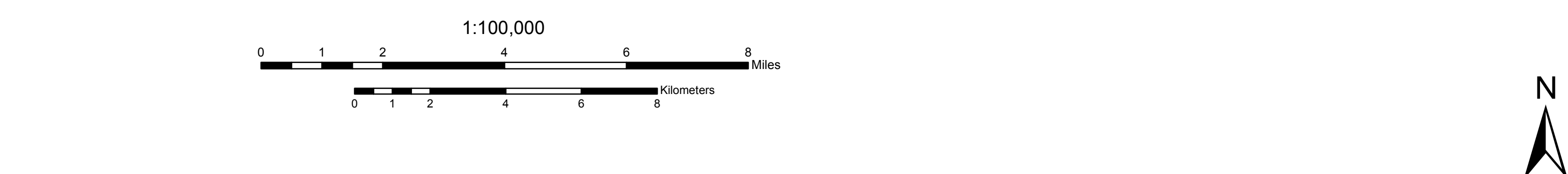
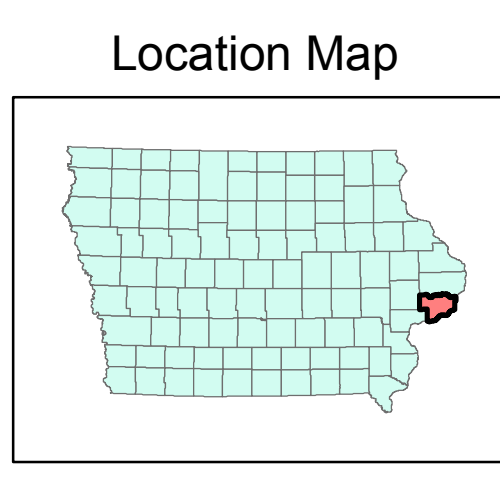
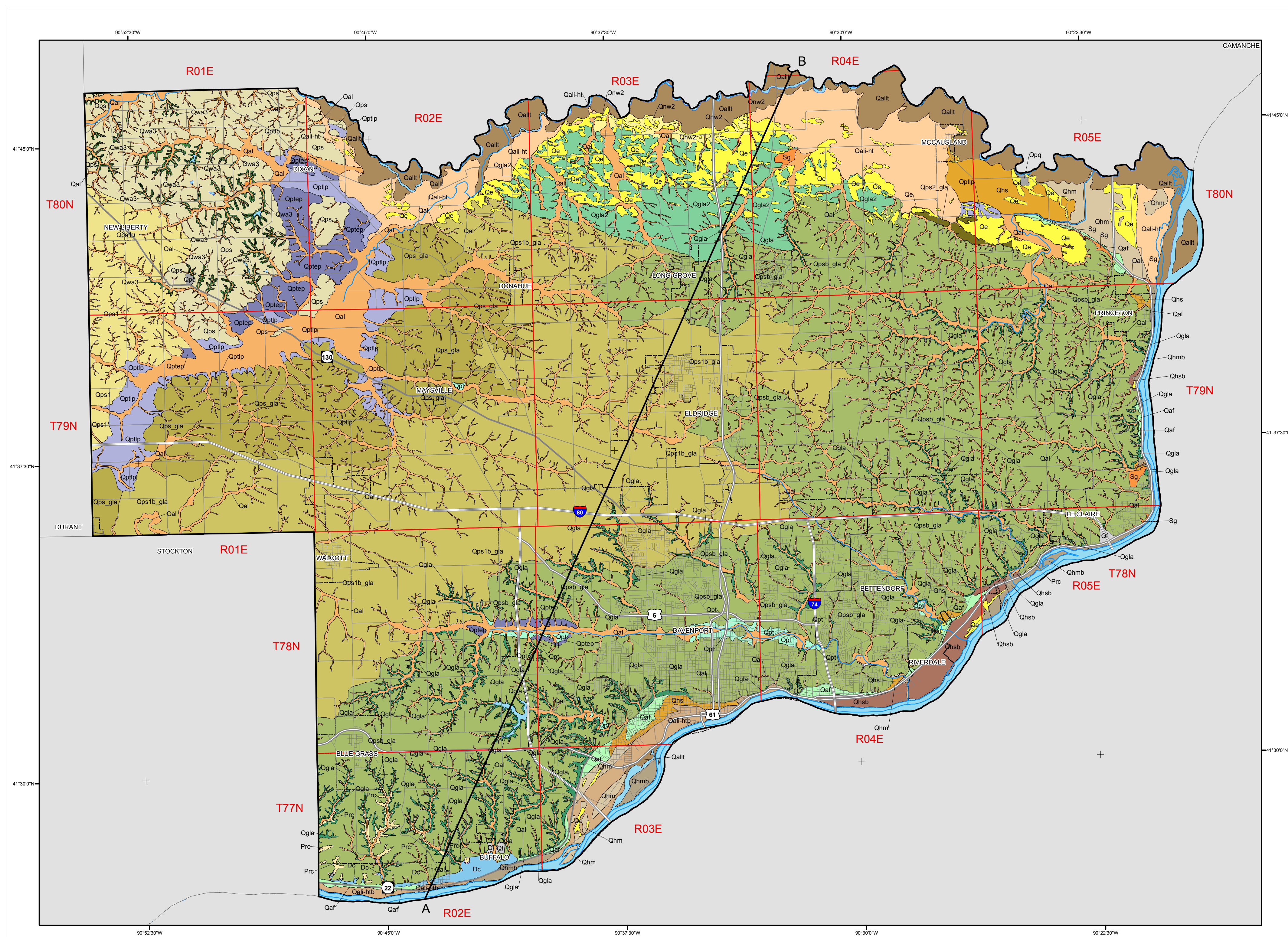
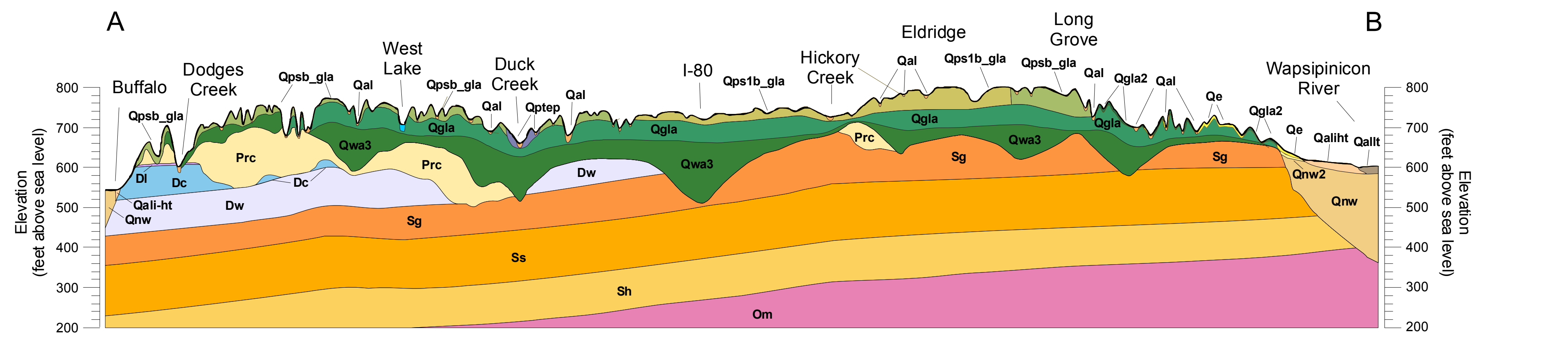


Surficial Geology of Scott County, Iowa



GEOLOGIC CROSS-SECTION A-B



LEGEND

CENOZOIC

QUATERNARY SYSTEM

HURON EPODE

Qal Qal-h
Qal-hb
Qal

Qnw2

Qe
Qhm
Qhmb
Qhs
Qhsb
Qpt
Qltp
Qltpn

Qps
Qps1
Qps-gla
Qpsb-gla
Qps1b-gla
Qps2-gla
Qgla2

Qgla

Qwa3

Prc

DI
Dc
Dw

Sg
Ss
Sh

Om

Qr

W

PALEOZOIC

PALEOZOIC BRACHIAN

PENNSYLVANIAN SYSTEM

DEVONIAN SYSTEM

SILURIAN SYSTEM

ORDOVICIAN SYSTEM

Introduction to the Surficial Geology of Scott County, Iowa

Scott County lies within a varied and unique surficial geologic region. Geomorphically, it contains the boundary between the Iowa Erosion Surface (IES) to the north and the Southern Iowa Drift Plain to the south (Prior and Kohrt, 2006). Stratigraphically, it contains Illinoian and glacial deposits, which are present only in a small area of southeastern Iowa. Surficial materials consist of a mix of thick colluvial deposits (loess), loamy sediments associated with the IES, glacial till (both Illinoian and pre-Illinoian), numerous sand and gravel deposits that may be very thick, terraces associated with both the Mississippi and Wapsipicon rivers and several areas of shallow rock. Multiple periods of Quaternary glacial and subglacial erosion have led to the landscape we see today. In general, the map area consists of loess and colluvial sand of variable thickness overlying Illinoian and Pre-Illinoian glacial sediments.

The surficial geology mapping has been completed as part of the STATEMAP program in Scott County (Quade et al., 2004 & 2005; Tassier-Surine et al., 2005, 2009 and 2009a) and to the west in adjacent County (Quade et al., 2008). The only other regional surficial map of the area consists of the Des Moines 4° x 6° Quadrangle at a scale of 1:50,000 (Hallberg et al., 1993). Norton (1899) first described and mapped the Quaternary and Paleozoic bedrock geology of the county, and discussed the stratigraphy of the Pennsylvanian, Devonian and Silurian strata that comprise the county's bedrock units. Statewide bedrock geology maps by Henshaw (1949) and most recently by Witzke, Anderson, and Pope (2010), depict the increased understanding of the distribution of geologic units at the bedrock surface across this region.

At least seven episodes of Pre-Illinoian glaciation occurred in this region between approximately 2.2 and 0.5 million years ago (Hedffoff, 1986a; Hallberg, 1980, 1986). Episodic erosion during the last 500,000 years has led to the destruction of pre-existing glacial landforms associated with Pre-Illinoian glaciation. In the central Iowa, Hallberg formally classified the units into two formations on the basis of differences in clay mineralogy: the Albion Formation (several unfederated members) and the younger Wolf Creek Formation (including the Whitport, Aurora and Hickory Hills members). Both formations are composed predominantly of fill deposits, but other materials are present. Paleosols are present in the upper part of these till units.

A limited area of southeastern Iowa was glaciated during the Illinoian Episode, around 300,000-130,000 years ago. Most of Scott County, with the exception of the area northwest of Mad Creek, was glaciated during the time. Leventz (1898, 1899) was the first to study the Illinoian glacial deposits in southeast Iowa. The Illinoian till was deposited by the advancing Lake Michigan lobe which moved across western Illinois into Iowa from the northeast (Leventz, 1899; Wickham, 1990). The Lake Michigan lobe incorporated Paleozoic bedrock materials from the Lake Michigan Basin which are distinguished by both the clay mineralogy of the matrix as well as the pebbles and clasts (Lindbeck, 1986; Wickham, 2000). Hallberg (1980) defined the formal stratigraphic nomenclature in Iowa. The only Illinoian Episode till present in Iowa is the Grand Tower Kelleysville Till Member (William and Frye, 1970).

Following Pre-Illinoian glaciation, most areas of eastern Iowa underwent extensive erosion and landscape development and erosion until the Wisconsin Episode loess began to be deposited. A period of intense cold occurred during the Wisconsin fall glacial episode from 21,000 to 16,500 years ago (Betts, 1989). A periglacial environment prevailed during this period with intensive freeze-thaw action, soilification, strong winds and a host of other glacial processes (Walters, 1996). This cold interval and ensuing upland erosion led to the development of the distinctive landform recognized as the IES (Prior, 1976).

In eastern Iowa, the highly eroded and dissected Illinoian and Pre-Illinoian upland and older terraces are mantled by Wisconsin loesses of variable thickness (Ruhe, 1969; Prior, 1976). These sediments are the youngest regionally extensive Quaternary materials and were deposited between 30,000 and 2,000 years ago. Two loess units were deposited across eastern Iowa, the older Pisgah Formation and the younger Pointa Pén Formation. The Pisgah is thin and includes loess and related slope sediments that have been altered by colluvial hillslope processes, pedogenic and periglacial processes. The unit is characterized by the presence of a weakly developed soil recognized as the Farnside Gossol. It is not uncommon to see the Farnside developed throughout the Pisgah and into the underlying older Sangamon Gossol. Most likely the Pisgah loess was deposited on the eastern Iowa landscape from 30,000 to 24,000 years ago (Betts, 1989). The Pisgah Formation is typically buried by Pointa Pén Formation loess. The Pointa Pén Formation loess accumulated on stable landforms in eastern Iowa from 25,000 to 21,000 years ago. The Pisgah Formation is in the same stratigraphic position as the Roxana Silt which is mapped in Illinois. The Pointa Pén Formation consists of all and sand facies.

Ewing (1984) conducted a regional study of extensive post-Illinoian alluvial deposits that had accumulated in major valleys in eastern Iowa. Three major terrace assemblages with differing stratigraphy and age were identified: Early Pisgah Terrace (EPHT), Late Pisgah Terrace (LPHT) and Low Terrace (LT). EPHT deposits are characterized by the presence of Pointa Pén and Pisgah Formation sediments overlying Sangamon Gossol in the underlying alluvium. Ewing theorized that these terraces are older than 40,000 years B.P. but younger than the Illinoian sediments in eastern Iowa (Betts, 1989). LPHT deposits are characterized by the presence of Pointa Pén Formation loess grading down into underlying alluvium with no paleosols. These terraces developed prior to 25,000 years ago and were buried by loess before 12,500 years ago. The LPHT terraces are typically inset into EPHT deposits. The LT is the youngest terrace and is not buried by Pointa Pén loess. Terrace deposits are present primarily along Mad Creek and Duck Creek in the map area.

Surficial deposits of Scott County represent seven formations: Df, Post, Noah Creek (Henry Fm. of Illinois), Pointa Pén, Albion, Grand Tower, Wolf Creek, and Albion formations, as well as an unnamed erosion surface sediment. Hudson-age deposits associated with fine-grained alluvial and colluvial sediments include the DeForest Formation which is subdivided into the Camp Creek, Roberts Creek, Gander and Corriam members. The Noah Creek Formation includes coarse sand and gravel located along the Wapsipicon River associated with the development of the Iowa Erosion Surface. Unnamed erosion surface sediments consist of reworked till and slopewash deposits formed by periglacial activity during the Wisconsin maximum and may be up to 25' thick. Pointa Pén Formation silt materials consist of fine sand and colluvial sand. Pointa Pén materials are present throughout the majority of Scott County except on the east south of the Wapsipicon River mapped as IES materials. Pisgah Formation loess is also present, being thicker near the Mississippi River and thinning to the west. The characteristics of the erosion deposits vary throughout Scott County. The southeast portion of the county adjacent to the Mississippi River is dominated by thick (up to 60') loess with almost no colluvial sand. An adjacent band comprises the center of the county and is dominated by interbedded silt and fine sand. Areas adjacent to Mad Creek are dominated by loess up to 20' thick. A buried sand sheet is present in the northwest part of the map area.

Correlation of Map Units

| Map Unit | Alluvial and Outwash Deposits | Eolian Deposits | Erosion Surface | Glacial Deposits | Bedrock |
|-----------|-------------------------------|-----------------|-----------------|------------------|---------|
| Qal | Qal | Qe | Qhs | Qgla | Qr |
| Qal-h | Qal-h | Qhm | Qhsb | Qgla2 | Qr |
| Qal-hb | Qal-hb | Qhmb | Qhs | Qgla | Qr |
| Qal | Qal | Qhm | Qhs | Qgla | Qr |
| Qnw2 | Qnw2 | Qe | Qhs | Qgla | Qr |
| Qe | Qe | Qhm | Qhs | Qgla | Qr |
| Qhm | Qhm | Qhmb | Qhs | Qgla | Qr |
| Qhmb | Qhmb | Qhs | Qhs | Qgla | Qr |
| Qhs | Qhs | Qhs | Qhs | Qgla | Qr |
| Qhsb | Qhsb | Qhs | Qhs | Qgla | Qr |
| Qpt | Qpt | Qhs | Qhs | Qgla | Qr |
| Qltp | Qltp | Qhs | Qhs | Qgla | Qr |
| Qltpn | Qltpn | Qhs | Qhs | Qgla | Qr |
| Qps | Qps | Qhs | Qhs | Qgla | Qr |
| Qps1 | Qps1 | Qhs | Qhs | Qgla | Qr |
| Qps-gla | Qps-gla | Qhs | Qhs | Qgla | Qr |
| Qpsb-gla | Qpsb-gla | Qhs | Qhs | Qgla | Qr |
| Qps1b-gla | Qps1b-gla | Qhs | Qhs | Qgla | Qr |
| Qps2-gla | Qps2-gla | Qhs | Qhs | Qgla | Qr |
| Qgla2 | Qgla2 | Qhs | Qhs | Qgla | Qr |
| Qgla | Qgla | Qhs | Qhs | Qgla | Qr |
| Qwa3 | Qwa3 | Qhs | Qhs | Qgla | Qr |
| Prc | Prc | Qhs | Qhs | Qgla | Qr |
| DI | DI | Qhs | Qhs | Qgla | Qr |
| Dc | Dc | Qhs | Qhs | Qgla | Qr |
| Dw | Dw | Qhs | Qhs | Qgla | Qr |
| Sg | Sg | Qhs | Qhs | Qgla | Qr |
| Ss | Ss | Qhs | Qhs | Qgla | Qr |
| Sh | Sh | Qhs | Qhs | Qgla | Qr |
| Om | Om | Qhs | Qhs | Qgla | Qr |
| Qr | Qr | Qhs | Qhs | Qgla | Qr |
| W | W | Qhs | Qhs | Qgla | Qr |

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SURFICIAL GEOLOGY OF SCOTT COUNTY, IOWA

IOWA GEOLOGICAL AND WATER SURVEY
Open File Map OFM-11-48
September 2011

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