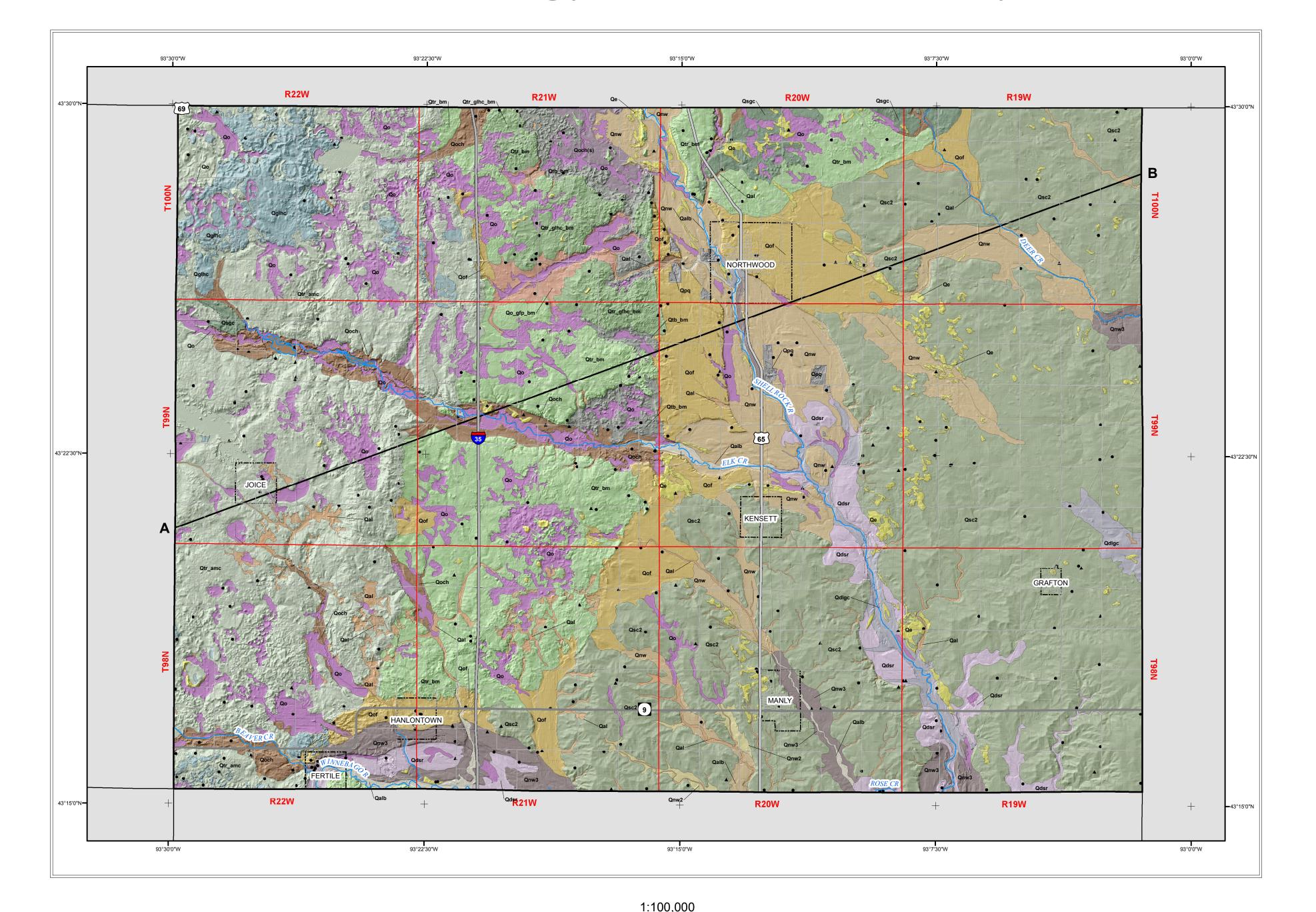
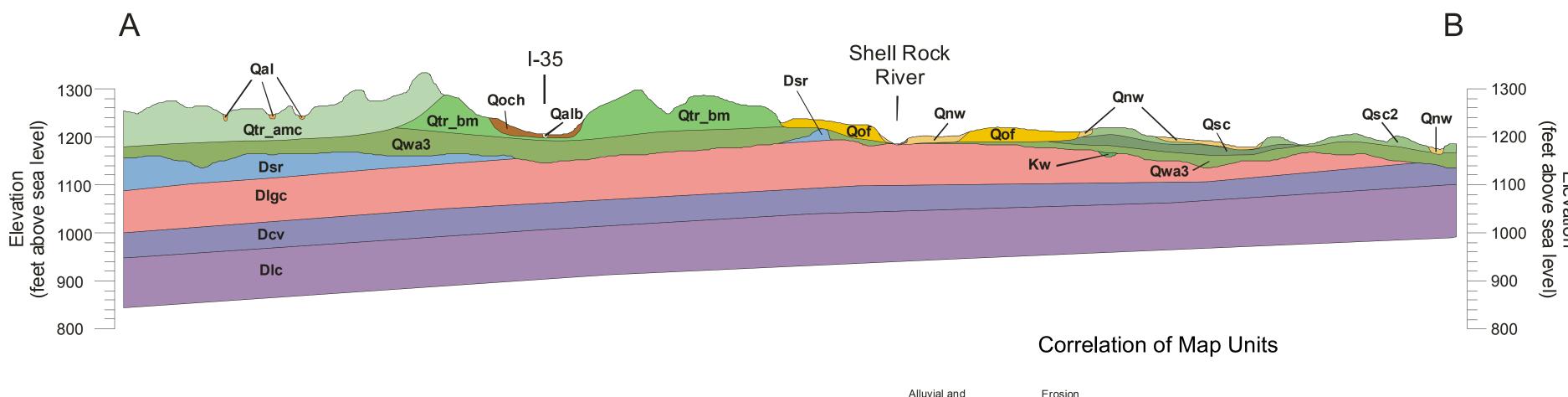
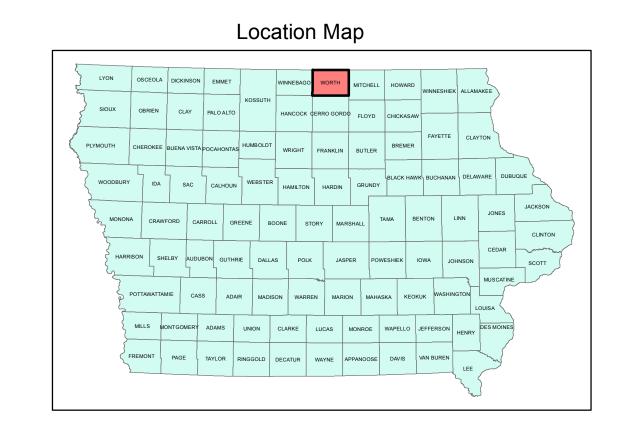
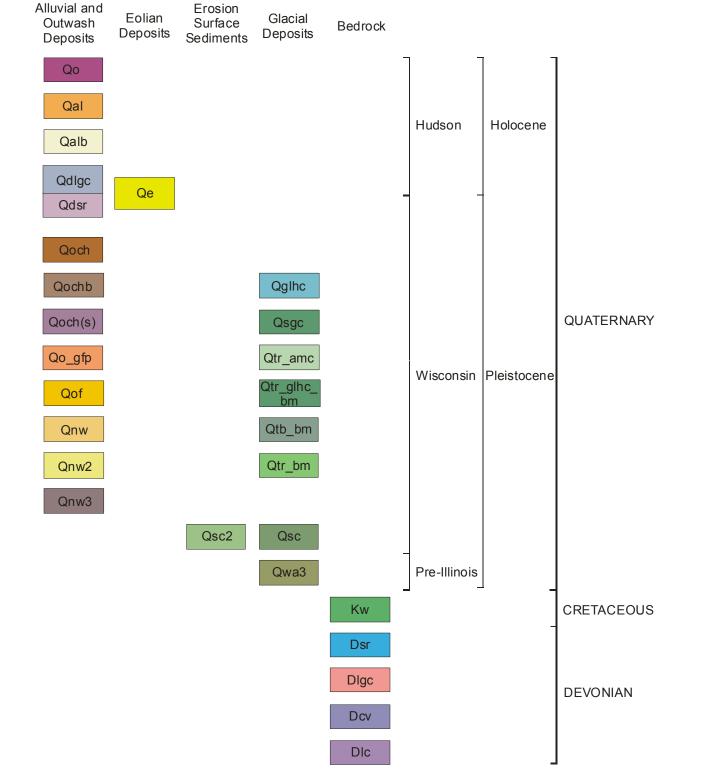
Surficial Geology of Worth County, Iowa



GEOLOGIC CROSS-SECTION A-B







LEGEND

CENOZOIC

QUATERNARY SYSTEM HUDSON EPISODE

- Qo De pre ssi ons (DeFore st Formation-Woden Mbr.). Generally 2.5 to 6 m (8-20 ft) of black to very dark gray, calcareous, muck, peat, silty clay loam col lu viums, and organ ic sedi ments in draine d and undrained closed and semi-close d depressions. Overlies gray, ca lcareous, I cam dia mi cton (Dows Fm.-Morgan/Alden Mbr.) or Noah Creek Fm sand and gravel or Sheldon Creek Fm loam diamicton. Associated with lowrelief featurest hat occupy depressions and low sags on the landscape. Supports wetland vegetation and can be permanently covered by water. High water table.
- Qal Alluvium (DeForest Formation-Undifferentiated). Variable thickness of less than 1 to 5 m (3-16 ft) of very dark gray to brown, noncal care ous to cal care ou s, massive to stratified silty clay loam, clay loam, loam to sandy loam all uvium and colluvium in stream valleys, on hill slopes, and in closed depressions. May overlie Noah Creek Formation. Wolf Creek or Alburnett formations, or fractured Devonian carbonate bedrock. Associated with low-relief modern flood plain, closed depressions, modern draina geways or toeslope positions on the landscape. Seasonal high water tableand potential for frequent
- Qalb Alluvium Shallowto Bedrock (DeFore st Formation-Undifferentiated). Variable thickness of less than 1 to 5 m (3-16 ft) of very dark gray to brown, no neal careous to cal careous, stratified silt ye lay loam, clay loam, loam to sandy loam all uvium and col luvium in stream valleys, on hillslopes and inclosed depressions May overlie Noah Creek Formation or Devonian carbonate bedrock. Bedrock surface is within 5 m (16 ft) of the land surface. Associated with low-relief modern flood plain, closed depressions, modern drainageways or toeslope positions on the landscape. Sea sonal high water table and potential for

HUDSON and WISCONS IN EPISODE

- Qe-Sand Dunes and Sand Shee ts (Peoria Formation-sand facies). Generally less than 3 m (10 ft) of yell owish brown, massive, calcareous, loamy sand to fine sand. It may overlie yellowish-brown, coarse-grained sand and gravel (Noah Creek Fm.), or it may overlie yellowish to grayish brown, usually cal care ou s, stratified I carm to silt I carm to sandy I carm diamicton (Dows Fm-Morgan Mbr.). Usually restricted to a narrow belt along major river valley bott oms or ad jace ntupl and s on the Des Moine's Lobe. Off the Des Moine's Lobe this unit is not restricted to dunes along valley a reas and may occur as sand stringers overlying unnamed ero sion surface loamy sediments.
- Qdsr Loa my Sedi ments Shallow to Lime stone, Dolomite, and Shale (DeForest-Noah Creek-Shell Rock Formation). 1 to 2 m (3-7 ft) of yellowish brown to gray, massive to weakly stratified, well to poorly sorted, loamy, sandy and silty alluvial sediment that overlies the Upper Devonian be drock surface., usually with a thickness of 12 to 18 m (40-60 ft) and it occurs in the southern part of the quad. It is characterized by fossi liferous carbonates with some shale. Layers with a bun dant subspherical and tabular stromatoporoids, which may be replaced by calcite crystal masses, commonly occur in the lower part of the formation. Around south west part of the quad, this formation is do minated by argillac eous do lo mite and do lo mite climestone, and the thickness
- Odl gc Lo amy Sediments Shallow to Dolomite, Limestone, and Shale (DeForest-Noah Creek-Lithographic City Formation). 1 to 2 m (3-7 ft) of yel lowish brown to gray, massive to weakly stratified, well to poorly sorted, loamy, sandy and silty alluvial sediment that overlies the Middle to Upper De von ian bedrock surface. This formation is the major top bedrock unit in the quad, the maximum thickness of this map unit is up to 3.3 m (11.0 ft), consisting of do lomite and dolomitic limestone, partially characterized by interbeds of unfossili ferous to sparsely fossil iferous lamin ated lithographic and sub lit ho graph ic li mesto ne and dol omiti c limestone, i npart argilla ceous or with slight shale. "Bird seye" carbo nate fa bii c i scommon in some la yers.

WISCONSIN EPISODE

- Qglhc- Colla pse d la ke se di me nts-hummoc ky i sola te d ic e-wal led la ke pla ins (flat-topped hummoc ks) (Dows Formation-Lake Mills Mbr./ Morgan Mbr.). Generally less than 3 m (10 ft) of dark gray ish brown, massive, calc areous silty clay loam, to silt loam; often overlying a thin basal increment of sand and gravel (<1 m). It overlies yellowish to gravish brown, usually calcareous, stratified loam to salt loam to sandy loam diamicton; textures can be quite variable (Dows Fm - Morgan Mbr.). Isolated i ce-wall ed lake plain's comprise a portion of this land form area. Other hummock's consist of less than 15 m of vel lowish to gra vish brown usually calcareous, stratified loam to silt loam to sandy loam diamicton (Dows Fm-Morgan Mbr.). Overlie's gray, calcareous, massive, dense loam diamict on (Dows Fm.-- Alden Mbr.). Modera te to high relief hummocky land form a rea with 3 to 10 m (10 to 32 ft) of local relief.
- Qnw-Sand and Gravel (Noah Creek Formation). Generally less than 8 m (26 ft), but there may be significantly thin ner coarseg as ined deposits in smaller stream valle ys. Yellowish brown to gray, poorly to well-sorted, massive to well stratified, coarse to fine feldspathic quartz sand, pebbly sand and grave l. In the map area the unit overlies middle Wisconsin-age Sheldon Creek Fm. This unit en compasses out wash deposits that accumulated in valley trains during
- Qnw2 Sand and Grave I (Noah Creek Formation). 2 to 12 m (7 to 40 ft) of yell owish brown to gray, poorly to well sorted, massive to well-stratified, coarse to fine feld spathic quartz sand, peb bly sand and gravel with few intervening layers of silty clay. Along many vall eys a thin mantle of losss, reworked loess, or fine-prained alluvium (Qal) may be present. This unit in cludes silty colluvial deposits derived from the adjacent map units. In places this unit is mant led with 1 to 3 m (3-10 ft) of fine to medium, well-sorted medium to fine sand derived from wind reworking of the all uvium. This unit encompasses deposits that accumulated in low-relief stream valleys during the Wisconsin Episode and Hudson Episode. Seasonal high water table and some potential for
 - Qnw3 Sand and Gravel Shallowto Bedrock (Noah Creek Formation). 1 to 3 m (3-10 ft) of yellowish brown to gray, poorly to well sorted, massive to well stratified coarse to fine feld spathic quartz sand, pebbly sand and gravel. May be overlain by up to 2 m (7 ft) of silty all uvial material. In place s mant led with fine to medium well-sorted feldspathic quartz sandderived from wind reworking of the alluvium. Fractured carbonate bedrock is less than 5 m (16 ft) be low the land surface. The unit encompasses deposits that accumulated in river and stream valleys during the late Wisconsin as well as exhumed Pre-Illi noi s Epi so de deposit s of the Wolf Creek and Albumett formations.
- Qo gfp bm Complex glaciofluvial plain (Noah Creek Formation). Greater than 3 m but less than 15 m (10 ft but less than 48 ft) of grayish brown, often cal care ou s, stratified sand and gravel. Overlies gray, calca reous, massive, dense loam diamicton (Dows Fm-Alden Mbr. or Sheldon Creek Fm) or in several locations Paleozoic limestone bedrock. It may have sand and grave lat the land surface. In depressions and sags it may be buried by several meters of DeForest Fm-Woden Mbr. This area is associated with lows and sags on the Bemis Moraine
- Qoch Vall ey train out wash (Noah Creek Formation). Generall y 3 m (10 ft) to less than 8 m (26 ft) of dark gray, dark grayish brown, dark brown to dark yel lowish brown medium to coarse sand, gravelly sand to pebbly gravel. Overlies gray, calcareous, massive, densel oam diamit on (Dows Fm-Alden Mbr.). n valley positions, it is at the land surface of older terraces. On the modern floodplain it is buried by DeForest Fm alluvium. Low-relief landforms expressed as broad terraces; long, narrow longitud in alternaces or cuspateshaped point terraces. Terraces associated with the major valleys are benched on a gray, cal care ous, massive, dense loam diamicton (Dows Fm.-Alden Mbr.).
- Qochb Valley train outwash shallow to be drock (Noah Creek Formation). Generally 1 to 5 meters (3 to 16 ft) of dark gray, dark gray ish brown, dark brown to dark yell owish brown medium to coarse sand, gravelly sand to pebbly gravel. Overlies gray, calcareous, massive, dense loam diamicton (Dows Fm -Alden Mbr.). It is at the land surface of older terraces. On the mode m flood pla in it is buried by DeForest Fm. alluvium Low-relief landforms expressed as broad terraces; long narrow longitudinal terraces or cuspate shaped point terraces. Unit is benched on Upper Devonian age Shell Rock Format ion limestone, dolomites or shale.
- Ooch(s) Slac kwater deposits overlying valley train outwash (Noah Creek Formation-silt facies). Generally less than 3 meters (10 ft) of dark grayish brown to yel low ish brown, massive to lamin ated, calca reous silt loam. Unit overlies < 5 meters of dark gray, dark grayish brown, dark brown to dark yellowish brown medium to coarse sand, gravelly sand to pe bbly gravel. Low-relief land forms expressed as broad terraces; long, narrow longitudinal terraces or c uspate-shaped point terraces. Unit is benched on a gray, cal careous, massive, dense loam diamicton (Dows Fm-Alden Mbr.). No flooding
- Qof Out wash fan (Noah Creek Formation). Thick ness can be quite variable from 3 to 12 m (10-39 ft) of yell owish-brown coarse-grained sand and gravel. May overlie gray, cal care ous, massive, dense loam diamict on (Dows Fm-Alden Mbr. or Sheldon Creek Fm.). Broad low-relief apron that gently slopes away from the morain e front. Seasonal high water table.
- Qsgc Suprag la cial complex (Dows Formation-Pilot Knob Mbr/ Morgan Mbr.) Greater than 3 m (10 ft) but less than 15 m (49 ft) of yellowish brown, often calcareous and fractured, stratified sand and gravel with interbedded stratified loam diamicton. Collapse features are usually evident. In depressions and sags on upland surfaces, the sand and gravel may be buried by DeForest Fm.-Woden Mbr. Overlies gray, calcareous, massive, dense loam diamicton (Dows Fm. - Alden Mbr.). Moderate to high relief 3 to 8 m (10 to 26 ft) hummocks, beaded ridges, kames and associated linked drainages on upland
- Qtr_gl hc_bm Coll apsed la ke sediments hummo cky iso lated i ce-walled lake pla ins to flat-topped hummocks-- (Dows Formation-Lake Mills Mbr. Mong an Mbr.). Generally less than 3 m (10 ft) of dark grayish brown, massive, calcareous silty clay loam, to silt loam overlying a thin basal increment of sand and gravel (<1 m or 3 ft). It over lies yellow is htograyish brown usually calcareous, stratified loam to salt loam to sandy loam diamicton; texture scan be quite variable (Dows Fm-Morgan Mbr) and Alden Mbr from 9 to 40 mt hick (27 to 128 ft). Isolated i ce-walled lake plains comprise a portion of this landform a rea. Other hummocks consist of less than 4 m of yel lowish to gray ish brown usually calcareous, stratified loam to silt loam to sandy loam diamict on (Dows Fm-Morgan Mbr.). Overlies a gray, calcareous, massive, dense loam diamict on (Dows Fm-- Alden Mbr.). Moderate to high relief hummocky landform area with 3 to 10 m (10 to 32 ft) of local relief. The se features are usually a spociated with the Altamont Moraine Complex butin the map a rea are present on the Be mis Moraine. Seasonalhigh water table.
- Qtr amc Till ridge (Dows Formation-Morgan Mbr. or Pilot Knob Mbr or Lake Mills Mbr.) Generall y9 to 15 meters (27 to 49 ft) of yell owish to grayish brown, usually calca reous and fractured, stratified loam to silt loam, stratified sands and gravels to sandy loam dia mictor; textures can be quite variable. Over lie s gray, cal care ous, massive, dense I oam dia micton (Dows Fm.-Al den Mbr.). The Al den Mbr. in this mapping unit can extend to depths in excess of 5 meters (85 ft). This sediment pack age over lies the Sheldon Creek Formation diamicton. Moderate to high relief hummocky landform features exceed 3 to 10 met ers (10 to 32 ft) of local relie f. This landform is characterized by moderate to highrelief hummocks, esker and kame features and ice-walled lakes; and is a sso ciated with the Altamont Mora ine Complex. Season all high water table.
- Qtb_bm Till be nch (Dows Formation-Morgan Mbr. or Pilot Knob Mbr or Lake Mills Mbr.). Generally 9 to 20 meters (27 to 64 ft) of yell owish to gravish brown, usually calca re ous and fractured, stratified loam to silt loam; stratified sands and gravels to sandy loam dia mictor; texture's can be quite variable. Over lie's gray, cal care ous, massive, dense I cam dia micton (Dows Fm.- Alden Mbr.). The Alden Mbr. in this mapping unit can extend to depths in excess of 20 meters (64 ft). This sediment package overlies the Sheldon Creek Formation diamicton. This bench is cut into the Bernis Moraine surface at an elevation of approximately 1250 feet. Appears to be related to a mel twater release a so ciated with the younger Altamont Mora ine Complex. Moderate relief hummocky land form features exceed 3 to 6 m (10 to 20 ft) of local relief. This land form is characterized by moderate to high relief hummocks and icewalled lakes; and is a sociated with the Altamont Moraine Complex. The surface pattern is primarily rounded circular to rounded irregularly shaped
- Qtr_bm Till ridge (Dows Formation-Morgan Mbr.) Generally 3 to 15 m (10 to 49 ft) of yellowish to grayish brown, usually calcareous and fractured, stratified loam to silt loam; stratified sands and gravels to sandy loam diamictor, textures can be quite variable. Overlies gray, cal careous, massive, dense loam diamic ton (Dows Fm.-Alden Mbr.). The Alden Mbr. in this mapping unit rarely extends to depths greater than 15 meters (49 ft); and overlies the Shel don Creek Formation dia micton. At the DML margin, this landform may be mantled with a thin layer of Peoria Formation silt. Low to moderate relief hummocky landform features exceed 3 to 5 m (1016 ft) of local relief. This land form is associated with the Bemis Moraine. The surface pattern is irre gul ar ly shape d p att em s. Season al high water table.
- Qsc2 Lo amy Sediments Shall ow to Glacial Till (Unnamed erosion surface sediment) 1 to 3 m (3 to 10 ft) of yellowish brown to gray, massive to we ally stratified, well to poorly sorted loamy, sandy and silty erosion surface sediment. Map unit includes some areas mantled with less than 1 m (3 ft) of Peoria Formation (silt or sand facies). Overlies massive, fractured, slightly firm glacial till of the Sheldon Creek Formation.
- Osc Glacial Till (Sheldon Creek Formation-undiff). Generally 3 to 15 m (10 to 49 ft) of a yellowish brown to gray, calcare ous fractured to massive clay loam; at dep th this unit can be variably textured and containsignificant sand and gravel bodies. The upper 3 to 7 m (10 to 20 ft) may be periglacially altered. It is not uncommon to see few Pierre Shale clasts in core samples. The presence of unaltered Sheldon Creek till in the eastern portion of the county is questionable. This unit overlies Pre-Illinois diamict on and is only shown on the cross-section.

PRE-ILLIN OIS EPISODE

Qwa3 - Till (Wolf Creek or Alburnett Formations). Generally, 3 to 23 m (10 to 75 ft) of very dense, massive, fractured, loamy glacial till of the Wolf Creek or Alburnett formations. This mapping unit can be builed by glacial sediments (Sheldon Creek Fm.), unnamederosion surface sediments, loess or alluvium

CRETACEOUS SYSTEM

Kw-Sandstone, Mudstone, and Side rite Pellets (Windrow Formation) "Mid"-Cretace ous This map unit occurs as erosional outliers and is only found occ asionally in well cuttings in the mapping a rea. The formation is characterized by red dish shally sandst one and mudst one or siderite pellets. Its thickness is variable, but usual lyless than 6 m (20 ft). This unit is shown only on the cross-section, not on the map.

PALEOZOIC

DEVONIAN SYSTEM

- Dsr Limestone, Dolomite, and Shale (Shell Rock Formation) Upper Devonian. This map unit usually has a thickness of 12 to 18 m (4060 ft), and occurs in the southern part of the quad. The unit is chara cterized by fossiliferous carbonates, with some grey to light green shale. Layers containing abundant subspherical and tabular stromatoporoids commonly occur in the lower part of the unit. Brachiopods, bry ozo ans, corals, and crinoids are abundant in some in tervals. This unit is shown only on the cross-section, not on the map.
- Dlgc Dolomite, Limestone, and Shale (Lithograph City Formation) Middle to Upper Devonian. This map unit forms the major uppermost bedrock in the quad, with a maximum thickness of up to 33 m (110 ft). This unit consists of dolomite and dolomitic limestone, partially characterized by interbeds of lami nated lithog raphic and sublithog raphic limestone and dolomitic limestone, in part argilla ceous or with slight shale. "Birdse ye" structures are common. Some intervals are fossil if erous and stromatoporoid in ch. This unit is shown only on the cross-section, not on the map.
- Dcv Lime stone and Dolo mite (Coral vil le Formation) Middle Devonian. The thickness of this map unit varies between 10 and 18 m (35-60 ft), and it is do minated by limestone, do lo mitic li mestone, and dolo mite, in part laminated and argilla ceous. Brach jop od s and coral's usu ally oc cur in the li mestone facie's This unit is shown only on the cross-section, not on the map.
- Dic Do lo mite and Li mestone (Little Cedar Formation) Middle Devonian. The thickness of this map unit ranges from 27 to 35 m (90 115 ft) in the county. The unit is do minated by slightly argillaceous to a rgill accous dolomite and dolomit ic limestone, usually vuggy and partially lamin ated and/or cherty. This unit is commonly fossiliferous, and brachi opods are especially abundant in the lower portion. This unit is shown only on the cross-section, not on the map.

Qpq - Pits and Quarries

▲ Dr ill Holes Water Well Logs

> Base map from iowa DOT Road Map Layers 2009. Shaded relief from Iowa Lidar Project 2007-2011. lowa Geological and Water Survey digital cartographic file Worth_Co_SurficialGeology2012.mxd, version 9/28/12 (ArcGIS 10.0) Map projection and coordinate system based on Universal Transverse Mercator (UTM) Zone 15, datum NAD83.

The map and cross section are based on interpretations of the best available information at the time of mapping. Map interpretations are not a substitute for detailed site specific studies.

SURFICIAL GEOLOGY OF WORTH COUNTY, IOWA

Iowa Geological and Water Survey **Open File Map OFM-12-02** September 2012

prepared by

Deborah Quade, Stephanie Tassier-Surine, Huaibao Liu, Robert McKay, Robert Rowden and James Giglierano

Iowa Geological and Water Survey, Iowa City, Iowa



Iowa Department of Natural Resources, Chuck Gipp, Director Iowa Geological and Water Survey, Robert D. Libra, State Geologist

Supported in part by the U.S. Geological Survey Cooperative Agreement Number G11AC20247 National Cooperative Geologic Mapping Program (STATEMAP)

ACKNOWLEDGMENTS

Special thanks to Stan Bachaus, Ken Benjegerdes, Leon Benjegerdes, Bill Blanchard, Dan Brunsvold, Dianna Brunsvold, Marguerite Brunsvold, Van Butler, Joyce Calgaard, Palmer Dahl, James Dobel, Patty Edenfield, Mark Fredricksen, Karen Gordon, Tim and Melinda Holden, Steve Jennings, Amil Johnson, LeRoy Luecht, Gary Marken, David Meyer, Arden Midtgaard, Arlen Midtgaard, Arlow Rugland, Don Severson, and Brian Tweeten for allowing us access to their properties; and to 3 H Account LLC., BMC Aggregates, L.R. Falk Construction Co., Falkstone LLC., Holcim Cement Inc., Lehigh Cement Company, Randall Ready Mix LLC., Trenhaile & Sons Inc., Ulland Brothers Inc., and Weaver Construction Co. for allowing us to work in their quarries. New subsurface geologic data was mostly generated by the University of Iowa (UI) student Kyle Bracken who produced descriptive logs of water well drilling samples. Michael Bounk of the Iowa Geological and Water Survey (IGWS) provided additional descriptive logging of water wells. Jason Vogelgesang (IGWS) prepared well samples for stratigraphic logging. Dr. Art Bettis (UI) provided valued information concerning the surficial geology. Natural Resources Conservation Service staff Robin Wisner, Kathy Woida, Leland Camp and Neil Sass assisted with field work and describing cores as part of the mapping update for the Iowa Cooperative Soil Survey update for Worth County. Especially, Dean Mueller, director of the Worth County Conservation Board who granted permission for numerous drilling locations on public lands. Shane Buenzow, Worth County Sanitarian, assisted with acquiring well drillers records. Also, a very special thanks to nearby landowner, Don Mathahs, who provided a vehicle tow out from Elk Creek Valley. In addition, Mary Howes and Casey Kohrt (IGWS) provided GIS mapping technical help, and Paul VanDorpe (IGWS) provided proof reading.

Introduction to the Surficial Geology of Worth County, Iowa

Worth County lies in north-central Iowa and straddles the eastern margin of the Late Wisconsin-age Des Moines Lobe Landform (DML), the most recently glaciated region of the state and the Wisconsin-age Iowan Erosion Surface (IES) Landform Region (Prior and Kohrt, 2006). Generally speaking, in the western portion of the map area, the DML consists of a complex suite of depositional landforms and sediment sequences related to supraglacial, subglacial, and proglacial sedimentation. Quaternary deposits in this region can reach a maximum thickness of 59 m (195 ft). The Des Moines Lobe is characterized by hummocky terrain that forms arcuate belts of moraine complexes and undulating plains with thick increments of supraglacial sediment (>3 m). In the map area, the most noteable features of the DML are the Bemis Moraine, the terminal moraine of the DML which is dated at approximately 14,500 to 14,000 RCYBP (radiocarbon years before present), and the slightly younger Altamont Moraine Complex which is dated at approximately 13,500 RCYBP. Supraglacial and proglacial sediments (coarse-grained glaciofluvial, ice-contact sediments associated with hummocky terrain, outwash fans, and channel deposits) encompass a large area of the eastern flank of the DML and are extensively mapped at the former ice margins and in the Shell Rock River, Elk Creek and Winnebago River valleys. The eastern portion of the map area is dominated by unnamed loamy sediments (IES materials) of variable thickness overlying Wisconsin-age Sheldon Creek Fm. glacial sediments, Pre-Illinoian glacial sediments or shallow rock. These deposits are regionally extensive and in this county, are on average less than 15 m (50 ft) in thickness. Significant areas of bedrock outcrop or areas with less than 5 m (15 ft) of loamy material over rock are present, especially along the Shell Rock River and Winnebago River valleys. Bisecting the eastern portion of the map area is the Shell Rock River valley. The valley is younger than the Bemis Moraine and was cut approximately 12,500 RCYBP during the catastrophic drainage of the younger Algona Moraine. The valley is dominated by a variable thickness of coarse-grained outwash associated with the "last-gasp" drainage of the DML.

Williams (1899) described and mapped the Quaternary and Paleozoic bedrock geology of the county and discussed the stratigraphy of Devonian strata that were exposed at the land surface. He also noted the extreme thinness of the "drift" along the Winnebago River and nearby Shell Rock River and the remarkable difference in surface features between the eastern and western portions of Worth County. Statewide bedrock geologic maps by Hershey (1969), and most recently by Witzke, Anderson, and Pope (2010), depict the increased understanding of the complex distribution of geologic units at the bedrock surface across this region, including Worth County. The only regional surficial map of the area consists of the Des Moines 4° x 6° Quadrangle at a scale of

The map area has a rich and complex geologic history punctuated by at least seven periods of glaciation between 2.2 million to 500,000 years ago (Boellstorff, 1978a, b; Hallberg, 1980, 1986). In this area, Pre-Illinois Episode glacial deposits and associated buried soils are overlain by much younger Wisconsin-age glacial deposits. During earlier and mid Wisconsin-age, ice advances dating from approximately 40,000 to 26,000 years before present were deposited throughout the map area. In Iowa, this glacial deposit is formally recognized as the Sheldon Creek Formation (Bettis et al., 1996, Bettis, 1997) and in earlier literature is referred to as the "Tazewell till" (Ruhe, 1950). A period of intense cold occurred during the Wisconsin full glacial episode from 21,000 to 16,500 years ago (Bettis, 1989). This cold episode and ensuing upland erosion led to the development of the distinctive landform recognized as the IES (Prior, 1976). A periglacial environment prevailed during this period with intensive freeze-thaw action, solifluction, strong winds, and a host of other periglacial processes (Walters, 1996). The result was that surface soils were removed from the IES, and the Sheldon Creek and Pre-Illinois till surface was significantly eroded, resulting in the development of a region-wide colluvial lag deposit referred to as a "stone line." Following the IES formation, the southern edge of the Laurentide Ice Sheet split into several lobes that each flowed down regional topographic lows. The Des Moines Lobe extended from central Canada through the Dakotas and Minnesota into Iowa, terminating at what is now the City of Des Moines. The Des Moines Lobe was active in Iowa between about 15,000 and 12,000 RCYBP, about 5,000 to 8,000 years later than glacial lobes to the east made their southernmost maximum advance (Johnson, 1986; Fullerton, 1986). The Lobe advance occurred well into a period of regional warming and was thus climatically out of equilibrium (Kemmis et al., 1994). Ice thickness reconstructions indicate that the lobe was probably thin and gently sloping (Mathews, 1974; Clark, 1992; Brevik, 2000; Hooyer and Iverson, 2002). Clark (1992) reconstructed the Lobe's thickness near Ames, Iowa, at ~80 m (262 ft). More recently, ice reconstructions by Hooyer and Iverson (2000) were based on a model assuming the Bemis Moraine was ice-cored, which yielded ice thickness estimates of ~250 m (820 ft). Despite these variations, all agree that the Des Moines Lobe ice sheet was extremely thin and gently sloping. This ice advance was rapid and episodic, and was most likely fueled by basal lubrication; in other words, a warm-based, non-deforming bed glacier. These assumptions are backed up by evidence of numerous plants (Baker et al., 1986) and trees (Bettis et al., 1996) found near the base of the DML package. Furthermore, the complex landform sediment assemblages found on the DML in Iowa seem more indicative and explained by regional stagnation caused by a surging-type glacier, not rapid recession.

Surficial deposits of the map area are composed of seven formations: DeForest, Dows, Noah Creek, Peoria, Sheldon Creek, Wolf Creek, and Alburnett formations, as well as unnamed erosion surface sediments. Hudson age deposits associated with finegrained alluvial, organic, and colluvial sediments include the DeForest Formation which is subdivided into the Camp Creek, Roberts Creek, Gunder, Corrington, Flack, and Woden members. The Dows Formation consists of upland glacial deposits and is subdivided into the Alden, Lake Mills, Morgan, and Pilot Knob members. The Noah Creek Formation includes coarse sand and gravel associated with outwash from the Des Moines Lobe. The Noah Creek Formation includes coarse to finer grained fluvial deposits associated with local stream and river valleys. Unnamed erosion surface sediments consist of reworked till and slopewash deposits associated with periglacial activity during the Wisconsin ice advance. Areas of Peoria Formation eolian materials are present along the the Shell Rock River valley as well as stringers on the IES surface. Eolian materials may also be intermittently present mantling most other mapping units, and are more abundant near stream valleys. Sheldon Creek Formation glacial deposits are undifferentiated and occur in northwest and north-central Iowa. The full extent of these deposits is still not fully understood. Pre-Illinoian glacial deposits in Iowa consist of two formations: the younger Wolf Creek Formation and the Alburnett Formation. The Wolf Creek is divided into the Winthrop, Aurora, and Hickory Hills members (oldest to youngest). The Alburnett Formation consists of several "undifferentiated"

Two bedrock mapping units (Devonian Shell Rock and Lithograph City) are exposed as outcrop in the map area. Bedrock outcrops occur primarily along the Shell Rock River and to some extent along Elk Creek and the Winnebago River. The Devonian rocks are dominated by carbonates varying between limestone and dolomite, accompanied with minor shale. Based on lithologic features and fossils, uppermost bedrock of Devonian in the mapping area can be subdivided into, in descending order, the Shell Rock, Lithograph City, and Coralville formations (Witzke et al., 2010).

Baker, R.G., Rhodes, R.S. II, Schwert, D.P., Ashworth, A.C., Frest, T.J., Hallberg, G.R., and Janssens, J.A., 1986, A full-glacial biota from southwestern Iowa, USA: Journal of Quaternary Science, v.1, p. 91-107. Bettis, E.A., III, 1989, Late Quaternary history of the Iowa River Valley in the Coralville Lake area in Plocher, O.W., Geologic Reconnaissance of the

- Coralville Lake area: Geological Society of Iowa Guidebook 51, p. 93-100. Bettis, E. A. III, Quade, D.J., and Kemmis, T.J., 1996, in Hogs, bogs, and logs: Quaternary deposits and environmental geology of the Des Moines lobe. Edited by: E.A. Bettis, III, D.J Quade, and T.J Kemmis. Iowa Department of Natural Resources-Geological Survey Bureau, Guidebook Series No. Bettis, E.A. III, 1997, Late-Middle and Early-Late Wisconsin Glaciation in North Central Iowa. Geological Society of America North-Central Section
- Meeting Abstracts with Programs 29 (4):5. Brevik, E., 2000. Limits of ice thickness in Iowa during the Late Wisconsian: Journal of the Iowa Academy of Science, v.107, p. 46-50. Boellstorff, J., 1978a, North American Pleistocene Stages reconsidered in light of probable Pliocene-Pleistocene continental glaciation: Science, v. 202 Boellstorff, J., 1978b, Chronology of some late Cenozoic deposits from the central United States and the ice ages: Transactions of the Nebraska
- Academy of Science, v. 6, p. 35-49. Clark, P.U., 1992, Surface form of the southern Laurentide ice sheet and its implications to ice-sheet dynamics. Geological Society of America Bulletin, Fullerton, D.S., 1986, Stratigraphy and correlation of glacial deposits from Indiana to New York and New Jersey, in Richmond, G.M. and Fullerton, D.S.,
- eds., Quaternary glaciations in the United States of America, in Sibrava, V., Bowen, D.Q., and Richmond, G.M., eds., Quaternary Glaciations in the Northern Hemisphere: Quaternary Science Reviews, v. 5, p 23-37.
- Hallberg, G.R., 1980, Pleis tocene stratigraphy in east-central Iowa: Iowa Geological Survey Technical Information Series 10, 168 p. Hallberg, G.R., 1986, Pre-Wisconsin glacial stratigraphy of the central plains region in Iowa, Nebraska, Kansas, and Missouri: in Richmond, G.M. and Fullerton, D.S., eds., Quaternary glaciations in the United States of America, Report of the International Correlation Programme-Project 24: in Sibrava, V., Bowen, D.Q., and Richmond, G.M., eds., Quaternary Science Reviews, Quaternary Glaciations in the Northern Hemisphere, v. 5, p. 11-
- Hallberg, G.R., Lineback, J.A., Mickelson, D.M., Knox, J.C., Goebel, J.E., Hobbs, H.C., Whitfield, J.W., Ward, R.A., Boellstorf, J.D., and Swinehart, J.B., 1991, Quaternary geologic map of the Des Moines 4° x 6° quadrangle, United States: U.S. Geological Survey, Miscellaneous Investigations Series, Map I-1420, 1:100,000 scale map sheet.
- Hershey, H. G., 1969, Geologic map of Iowa, Iowa Geological Survey, scale 1:500,000 (out of print). Hooyer, T.S. and Iverson, N.R., 2000, Clast-fabric development in a shearing granular material: Implications for subglacial till and fault gouge. Geological. Society of America Bulletin, 112(5), 683-692. Hoo yer, T.S. and Iverson, N.R., 2002, Flow mechanism of the Des Moines lobe of the Laurentide ice sheet, Journal of Glaciology., 48(163), 575-586.
- Johnson, W.H., 1986, Stratigraphy and correlation of the glacial deposits of the Lake Michigan lobe prior to 14 ka BP, in Richmond, G.M. and Fullerton, D.S., eds., Quaternary Glaciations in the United States of America, in Sibrava, V., Bowen, D.O., and Richmond, G.M., eds., Quaternary Glaciations of the Northern Hemisphere: Quaternary Science Reviews, v. 5, p.17-22. Kemmis, T.J, Bettis, E.A. III, and Quade, D.J., 1994, The Des Moines lobe in Iowa: a surging Wisconsinan glacier: American Quaternary Association, Program and Abstracts, 13th Biennial Meeting,
- Mathews, W.H., 1974, Surface profile of the Laurentide ice sheet in its marginal areas: Journal of Glaciology, v. 13, p. 37-43. Minneapolis, Minnesota. Prior, J.C., 1976, Landforms of Iowa: Iowa City, Iowa. University of Iowa Press, 154 p.
- Prior, J.C. and Korht, C.J., 2006, The Landform Regions of Iowa, Iowa Geological Survey, digital map, available on IDNR GIS Libraryhttp://www.igsb.uiowa.edu/nrgislibx/
- Ruhe, R.V., 1950, Reclassification and correlation of the glacial drifts of northwestern Iowa and adjacent areas: University of Iowa Department of
- Geology, Iowa City, unpublished Ph.D. thesis, 124 p. Walters, J.C., 1996, General and environmental geology of the Cedar Falls/Waterloo Area, The Iowan Surface, in General and environmental geology of Cedar Falls/Waterloo and surrounding area, Northeast Iowa, Iowa Geological Survey Guidebook Series No. 22, p. 7-9.
- Williams, I. A., 1899, Geology of Worth County: Iow a Geological Survey, Annual Report, v. 10, p. 315-378. Witzke, B.J., Anderson, R.R. and Pope, J.P., 2010, Bedrock Geologic Map of Iowa, scale: 1:500,000, Iowa Geological and Water Survey, Open File Digital Map OFM-10-1.