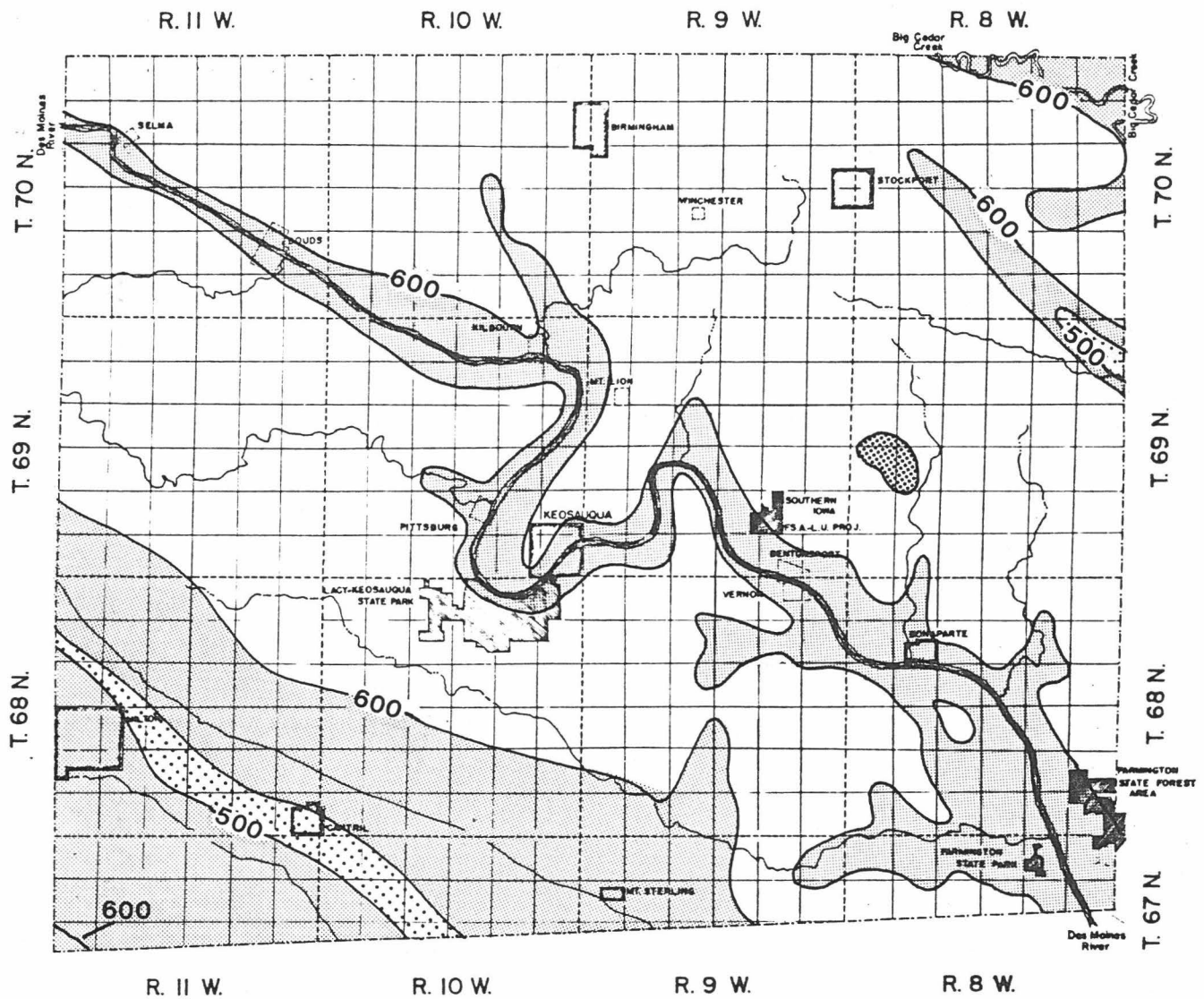
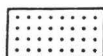



Figure 5

ELEVATION OF BEDROCK SURFACE IN FEET ABOVE MEAN SEA LEVEL



 Below 500  
 500-600

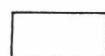

 600-700  
 Over 700

Figure 6

RANGE IN DEPTH TO VAN BUREN COUNTY'S PRINCIPAL ROCK AQUIFERS

	R-11W	R-10W	R-9W	R-8W	
T-70N	BEDROCK 0-200 MISSISSIPPIAN 0-250 DEVONIAN 650-850 CAMBRO-ORDOVICIAN 1800-1900	BEDROCK 50-150 MISSISSIPPIAN 100-250 DEVONIAN 700-850 CAMBRO-ORDOVICIAN 1700-1900	BEDROCK 50-150 MISSISSIPPIAN 100-250 DEVONIAN 700-850 CAMBRO-ORDOVICIAN 1750-1850	BEDROCK 0-150 MISSISSIPPIAN 0-200 DEVONIAN 600-800 CAMBRO-ORDOVICIAN 1700-1900	T-70N
T-69N	BEDROCK 50-250 MISSISSIPPIAN 50-300 DEVONIAN 800-900 CAMBRO-ORDOVICIAN 1800-1900	BEDROCK 0-150 MISSISSIPPIAN 0-200 DEVONIAN 650-900 CAMBRO-ORDOVICIAN 1700-1900	BEDROCK 0-150 MISSISSIPPIAN 0-200 DEVONIAN 500-700 CAMBRO-ORDOVICIAN 1700-1900	BEDROCK 50-150 MISSISSIPPIAN 50-200 DEVONIAN 600-700 CAMBRO-ORDOVICIAN 1750-1850	T-69N
T-68N	BEDROCK 150-250 MISSISSIPPIAN 200-300 DEVONIAN 850-950 CAMBRO-ORDOVICIAN 1800-1900	BEDROCK 50-250 MISSISSIPPIAN 50-300 DEVONIAN 800-900 CAMBRO-ORDOVICIAN 1800-1900	BEDROCK 50-150 MISSISSIPPIAN 50-250 DEVONIAN 600-800 CAMBRO-ORDOVICIAN 1700-1900	BEDROCK 0-150 MISSISSIPPIAN 0-200 DEVONIAN 500-700 CAMBRO-ORDOVICIAN 1700-1900	T-68N
T-67N	BEDROCK 150-250 MISSISSIPPIAN 200-300 DEVONIAN 850-950 CAMBRO-ORDOVICIAN 1800-1900	BEDROCK 50-150 MISSISSIPPIAN 100-250 DEVONIAN 850-950 CAMBRO-ORDOVICIAN 1750-1850	BEDROCK 50-150 MISSISSIPPIAN 50-250 DEVONIAN 800-850 CAMBRO-ORDOVICIAN 1700-1900	BEDROCK 0-150 MISSISSIPPIAN 0-200 DEVONIAN 500-800 CAMBRO-ORDOVICIAN 1700-1900	T-67N
	R-11W	R-10W	R-9W	R-8W	

Figure 7

INDEX MAP FOR TYPICAL WELLS IN VAN BUREN COUNTY

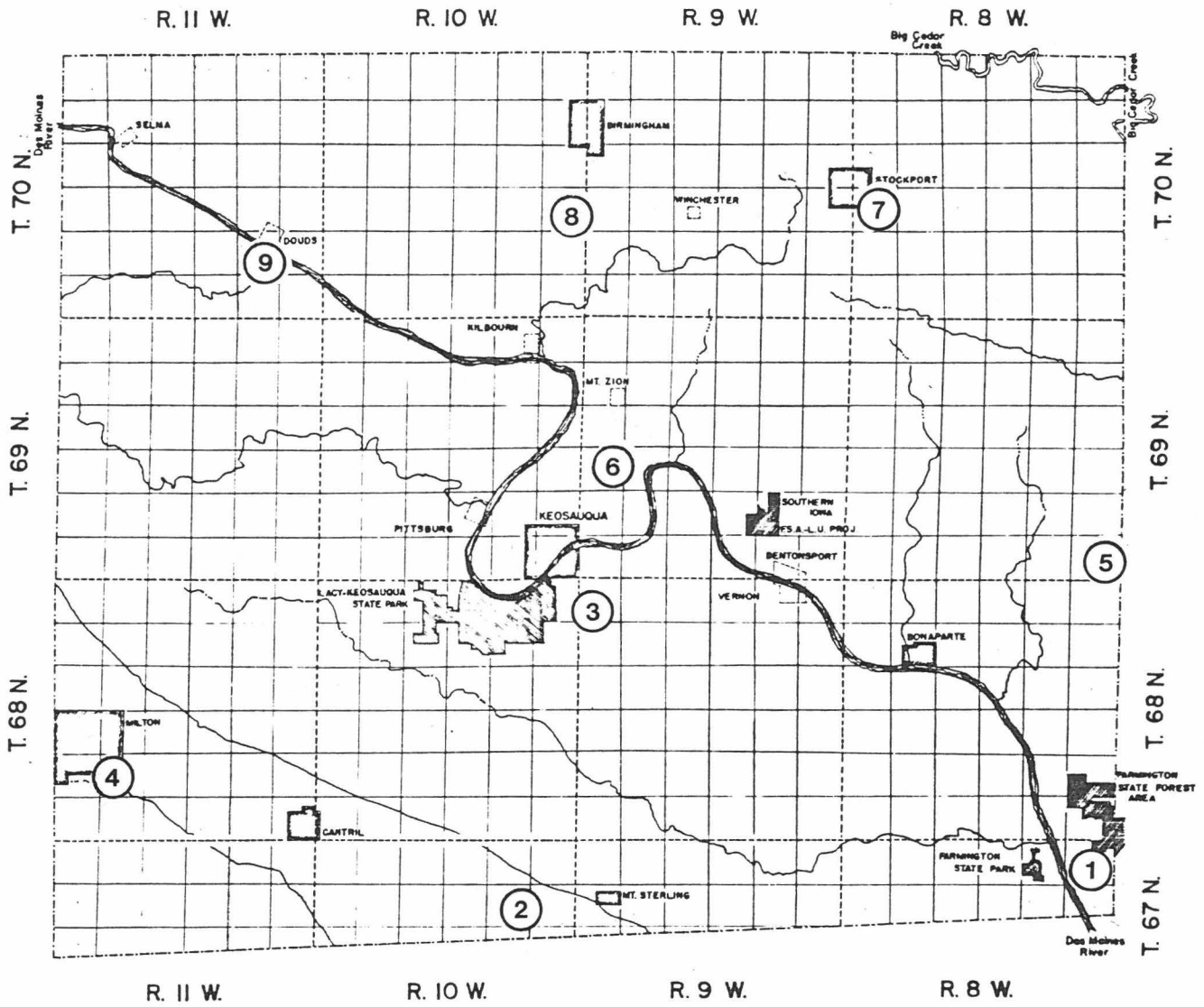


Figure 8

TYPICAL WELLS IN VAN BUREN COUNTY

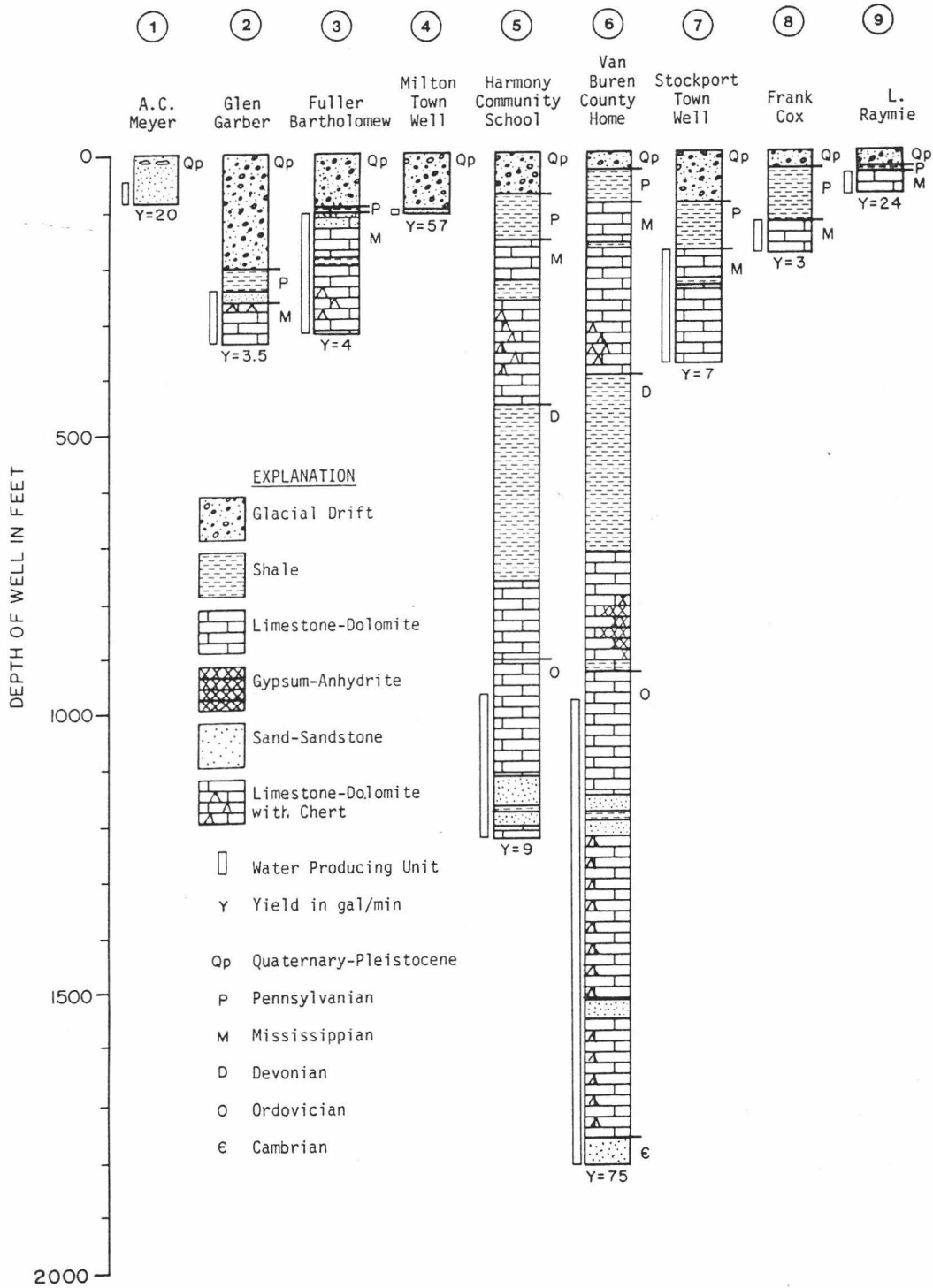
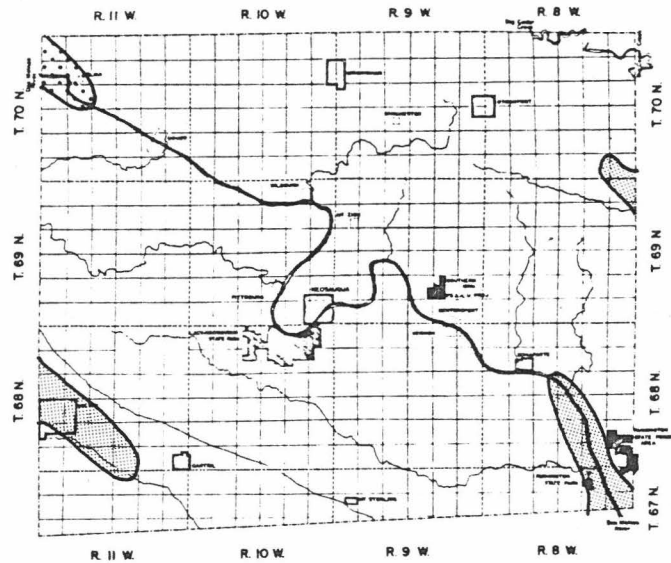


Figure 9  
SURFICIAL AQUIFERS

Water Levels

Water levels in the surficial aquifers are difficult to analyze. Water rises to different levels in wells drilled into alluvial, buried-channel, and drift aquifers. The water table in the drift aquifer generally slopes from high land areas toward the streams and, changes noticeably throughout the year. Levels in drift and buried-channel aquifers respond rapidly to recharge from precipitation. Water levels in the alluvial aquifer fluctuate somewhat in the same way as those in the drift and buried-channel aquifers; however, the main influence on the alluvial aquifer is the stage (level) of the associated streams. Water levels will be high during periods of high stream stage and low during the low-stage periods.

Water levels in the drift aquifers commonly are from 10 to 50 feet below the land surface, and those in the buried-channel aquifer have been reported to be as low as 175 feet below the land surface. The water levels in alluvial wells are from 4 to 20 feet below the flood-plain surface, and the depth to the water surface will be accordingly deeper in wells located on terrace surfaces.



Water yields to wells in gallons per minute

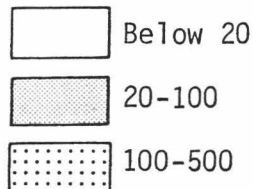
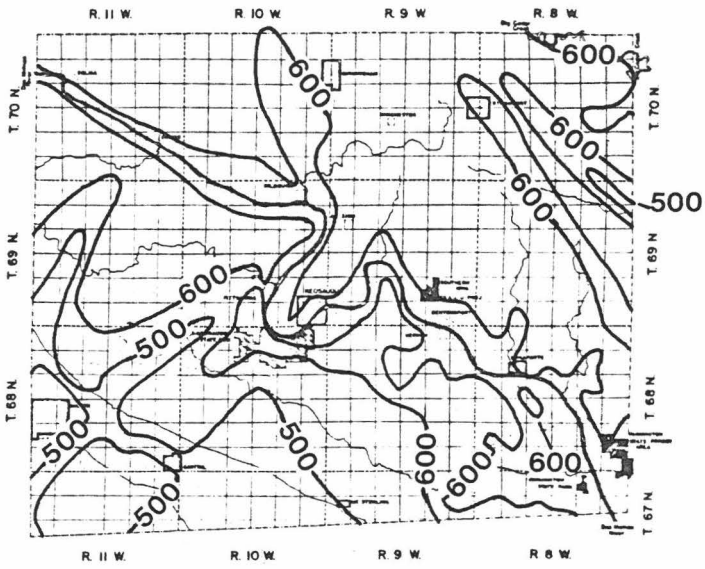
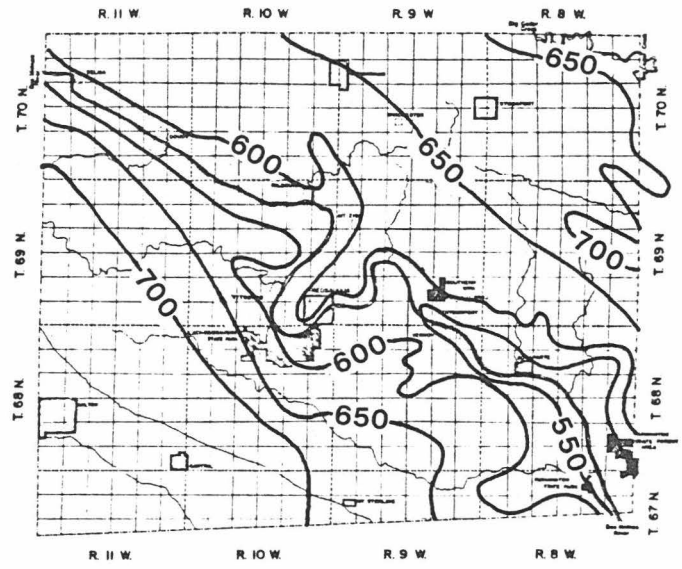


Figure 10

MISSISSIPPIAN AQUIFER

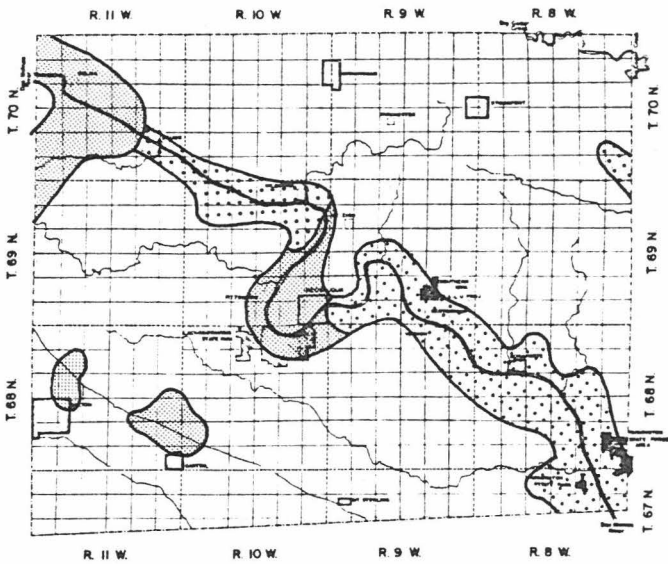


Elevation of Mississippian Aquifer in feet above mean sea level

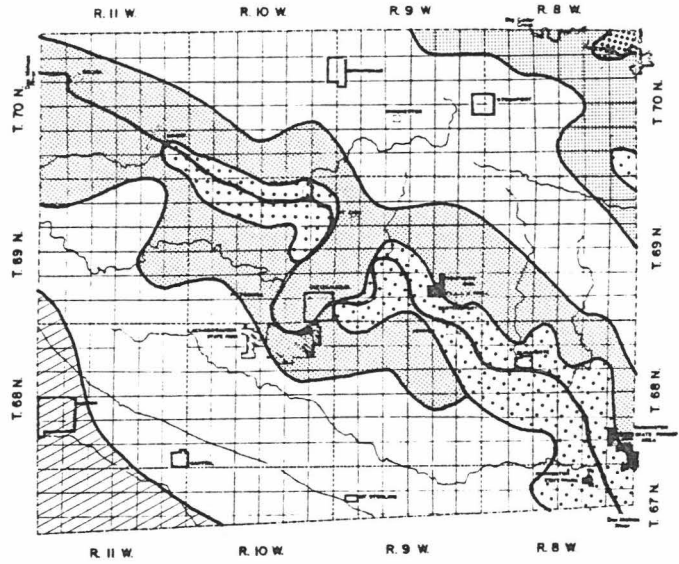
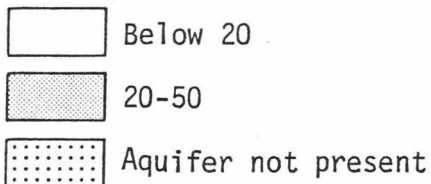


Water levels in wells in feet above mean sea level

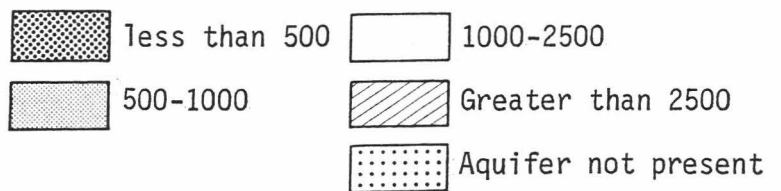
Upper Part



Water yields to wells in gallons per minute



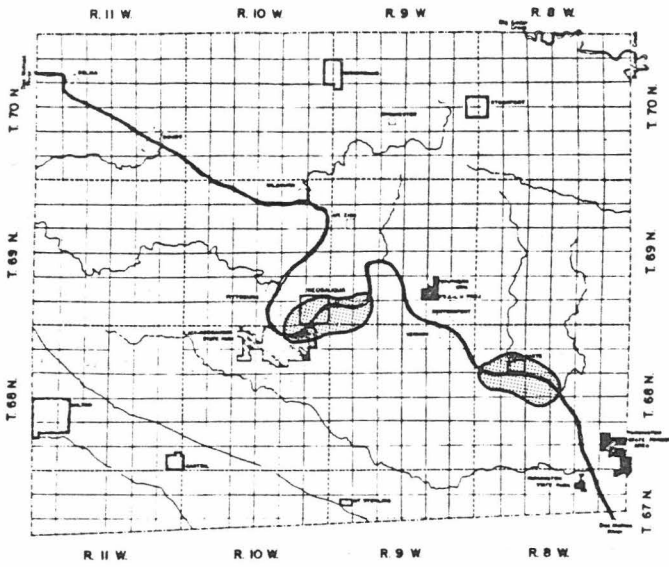
Dissolved solids content in milligrams per liter (mg/l)\*



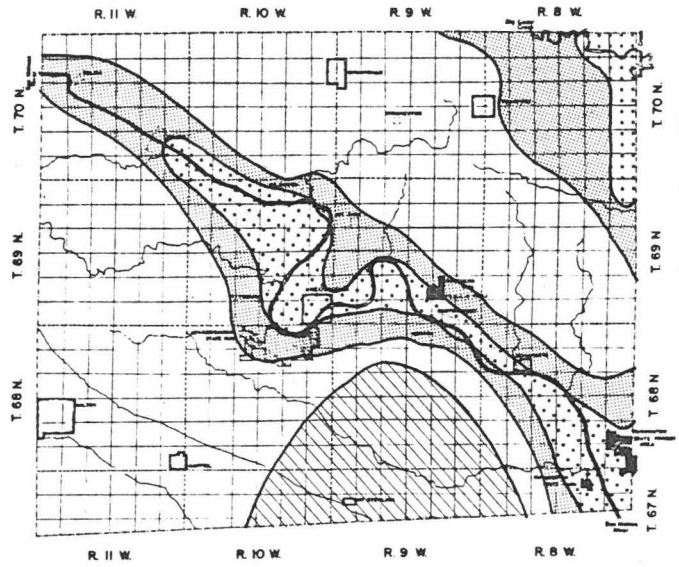
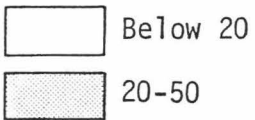
\*Other water quality data in Figure 13

# MISSISSIPPIAN AQUIFER

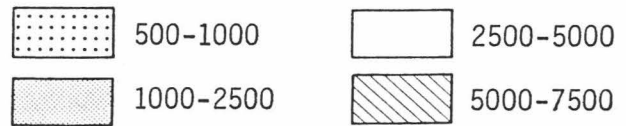
## Lower Part



Water yields to wells in gallons per minute



Dissolved solids content in milligrams per liter (mg/l)\*

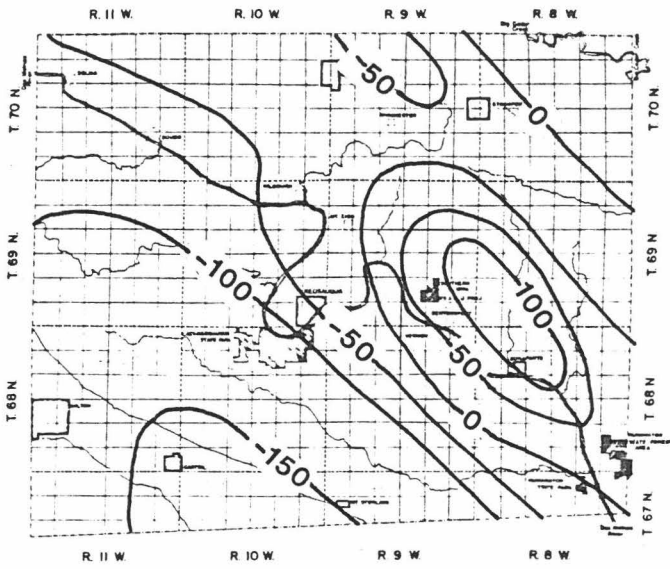


\*Other water quality data in Figure 14

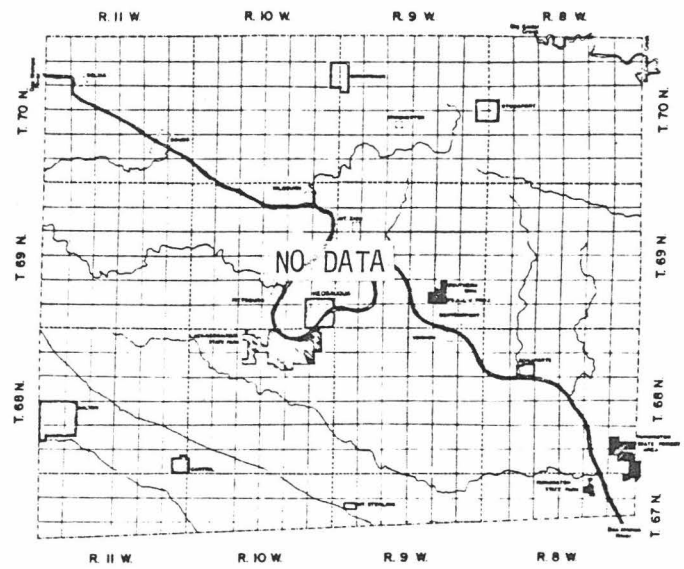


Figure 11

DEVONIAN AQUIFER

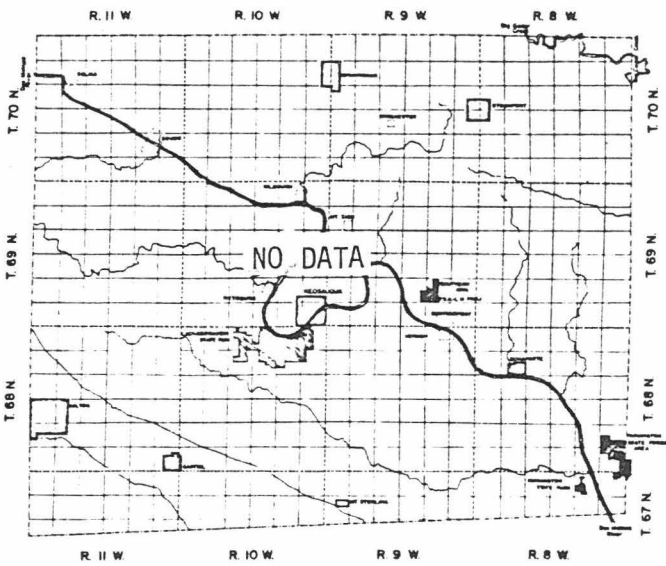


Elevation of Devonian Aquifer in feet above mean sea level



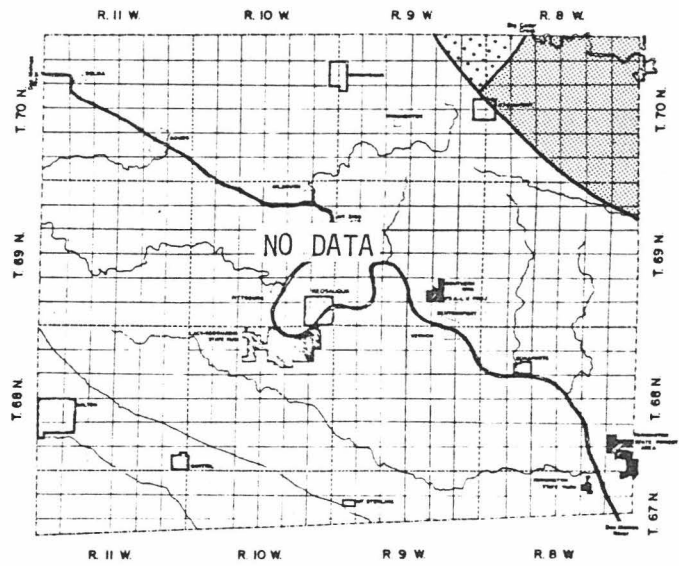
Water levels in wells in feet above mean sea level

Estimated to be approximately 650 feet above mean sea level

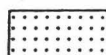


Water yields to wells in gallons per minute

Estimated to be below 20 gallons per minute



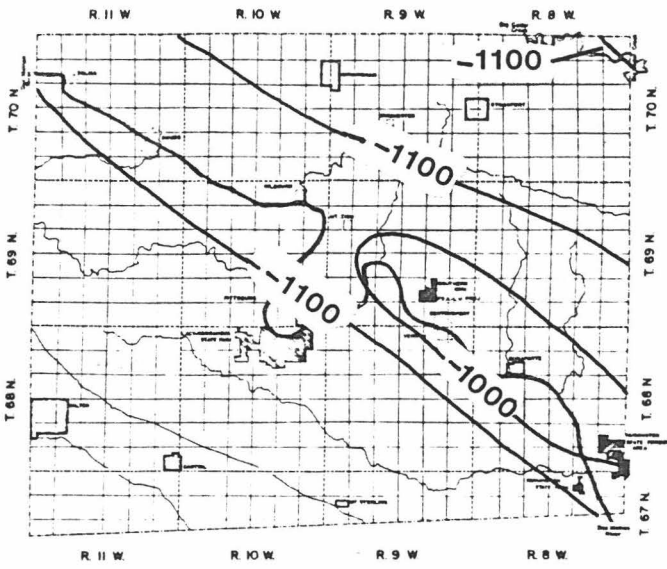
Dissolved solids content in milligrams per liter (mg/l)\*

-  5000-7500
-  Over 10,000

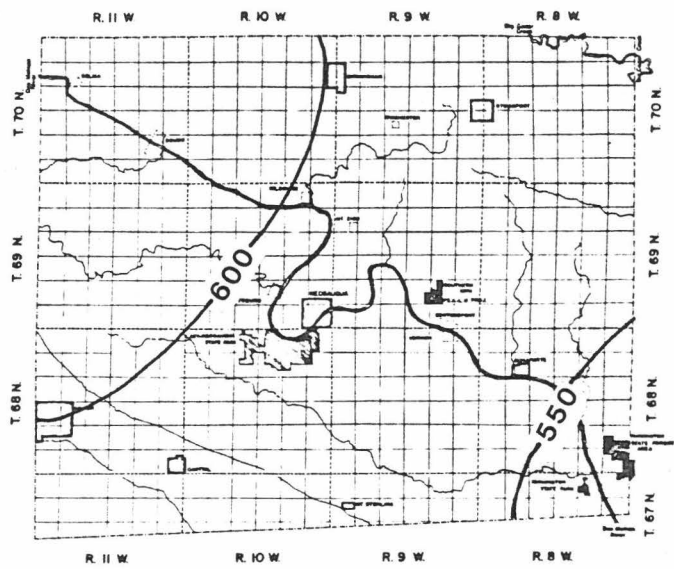
\*Other water quality data in Figure 15

Figure 12

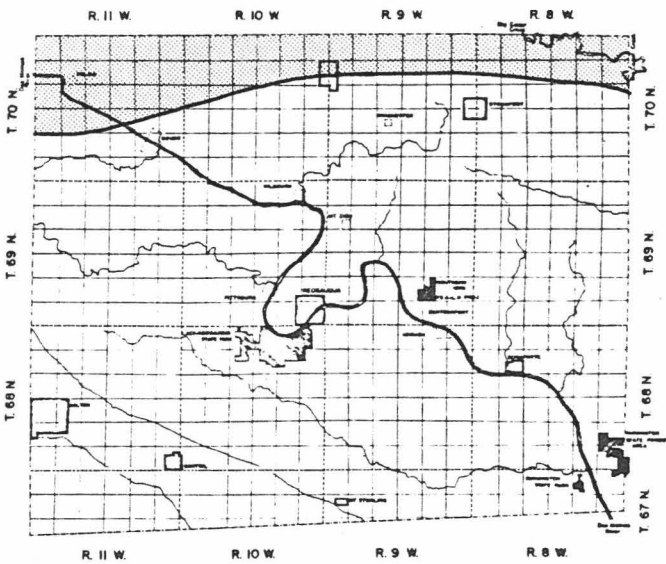
CAMBRO-ORDOVICIAN (JORDAN) AQUIFER



Elevation of Jordan aquifer in feet above sea level

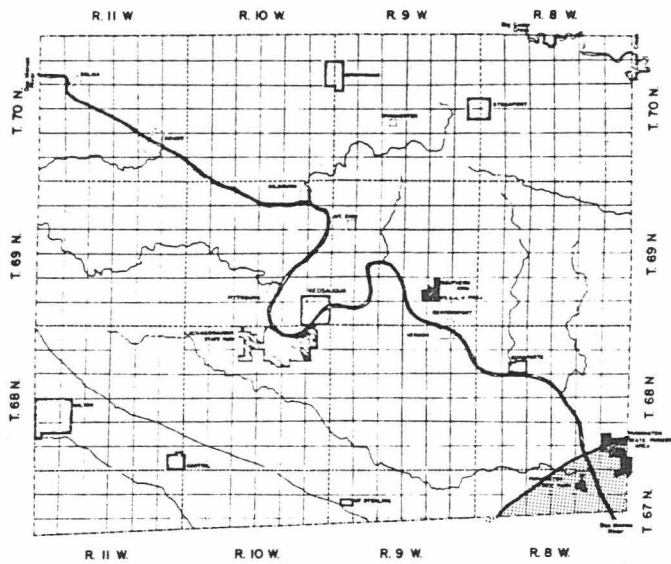


Water levels in wells in feet above mean sea level



Water yields to wells in gallons per minute

- 500-1000
- More than 1000



Dissolved solids content in milligrams per liter (mg/l)\*

- 1000-1500
- Greater than 1500

\*Other water quality data in Figure 15

Table 2

## SIGNIFICANCE OF MINERAL CONSTITUENTS AND PHYSICAL PROPERTIES OF WATER

Constituent or Property	Maximum Recommended Concentration	Significance
Iron (Fe)	0.3 mg/l	Objectionable as it causes red and brown staining of clothing and porcelain. High concentrations affect the color and taste of beverages. Iron is not listed in the following tables, as there are often major differences between reported and actual concentrations. It may be added to water from well casings, pumps, and pipes. The concentration also is affected by micro-organisms. Special sampling and analytical techniques are needed for accurate study.
Manganese (Mn)	0.05 mg/l	Objectional for the same reason as iron. When both iron and manganese are present, it is recommended that the total concentration not exceed 0.3 mg/l. Micro-organisms also affect the concentration. Special techniques are needed for an accurate study.
Calcium (Ca) and Magnesium (Mg)		Principal causes for hardness and scale-forming properties of water. They reduce the lathering ability of soap.
Sodium (Na) and Potassium (K)		Impart a salty or brackish taste when combined with chloride. Sodium salts cause foaming in boilers.
Sulfate (SO <sub>4</sub> )	250 mg/l	Commonly has a laxitive effect when the concentration is 600 to 1,000 mg/l; particularly when combined with magnesium or sodium. The effect is much less when combined with calcium. This laxitive effect is commonly noted by newcomers, but they become acclimated to the water in a short time. The effect is noticeable in almost all persons when concentrations exceed 750 mg/l. Sulfate combined with calcium forms a hard scale in boilers and water heaters.
Chloride (Cl)	250 mg/l	Large amounts combined with sodium impart a salty taste.
Fluoride (F)	2.0 mg/l	Concentrations of 0.8 to 1.3 mg/l are considered to play a part in the reduction of tooth decay. However, concentrations over 2.0 mg/l will cause the mottling of the enamel of children's teeth.
Nitrate (NO <sub>3</sub> )	45 mg/l	Waters with high nitrate content should not be used for infant feeding as it may cause methemoglobinemia or cyanosis. High concentrations suggest organic pollution from sewage, decayed organic matter, nitrate in the soil, or chemical fertilizer. High nitrates in the natural waters of central Iowa are limited to isolated occurrences, usually from shallow dug wells on farms. Since the high concentrations are characteristic of individual wells and not of any one aquifer, nitrate will not be discussed in this report.
Dissolved Solids	500 mg/l	This refers to all of the material in water that is in solution. It affects the chemical and physical properties of water for many uses. Amounts over 2,000 mg/l will have a laxitive effect on most persons. Amounts up to 1,000 mg/l are generally considered acceptable for drinking purposes if no other water is available.
Hardness (as CaCO <sub>3</sub> )		This affects the lathering ability of soap. It is generally produced by calcium and magnesium. Hardness is expressed in parts per million equivalent to CaCO <sub>3</sub> as if all the hardness were caused by this compound. Water becomes objectionable for domestic use when the hardness is above 100 mg/l; however, it can be treated readily by softening.
Temperature		Affects the desirability and economy of water use, especially for industrial cooling and air conditioning. Most users want a water with a low and constant temperature.
Suspended Sediment		Causes water to have a cloudy or muddy appearance. It must be settled or filtered out before the water is used. It is the material that "silts-up" reservoirs, and it is the major cause of the reduction of reservoir life.

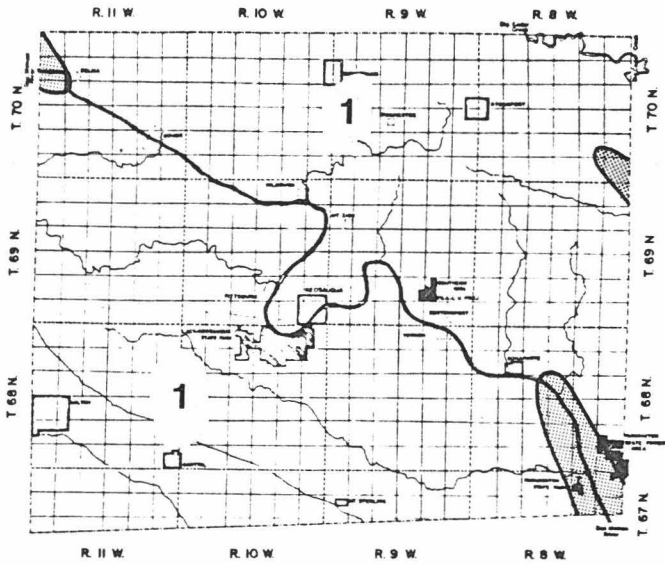
To the user, the quality of ground water is as important as the amount of water that an aquifer will yield. As ground water moves through soil and rock materials, it dissolves some of the minerals which, in turn, affect water quality. In addition to mineral content, bacterial and chemical contamination may be introduced through poorly constructed wells and seepage from other pollution sources.

Recommended standards for common water constituents are described in the table above. These are rationally accepted as guidelines for acceptable drinking water supplies. Limits for uses other than drinking often differ from these. For instance, water that is unacceptable for drinking and household use may be completely satisfactory for industrial cooling.

From analyses of ground water averages (A) and ranges (R) of values in milligrams per liter (mg/l) for several constituents are summarized in Tables 3 and 4 for the surficial and bedrock aquifers in Van Buren County. Recommended concentrations for some constituents are often exceeded without obvious ill effects, although the water may be unpalatable. Water quality analyses for individual wells should be obtained to determine if concentrations of constituents that affect health are exceeded.

Figure 13  
CHEMICAL CHARACTER OF GROUND WATER

Surficial Aquifers

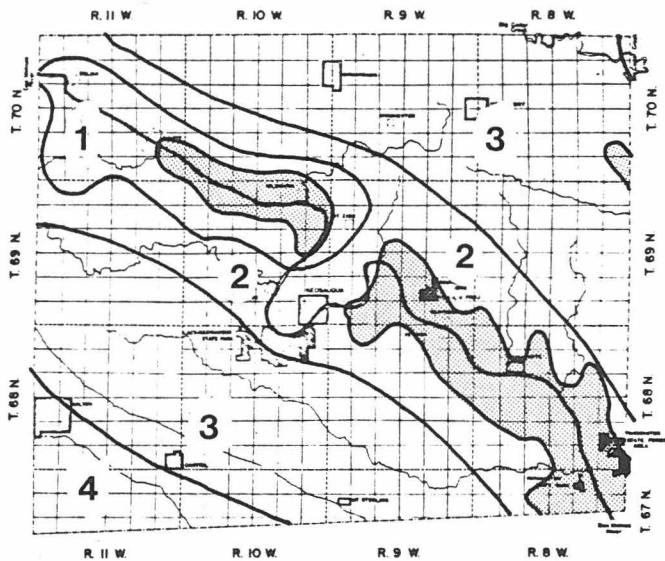


Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium Potassium (Na+K)	Bicarbonate (HCO <sub>3</sub> )	Sulfate (SO <sub>4</sub> )	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as CaCO <sub>3</sub> )
1	A	103	36	40	Drift Aquifer					
	R	77-150	24-61	13-88	304-739	11-393	0.5-7.5	0.2-0.7	357-982	292-625
2	A	108	31	21	Alluvial Aquifer					
	R	84-143	20-41	11-28	295	169	3.5-17	.4	440-726	304-525
3	A	53	20	153	Buried-Channel Aquifer					
	R	24-88	3.1-39	87-244	579	46	2.5-28	0-1.0	357-715	78-360

The alluvial and drift aquifers yield good quality water. Water from buried channels has a higher dissolved solids content, but it is less than that from rock aquifers. Water temperatures average 54°F and normally do not vary more than 6°F.

Mississippian Aquifer

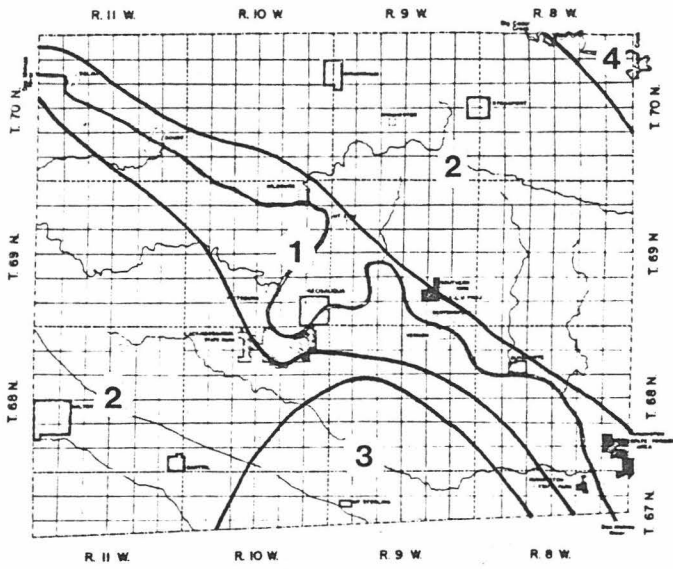
Upper Part



Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium Potassium (Na+K)	Bicarbonate (HCO <sub>3</sub> )	Sulfate (SO <sub>4</sub> )	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as CaCO <sub>3</sub> )
1	A	104	32	43	472	91	4.5	0.4	537	399
	R	66-156	19-50	8.8-118	322-802	5.8-240	0.5-24	0-1.8	345-737	315-545
2	A	155	46	84	423	397	9.9	.6	955	578
	R	104-210	21-61	8-177	349-498	321-480	.5-35	0-1.8	793-1050	385-660
3	A	299	92	125	281	1100	8.5	.8	1950	1130
	R	197-547	47-124	33-214	44-434	750-1590	2-22	.3-1.6	1420-2740	723-1560
4	A	327	76	679	284	2160	68	1.4	3620	1130
	R	154-490	54-99	474-1080	212-350	1500-2850	27-186	1.2-2.4	2710-4880	605-1520

Water in the upper part of the Mississippian aquifer is more highly mineralized than that typically found in the surficial aquifers and is usually very hard. With exception of the areas adjacent to the Des Moines River, dissolved solids content is high, particularly in the northeastern and southwestern portions of the county. For most of the county, sulfate concentrations exceed recommended standards. Average water temperature is 55°F with a range of 51°-60°F.

Figure 14  
Mississippian Aquifer  
Lower Part



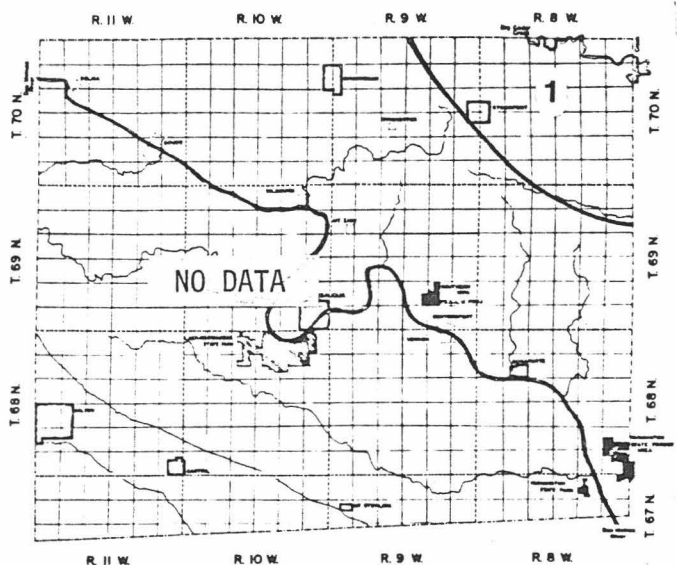
Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium Potassium (Na+K)	Bicarbonate (HCO <sub>3</sub> )	Sulfate (SO <sub>4</sub> )	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as CaCO <sub>3</sub> )
1	A	152	44	79	432	322	45	.4	964	565
	R	132-174	28-98	17-162	334-608	84-658	13-93	.2-.7	566-1610	466-840
2	A	102	52	718	459	1340	169	2.6	2710	469
	R	35-193	15-107	451-994	266-595	920-1660	19-365	.5-6	2220-3250	148-891
3	A	215	82	1375	956	2520	610	1.8	5280	875
	R	102-452	58-120	1040-1630	378-3640	1800-4020	290-930	1.4-3	4070-7330	497-1620
4	A	78	38	276	592	354	57	1.6	1110	355
	R	38-128	19-68	143-489	465-754	260-560	.5-150	.5-2.5	879-1480	176-581

Water in the lower part of the Mississippian aquifer is generally of poorer quality than that found in the upper part. Throughout the county the water is exceptionally hard, and with the exception of the area adjacent to the Des Moines River, greatly exceeds recommended standards for sulfate and total dissolved solids. Average water temperature is 55°F with a range of 51°-60°F.

Figure 15

CHEMICAL CHARACTER OF GROUND WATER

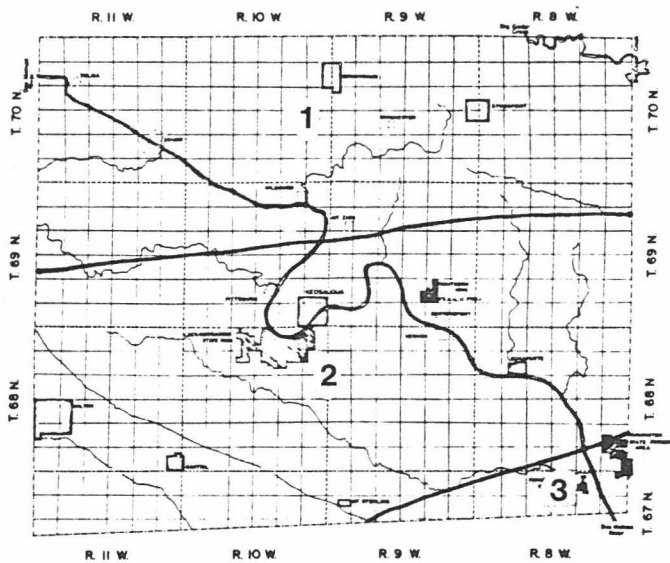
Devonian Aquifer



Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium Potassium (Na+K)	Bicarbonate (HCO <sub>3</sub> )	Sulfate (SO <sub>4</sub> )	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as CaCO <sub>3</sub> )
1 A		492	138	2220	289	4580	1100	2.8	9570	1800
1 R		441-617	97-155	1850-2700	183-378	4030-5000	550-1740	1.9-5.0	8240-11,100	1580-2180

Water in the Devonian aquifer is highly mineralized and objectionably hard wherever it has been encountered in the county. Because of extremely high total dissolved solids, primarily sodium-potassium, chloride, and sulfate, the water from the Devonian is unfit for human or animal consumption. The average water temperature is 60°F and ranges between 54° and 64°F in southeast Iowa.

Cambro-Ordovician Aquifer



Area	Average and Range	Calcium (Ca)	Magnesium (Mg)	Sodium Potassium (Na+K)	Bicarbonate (HCO <sub>3</sub> )	Sulfate (SO <sub>4</sub> )	Chloride (Cl)	Fluoride (F)	Dissolved Solids	Hardness (as CaCO <sub>3</sub> )
1 A		84	34	267	298	476	124	1.7	1160	349
1 R		78-92	26-41	247-283	183-317	455-500	100-148	1.0-2.2	1100-1220	322-388
2 A		71	30	322	304	394	216	2.4	1210	300
2 R		65-76	25-32	310-331	273-326	338-420	209-222	2.0-2.9	1080-1310	290-321
3* A		97	39	430	300	440	440	3.2	1650	403
3* R										

\* based on one analysis

This deep aquifer yields water of fair to poor quality. The water is noticeably hard and generally exceeds recommended standards for sulfate and dissolved solids. Water temperatures are higher than other rock aquifer sources, averaging 72°F and ranging up to about 75°F.

## RECOMMENDATIONS FOR PRIVATE WATER WELLS

### Contracting for Well Construction

To protect your investment and guarantee satisfactory well completion, it is a good idea to have a written agreement with the well driller. The agreement should specify in detail:

- size of well, casing specifications, and types of screen and well seal

- methods of eliminating surface and subsurface contamination

- disinfection procedure to be used

- type of well development if necessary

- test pumping procedure to be used

- date for completion

- itemized cost list including charges for drilling per foot, for materials per unit, and for other operations such as developing and test pumping

- guarantee of materials, workmanship, and that all work will comply with current recommended methods

- liability insurance for owner and driller

### Well Location

A well should be located where it will be least subject to contamination from nearby sources of pollution. The Iowa State Department of Health recommends minimum distances between a new well and pollution sources, such as cesspools (150 ft.), septic tanks (50 ft.), and barnyards (50-100 ft. and down slope from well). Greater distances should be provided where possible.

The well location should not be subject to flooding or surface water contamination. Select a well-drained site, extend the well casing a few feet above the ground, and mound earth around it. Diversion terraces or ditches may be necessary on slopes above a well to divert surface runoff around the well site.

In the construction of all wells care should be taken to seal or grout the area between the well bore and the well casing (the annulus) as appropriate so that surface water and other pollutants cannot seep into the well and contaminate the aquifer.

## Water Treatment

Water taken from a private well should ideally be tested every six months. The University Hygienic Laboratory will do tests for coliform bacteria, nitrate, iron, hardness, and iron bacteria in drinking water for private individuals. Special bottles must be used for collecting and sending water samples to the laboratory. A sample kit can be obtained by writing to the University Hygienic Laboratory, University of Iowa, Oakdale Campus, Iowa City, Iowa 52242. Indicate whether your water has been treated with chlorine, iodine, or bromine; as different sample bottles must be used for treated and untreated water. The charge for the bacterial test is \$3.00; for iron, hardness, and nitrate, it is \$3.00; and for iron bacteria, \$5.00. If your well is determined to be unsafe, advice for correcting the problem can be obtained from your county or state Department of Health. Several certified private laboratories also run water analyses.

Shock chlorination is recommended following the construction and installation of a well and distribution system and any time these are opened for repairs or remodeling. A strong chlorine solution is placed in the well and complete distribution system to kill nuisance and disease-causing organisms. If the first shock chlorination does not rid the water supply of bacteria it should be repeated, if this does not solve the problem the well should be abandoned or the water should be continuously disinfected with proper chlorination equipment.

Since most of the ground waters in Van Buren County are mineralized, water softening and iron removal equipment may make water more palatable and pleasant to use. Softened water contains increased sodium; contact your physician before using a softener if you are on a sodium-restricted diet. Chlorination followed by filtration will remove most forms of iron and iron bacteria. Iron bacteria has no adverse effect on health but will plug wells, water lines, and equipment and cause tastes and odors. Iron removal equipment can be used if problems persist.

## Well Abandonment

Wells taken out of service provide easy access for pollution to enter aquifers supplying water to other wells in the vicinity. Unprotected wells may also cause personal injury. Proper abandonment procedures should be followed to restore conditions to those that existed before the well was constructed, and to prevent source contamination. Permanent abandonment requires careful sealing. The well should be filled with concrete, cement grout, or sealing clays throughout its entire length. Before dug or bored wells are filled at least the top 10 feet of lining should be removed so surface waters will not penetrate the subsurface through a porous lining or follow cracks in or around the lining. The site should be completely filled and mounded with compacted earth.

ABANDONED WELLS SHOULD NEVER BE USED FOR DISPOSAL OR SEWAGE OR OTHER WASTES.



## SOURCES OF ADDITIONAL INFORMATION

In planning the development of a ground water supply or contracting for the drilling of a new well additional or more specific information is often required. This report section lists several sources and types of additional information.

### State Agencies That May Be Consulted

Iowa Geological Survey <sup>1</sup>	123 North Capitol Iowa City 52242	(319) 338-1173
State Health Department <sup>2,6</sup>	Lucas Building Des Moines 50319	(515) 281-5787
Iowa Natural Resources Council <sup>3</sup>	Wallace Building Des Moines 50319	(515) 281-5914
Iowa Dept. of Environ. Quality <sup>4</sup>	Wallace Building Des Moines 50319	(515) 281-8854
University Hygienic Laboratory <sup>5</sup>	U. of IA, Oakdale Campus Iowa City 52242	(319) 353-5990
Cooperative Extension Service in Agriculture and Home Economics <sup>6</sup>	110 Curtis Hall, ISU Ames 50011	(515) 294-4569

### Functions:

- <sup>1</sup> Geologic and ground water data repository, consultant on well problems, water development and related services.
- <sup>2</sup> Drinking water quality, public and private water supplies.
- <sup>3</sup> Water withdrawal regulation and Water Permits for well withdrawing more than 5000 gpd.
- <sup>4</sup> Municipal supply regulation and well construction permits
- <sup>5</sup> Water quality analysis
- <sup>6</sup> Advice on water systems design and maintenance

## Well Drillers and Contractors

The listing provided here was drawn from an Iowa Geological Survey mailing list, those selected are within a radius of 50 miles of Van Buren County. For a state-wide listing contact either the Iowa Water Well Drillers Association, 4350 Hopewell Ave., Bettendorf, Iowa 51712, (319) 355-7528 or the Iowa Geological Survey.

Bailey Well Co.  
203 East Main  
New London, Iowa 52645

Schlicher Bros. Well  
Hwy. 34 West  
Fairfield, Iowa 52556

Detrick Well Co.  
R.R. #1  
New London, 52645

Schlicher Well Co.  
P.O. Box 207  
Donnelson, Iowa 52625

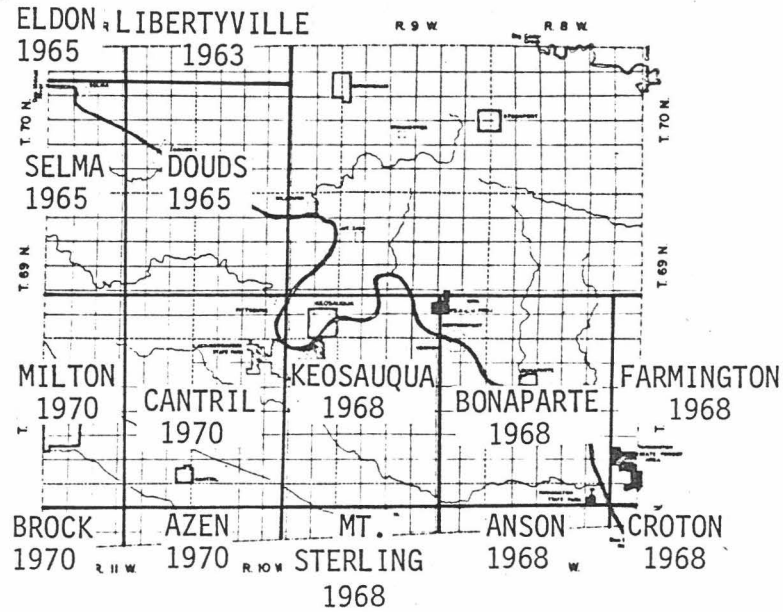
Kramer Well Co.  
Mt. Pleasant, Iowa 52641

Schmeiser Well Co.  
1111 Hageman St.  
Burlington, Iowa 52601

Lyon Well Co.  
Salem, Iowa 52649

Wilson Well Co.  
R.R. #3  
Burlington, Iowa 52601

Topographic Maps (Available from the Iowa Geological Survey)



<u>Map Title</u>	<u>Date</u>	<u>Scale</u>	<u>Contour Interval</u>
Eldon	1965	1:24,000	10 ft.
Libertyville	1963	1:24,000	10 ft.
Selma	1965	1:24,000	10 ft.
Douds	1965	1:24,000	10 ft.
Milton	1970	1:24,000	10 ft.
Cantril	1970	1:24,000	10 ft.
Keosauqua	1968	1:24,000	10 ft.
Bonaparte	1968	1:24,000	10 ft.
Farmington	1968	1:24,000	10 ft.
Brock	1970	1:24,000	10 ft.
Azen	1970	1:24,000	10 ft.
Mt. Sterling	1968	1:24,000	10 ft.
Anson	1968	1:24,000	10 ft.
Croton	1968	1:24,000	10 ft.

### Useful Reference Materials

- Coble, R.W., and Roberts, J.V., 1971, The water resources of Southeast Iowa, Iowa Geological Survey, Water Atlas No. 4.
- Horick, P.J., and Steinhilber, W.L., 1973, Mississippian aquifer of Iowa, Iowa Geological Survey, Misc. Map Series No. 3.
- Horick, P.J., and Steinhilber, W.L., 1978, Jordan aquifer of Iowa, Iowa Geological Survey, Misc. Map Series No. 6.
- Iowa State Department of Health, 1971, Sanitary standards for water wells, State Department of Health, Environmental Engineering Service.
- Van Eck, O.J., 1971, Optimal well plugging procedures, Iowa Geological Survey Public Information Circular No. 1.
- Van Eck, O.J., 1978, Plugging procedures for domestic wells, Iowa Geological Survey, Public Information Circular No. 11.